

Representing effects of aqueous-phase reactions in shallow cumuli in global models

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Beyond effects on dynamics, moist convection also affects atmospheric chemistry

- Vertical transport of chemical species (including scavenging)
- Turbulent mixing of chemical species
- Photochemistry by changing the radiation field
- Lightning production of NO_x
- Aqueous phase reactions

Aqueous phase oxidation of SO₂ accounts for a majority of sulfate production

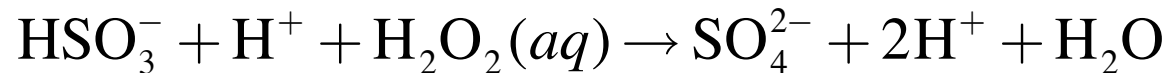
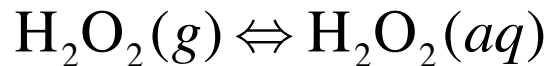
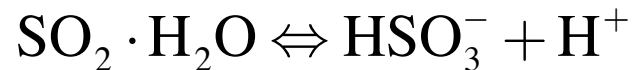
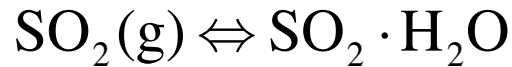
Table 2. GEOS-CHEM Global Budgets for Sulfate Produced by Different Oxidation Pathways

	Source, Tg S yr ⁻¹
Total sulfate	31.0
SO ₂ + OH (gas phase)	8.2
S(IV) + H ₂ O ₂ (in cloud)	15.7
S(IV) + O ₃ (in cloud)	2.3
S(IV) + O ₃ (fine sea salt)	0.4
S(IV) + O ₃ (coarse sea salt)	2.3
Primary anthropogenic	2.0

Alexander et al., 2005, JGR

It also increases the scattering efficiency of sulfate aerosols (Lelieveld and Heintzenberg, Science, 1992)

The reactions



Seinfeld and Pandis, 2006

Sulfate production rate in mixing ratio relative to air:

$$\frac{d[\text{SO}_4^{2-}]}{dt} = k [\text{SO}_2]_{\text{g}} [\text{H}_2\text{O}_2]_{\text{g}} q_{\text{c}} p_{\text{air}} \rho_{\text{air}} \quad q_{\text{c}}: \text{cloud water content}$$

$$k \approx 2 \times 10^{-3} \text{ s}^{-1} \text{ ppb}^{-1} (\text{g} / \text{kg})^{-1} (\text{bar})^{-1} (\text{kgm}^{-3})^{-1}$$

Currently, in global chemical transport models (CTM), SO_2 and H_2O_2 are titrated over a CTM time step within the cloudy volume.

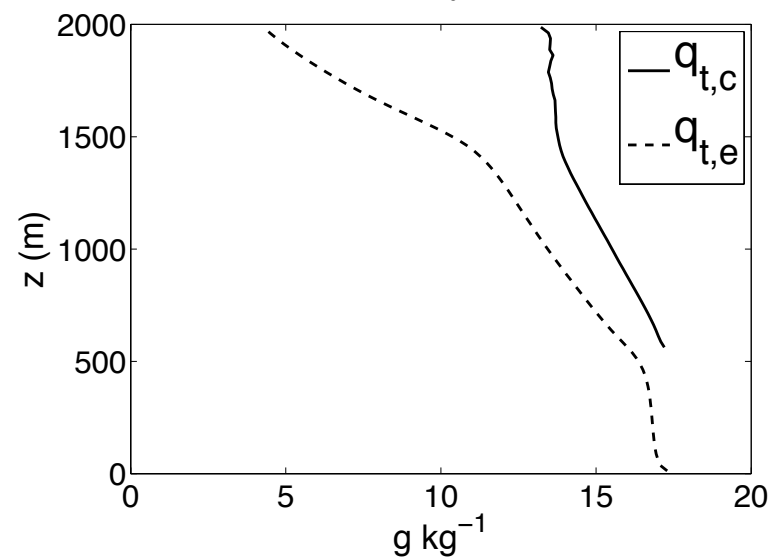
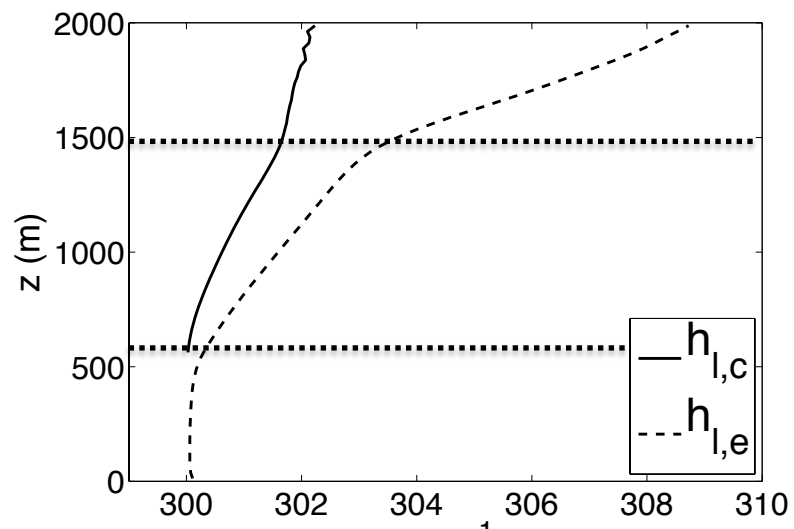
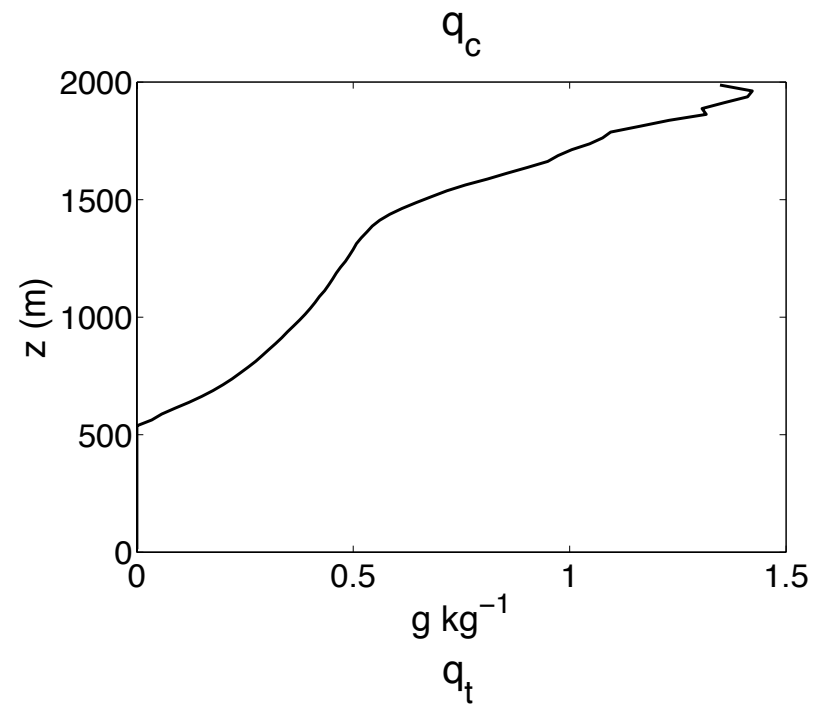
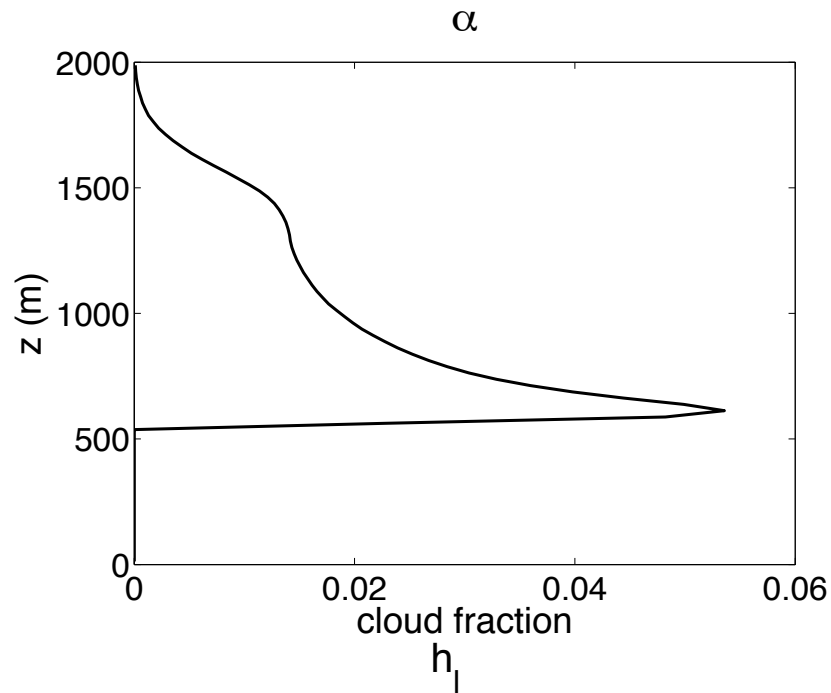
LES of SO₂ oxidation by H₂O₂

- Large-scale meteorological forcing from BOMEX
- 6.4kmX6.4kmX3km with a resolution of 25mX25mX25m

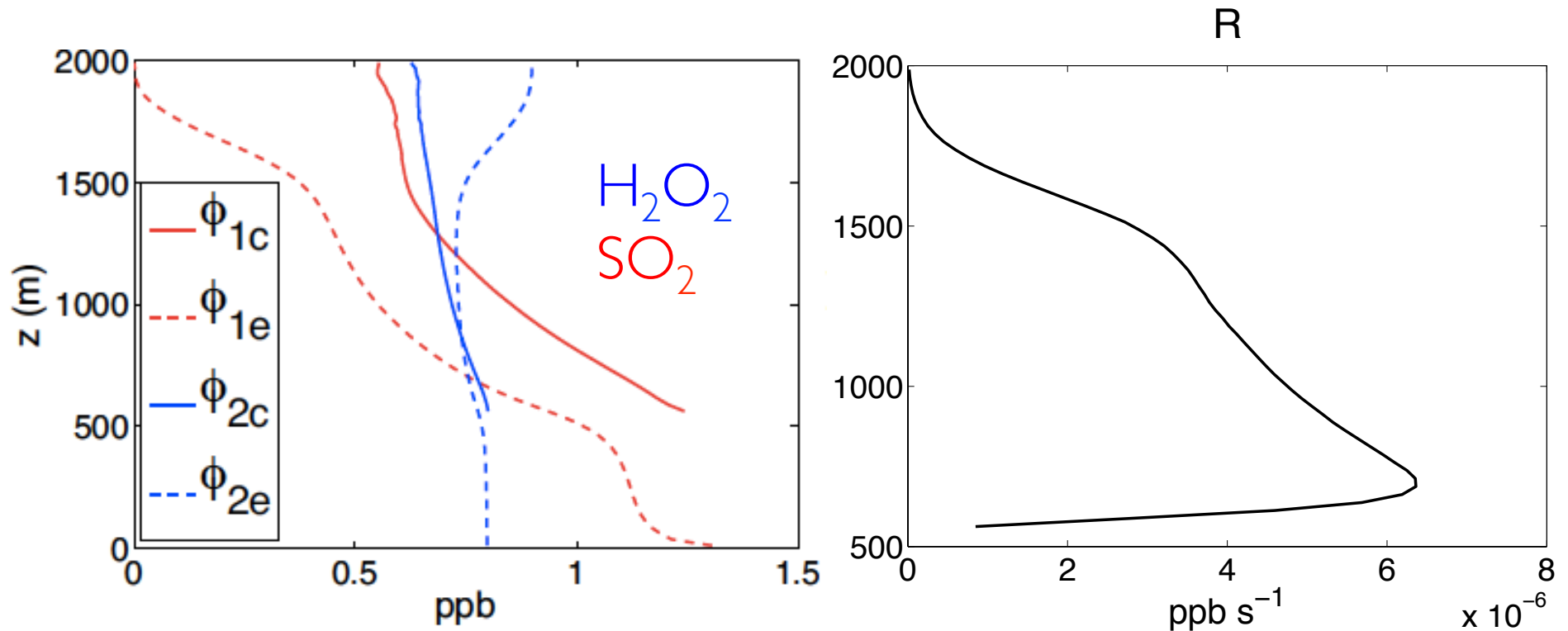
“Chemistry”:

- Tracer 1 (“SO₂”) is released from surface with a fixed flux,
- Tracer 2 (“H₂O₂”) is relaxed to a constant reference profile over 1 day
- The two tracers react to form tracer 3 (“H₂SO₄”) with a specified reaction rate k
$$\frac{d[\text{SO}_4^{2-}]}{dt} = k[\text{SO}_2]_g[\text{H}_2\text{O}_2]_g q_c$$
- An additional sink of tracer 1 with a timescale of one day (mimicking dry deposition and gaseous oxidation).

General characteristics of BOMEX



LES results on SO₂ oxidation



Reference case:

SO₂ flux: $0.024 \text{ ppb kg m}^{-2} \text{ s}^{-1}$

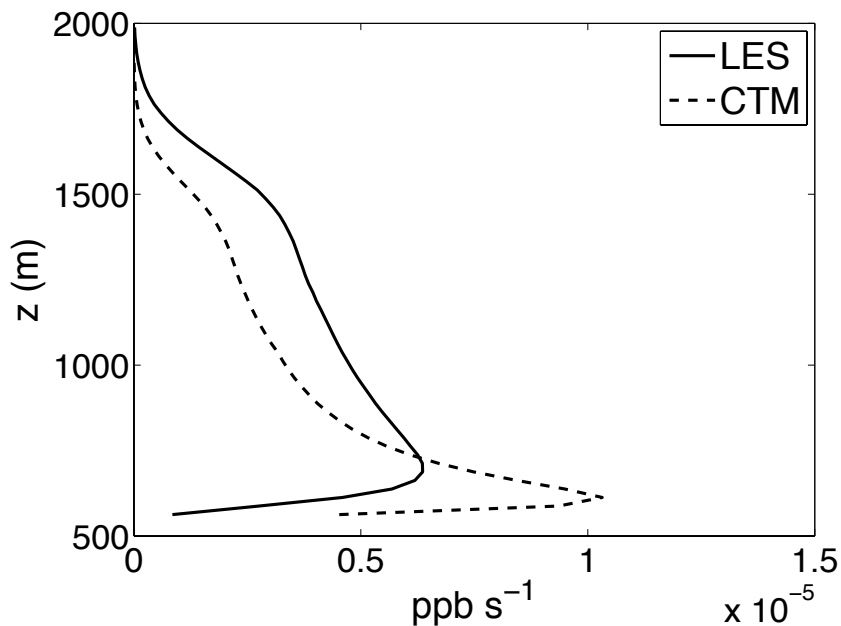
H₂O₂ value: 0.9 ppb

$$k = 1 \times 10^{-3} \text{ s}^{-1} \text{ ppb}^{-1} (\text{g / kg})^{-1}$$

Performance of current chemical transport model treatment

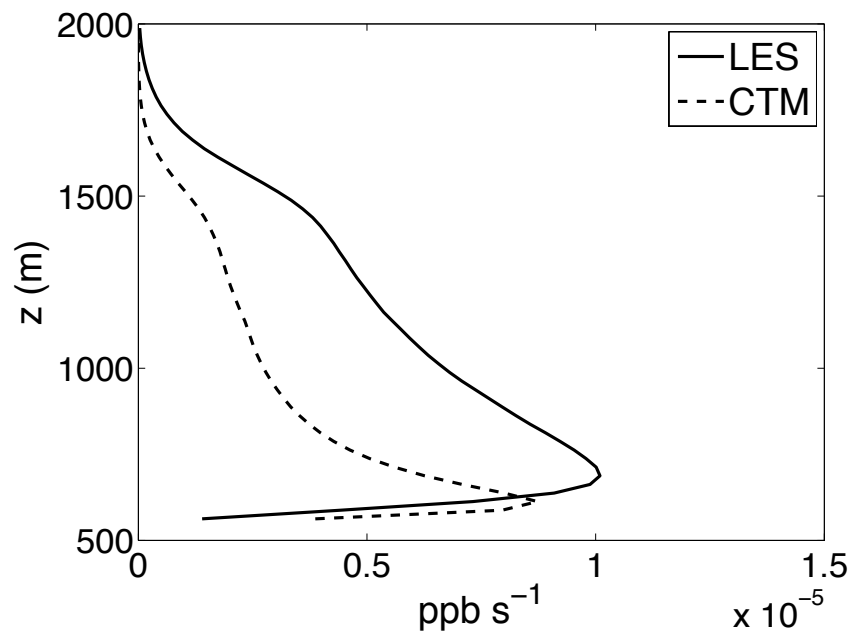
$$k = 1 \times 10^{-3} s^{-1} ppb^{-1} (g/kg)^{-1}$$

R



$$k = 2 \times 10^{-3} s^{-1} ppb^{-1} (g/kg)^{-1}$$

R



Uses LES simulated cloud fraction and horizontal mean tracer values with a 1-hr global model time step.

How well can the eddy diffusivity and mass flux (EDMF) approach represent the aqueous phase reaction ?

Sub-grid scale transport

$$\overline{w'\phi'} \simeq -K \frac{\partial \bar{\phi}}{\partial z} + M(\phi_u - \bar{\phi})$$

Soares et al., QJ, 2004

where K is the eddy diffusivity and M is the updraft mass flux. The mass flux component is modeled with a bulk entraining/detraining plume

Reaction



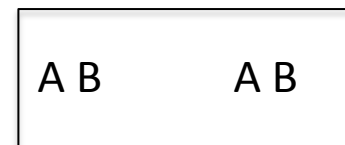
where over-bar indicates averages over cloudy updrafts

Effective entrainment/detrainment rates and eddy diffusivities are diagnosed using conserved tracers.

This approach isolates errors due to chemistry and can be used in a super-parameterized global model.

Potential errors:

1. Effective entrainment and detrainment rates are tracer-dependent
2. Segregation error



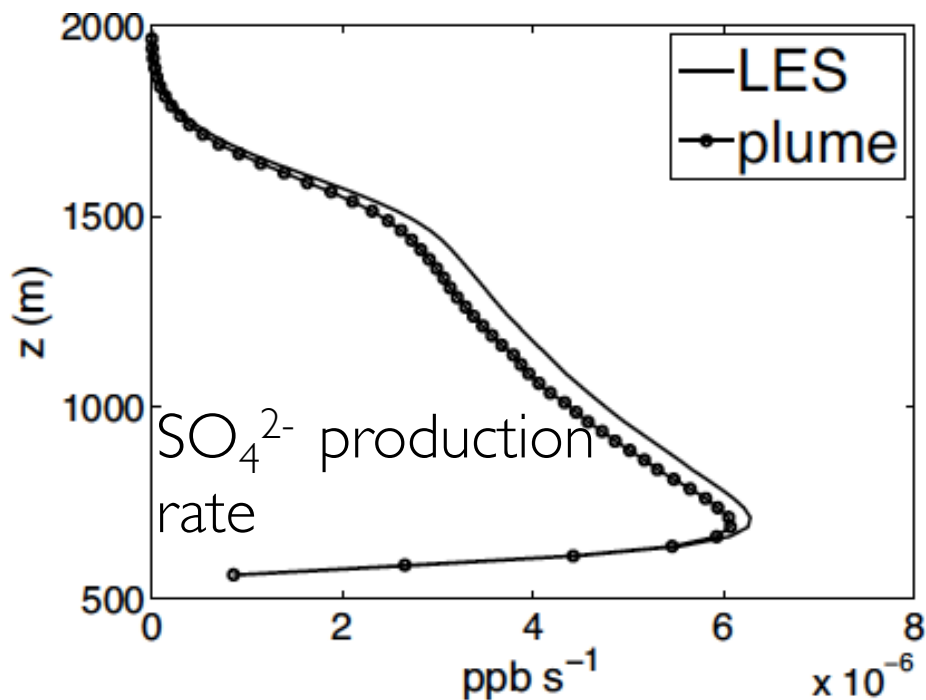
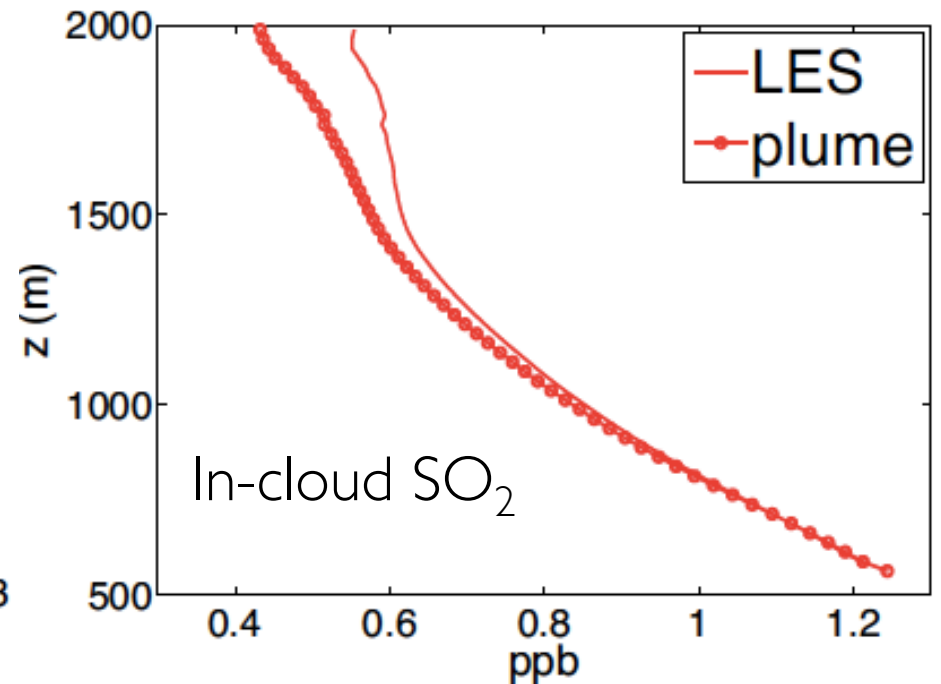
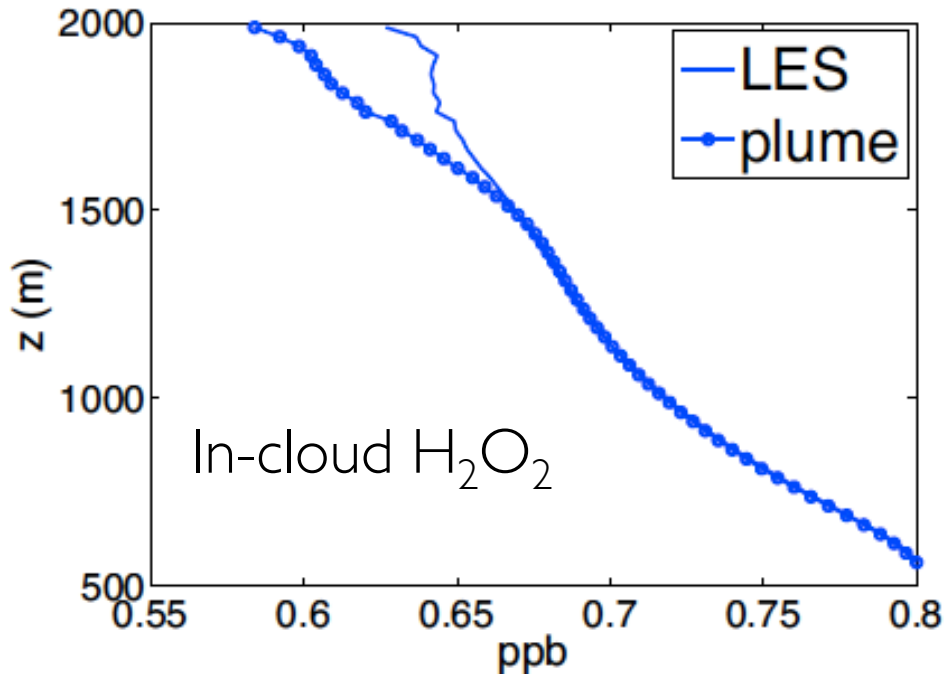
Single step errors

Reference case:

SO₂ flux: 0.024 ppb kg m⁻² s⁻¹

H₂O₂ value: 0.9ppb

$$k = 1 \times 10^{-3} s^{-1} ppb^{-1} (g / kg)^{-1}$$

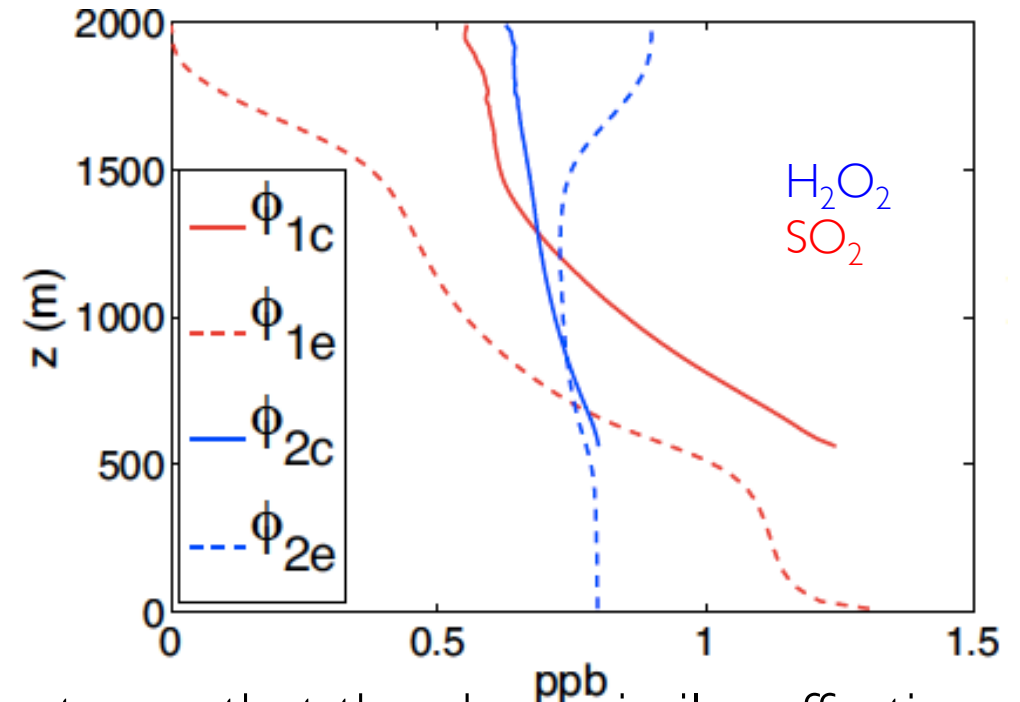


In-cloud tracer values are well represented below the inversion layer

The bulk plume underestimates the reaction rates by $\sim 6\%$

Why are in-cloud tracer values well-modeled in the bulk plume?

While this reaction is considered fast from a global model perspective, it's relatively slow (tens of minutes) compared to the timescale of eddy mixing in cumulus clouds so that:



SO_2 is well correlated with total water so that they have similar effective entrainment/detrainment rates.

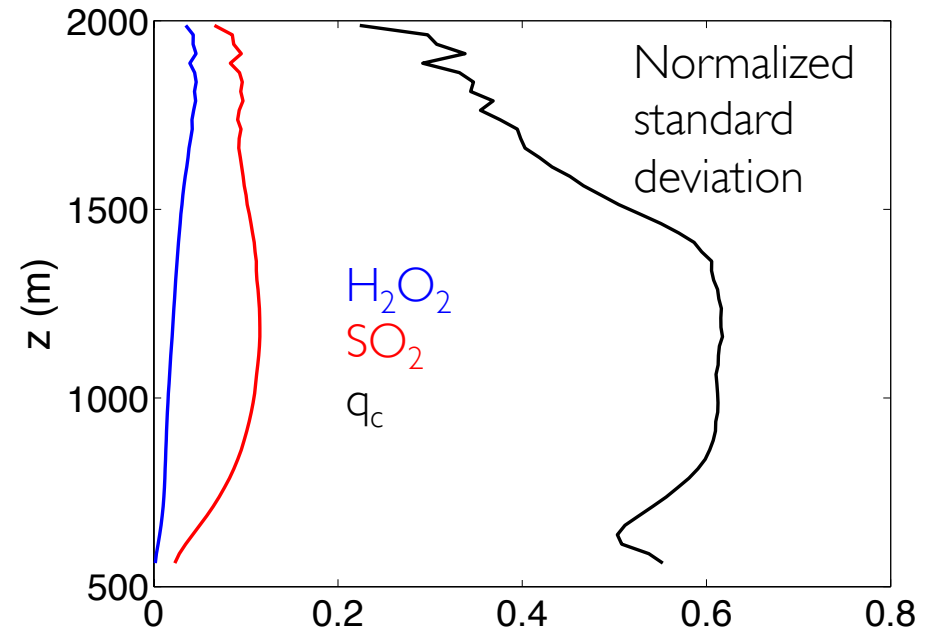
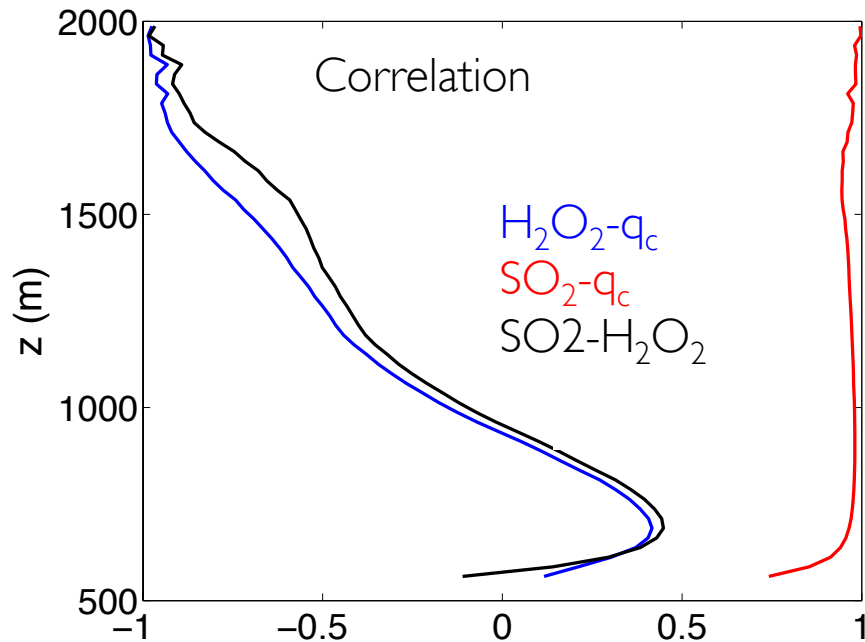
Because of the slow reaction and the small vertical gradient, H_2O_2 values in and out of clouds are similar so entrainment has less of an effect.

Why plume model underestimates the reaction rates

$$1 - \frac{R_{LES}}{R_{plume}} \approx -C_{SO_2, q_c} \sigma_{SO_2} \sigma_{q_c} - C_{H_2O_2, q_c} \sigma_{H_2O_2} \sigma_{q_c} - C_{H_2O_2, SO_2} \sigma_{H_2O_2} \sigma_{SO_2}$$

$\approx -\sigma_{SO_2} \sigma_{q_c}$ Where σ is the standard deviation divided by the mean

- The reaction is “slow” so that SO_2 is well correlated with cloud liquid water
- Fractional variance of H_2O_2 is small compared to that of SO_2



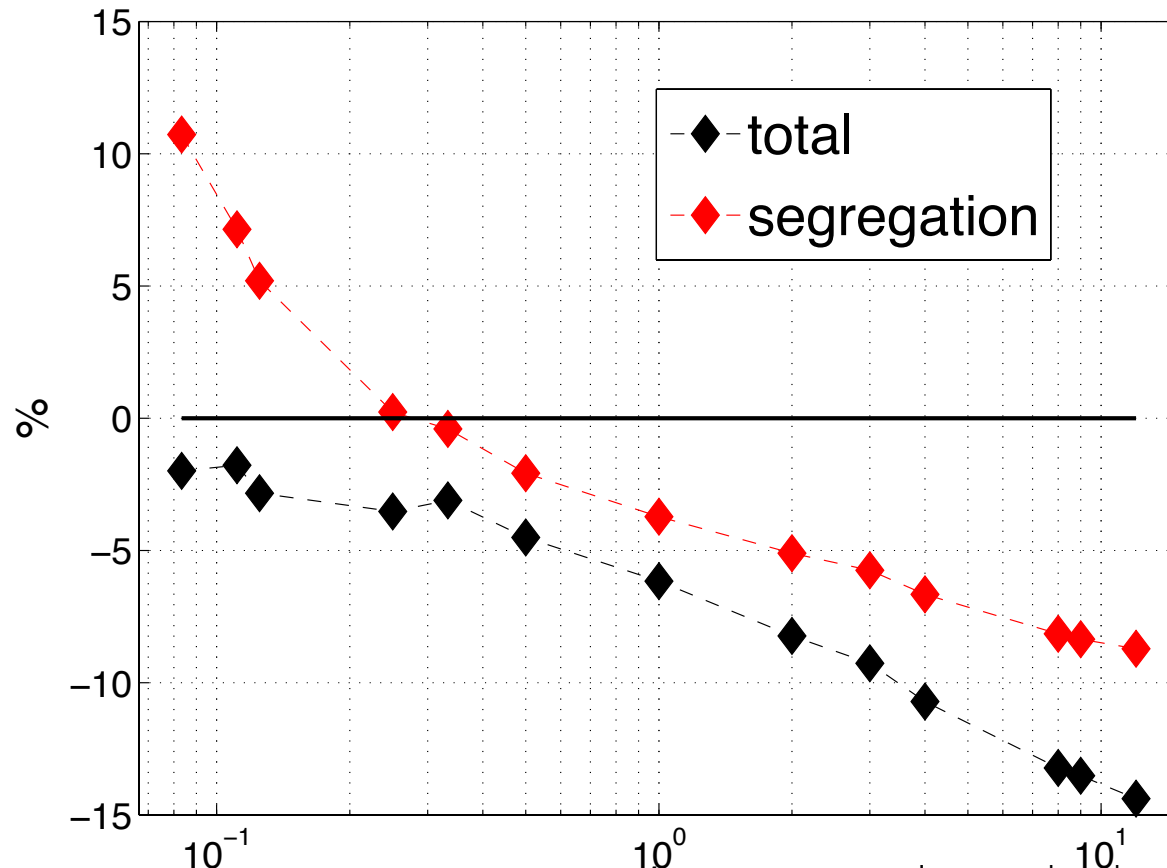
Varying two control parameters

Reaction rate constant: k

Relative magnitudes of sources of SO_2 and H_2O_2

Varying relative magnitudes of sources of SO_2 and H_2O_2

R_{err}



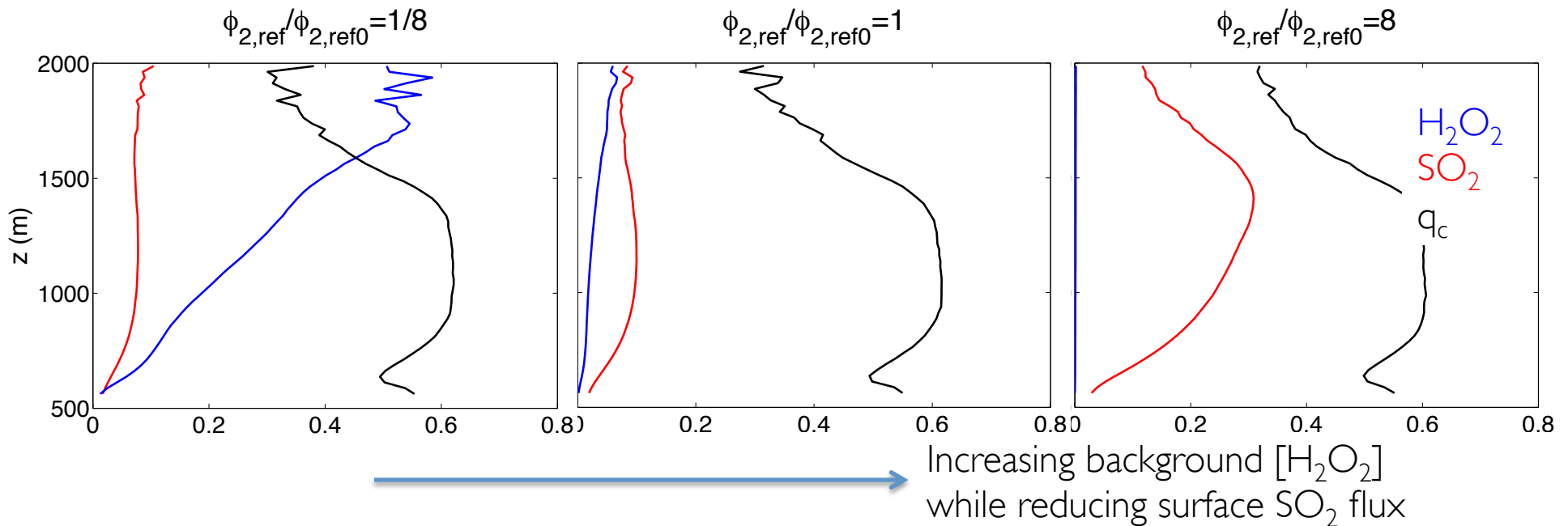
$\phi_{2,\text{ref}}/\phi_{2,\text{ref}0}$

Increasing background $[\text{H}_2\text{O}_2]$
while reducing surface SO_2 flux by
the same factor

Varying relative magnitudes of sources of SO₂ and H₂O₂

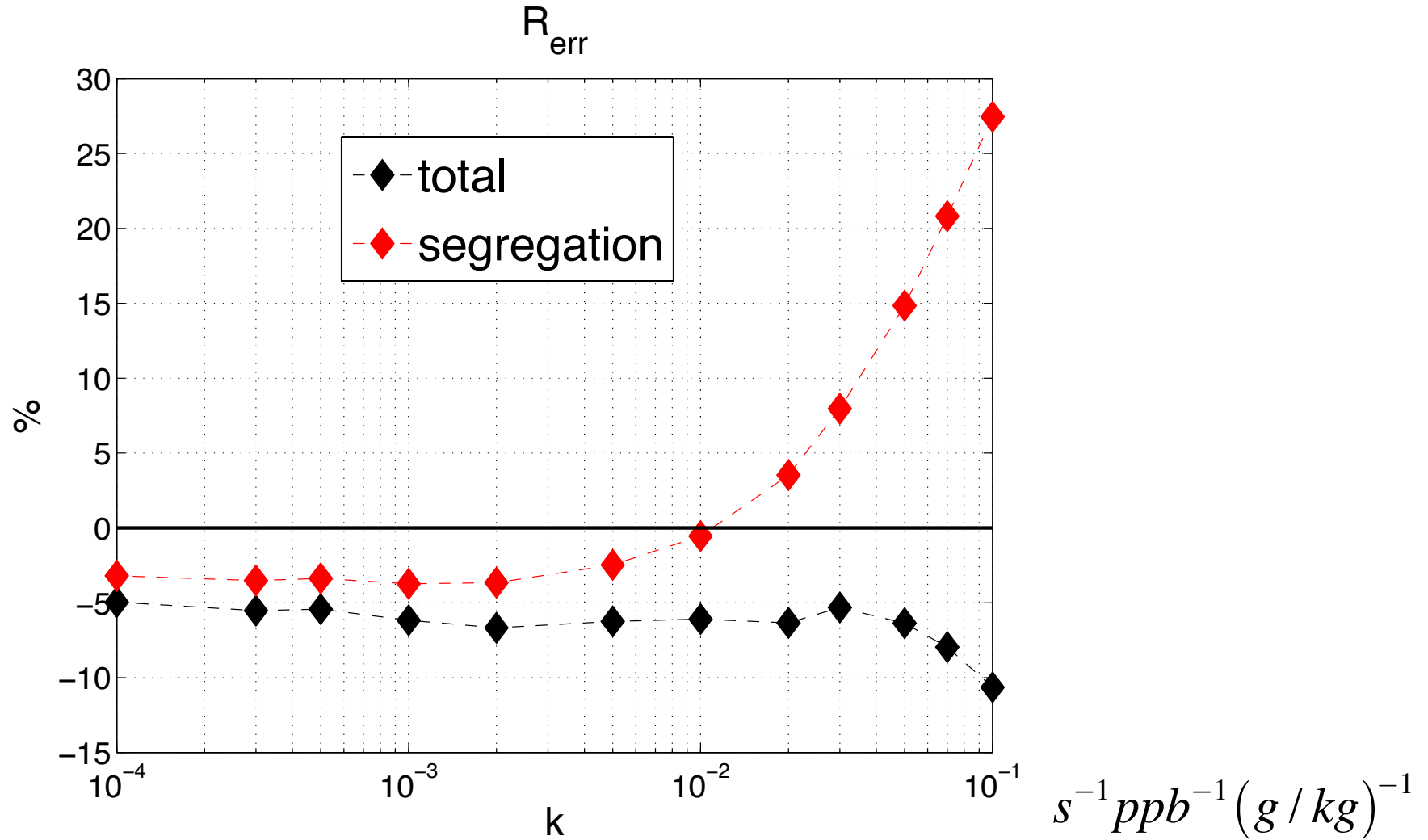
$$1 - \frac{R_{LES}}{R_{plume}} \approx \underbrace{-C_{SO_2, q_c} \sigma_{SO_2} \sigma_{q_c}}_{\text{Negative}} \underbrace{-C_{H_2O_2, q_c} \sigma_{H_2O_2} \sigma_{q_c} - C_{H_2O_2, SO_2} \sigma_{H_2O_2} \sigma_{SO_2}}_{\text{Positive}}$$

Normalized standard deviation

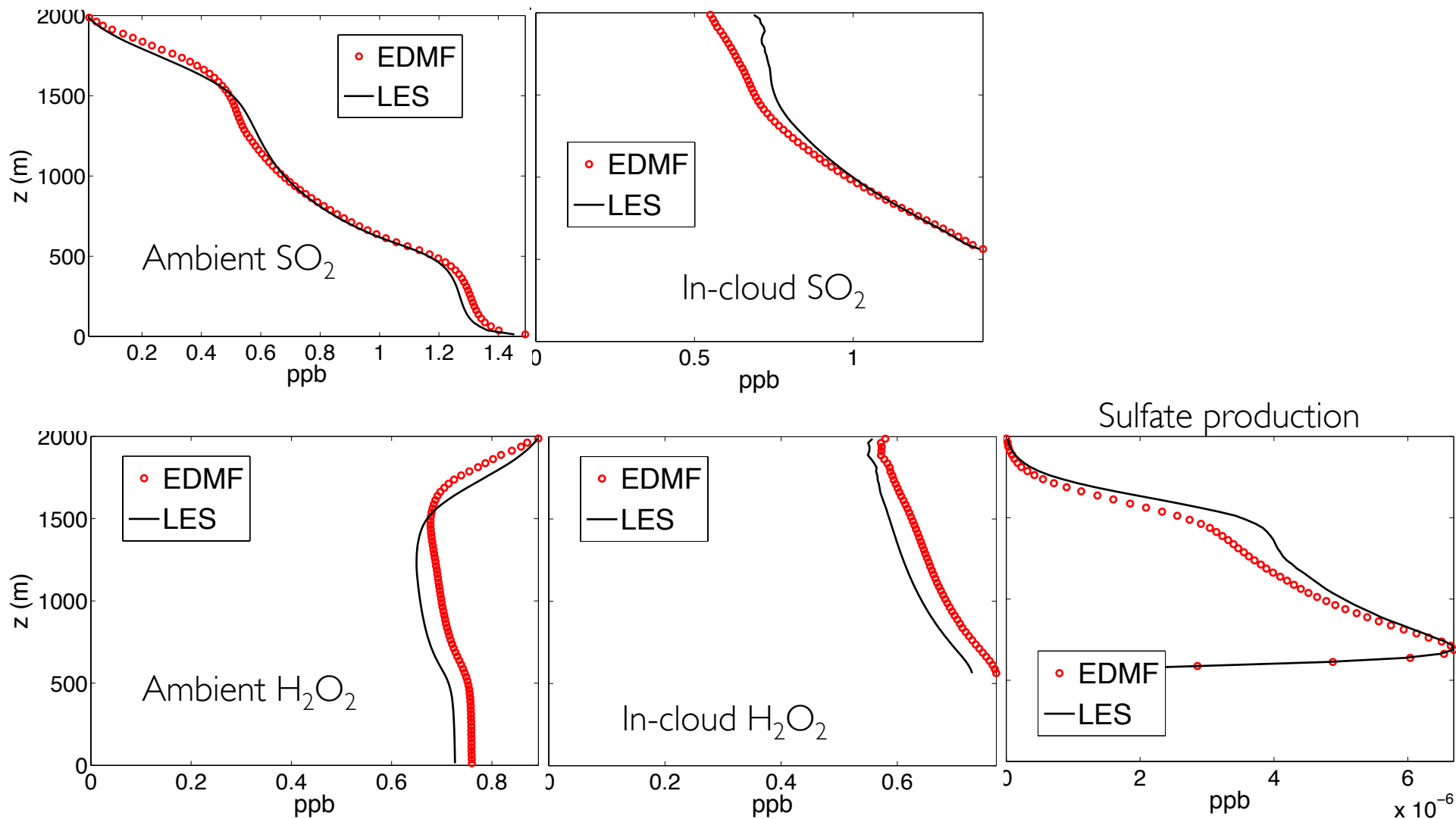


As H₂O₂ increases relative to SO₂, there is a stronger vertical gradient in SO₂ and stronger fractional in-cloud variance.

Varying the reaction rate constant



Equilibrium state using the EDMF



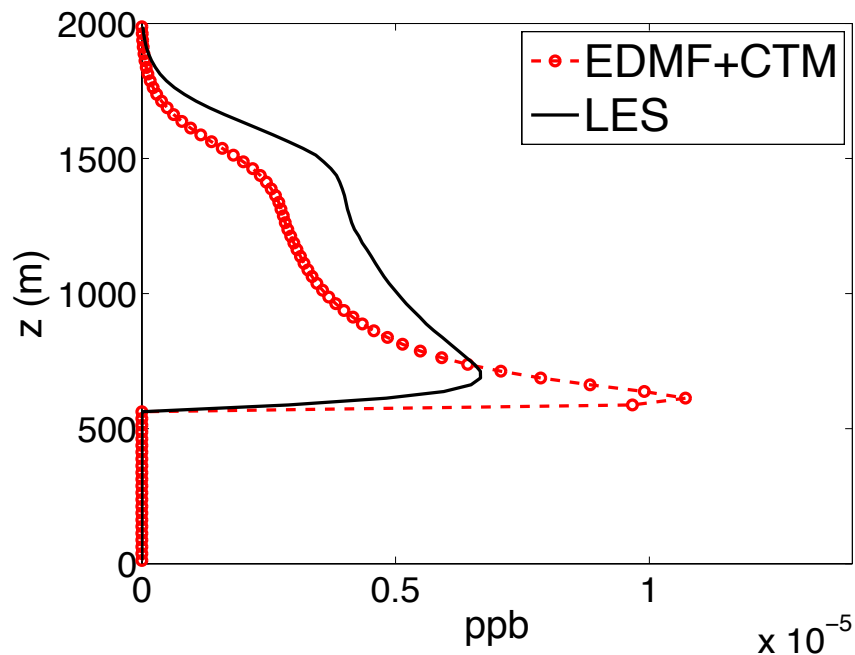
Conclusions

- The EDMF approach with a bulk plume can represent aqueous reaction in shallow cumuli quite well when entrainment/detrainment rates and eddy diffusivity are diagnosed using conservative tracers like total water.
- This is because the aqueous reactions are slow compared to eddy mixing timescale in shallow cumuli.
- The bulk plume underestimates the reaction rate by 5-10% and errors are larger with faster reactions and in H_2O_2 dominated cases (for understood reasons)
- This approach can be used with super-parameterization.

Performance of current chemical transport model treatment

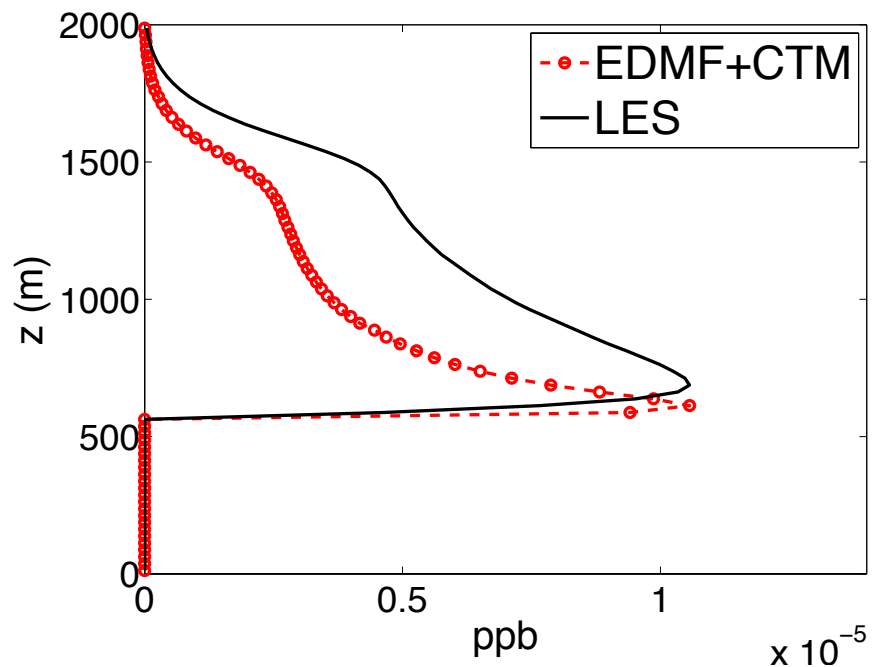
$$k = 1 \times 10^{-3} s^{-1} ppb^{-1} (g/kg)^{-1}$$

R



$$k = 2 \times 10^{-3} s^{-1} ppb^{-1} (g/kg)^{-1}$$

R

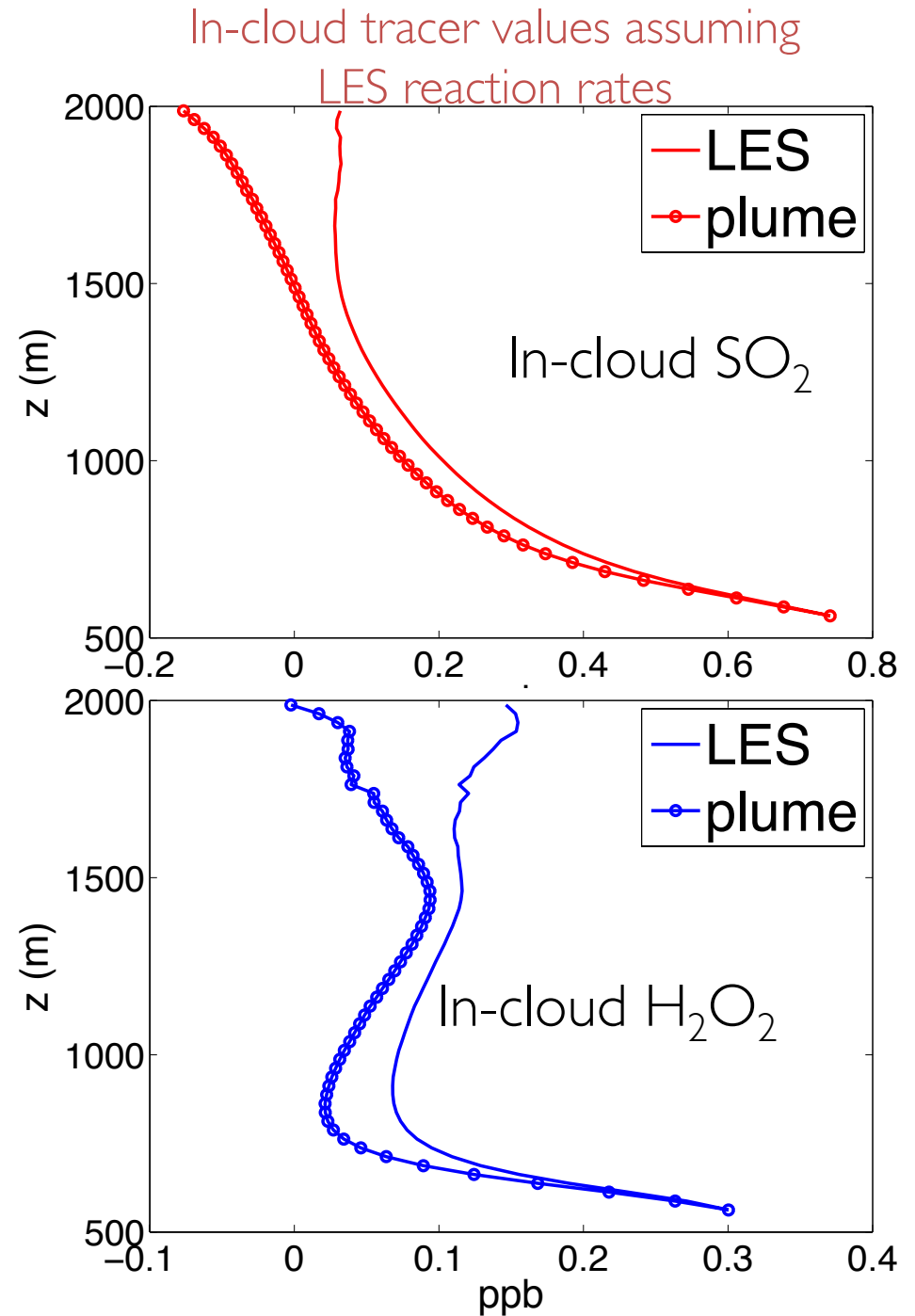


Uses cloud fraction and vertical tracer fluxes diagnosed from the LES with a 1-hr global model time step.

Errors in entrainment/ detrainment rates

$$k = 0.1 s^{-1} ppb^{-1} (g/kg)^{-1}$$

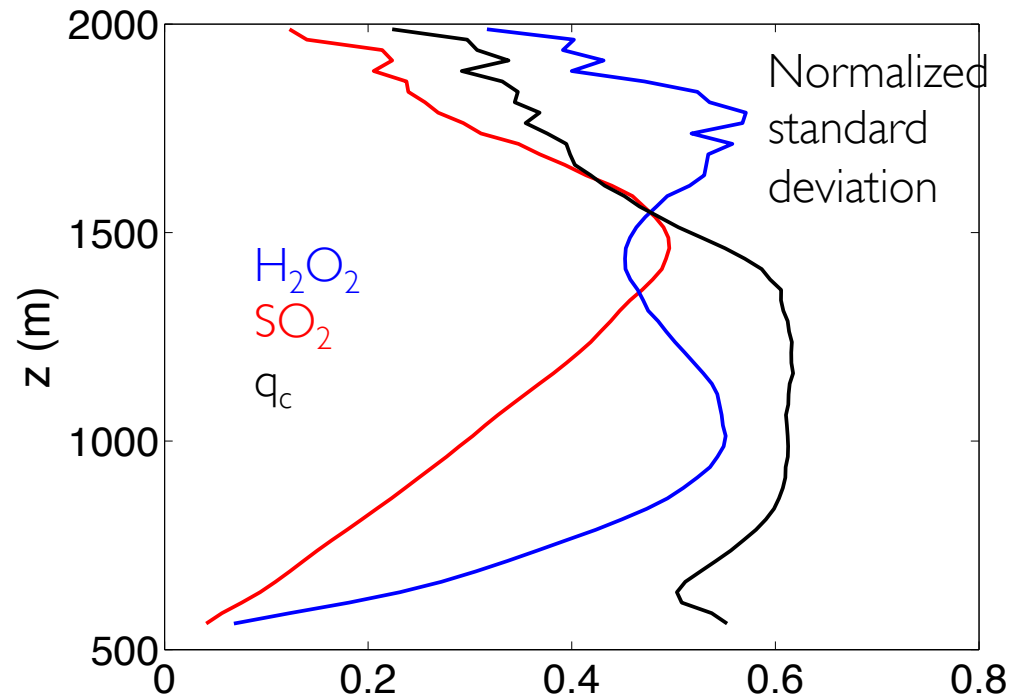
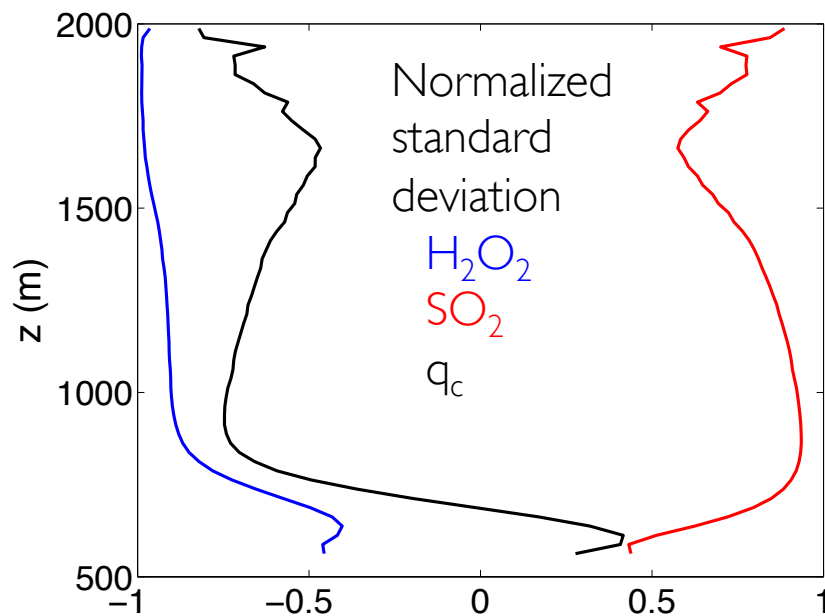
For large k , effective entrainment and detrainment rates for the tracers become more different from those for conserved variables.



Segregation error $k = 0.1 s^{-1} ppb^{-1} (g/kg)^{-1}$

$$\frac{R_{LES}}{R_{plume}} - 1 \approx \sigma_{SO_2} \sigma_{q_c} - \sigma_{H_2O_2} \sigma_{q_c} - \sigma_{H_2O_2} \sigma_{SO_2}$$

As k increases, H_2O_2 becomes anti-correlated with SO_2 and cloud water, and fraction variance of H_2O_2 becomes large and the second and third terms become more dominant

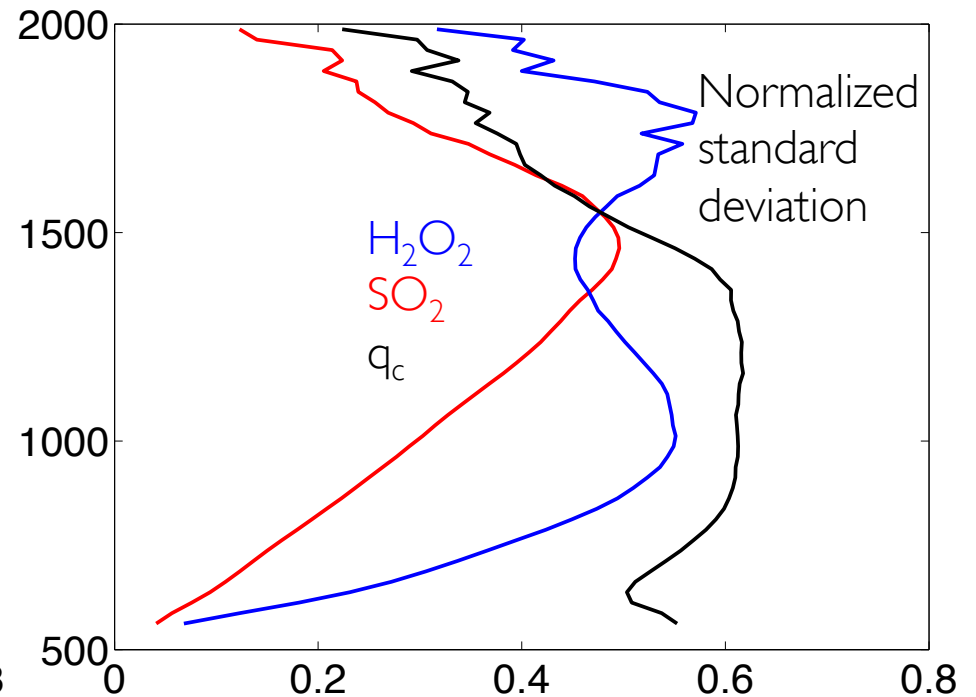
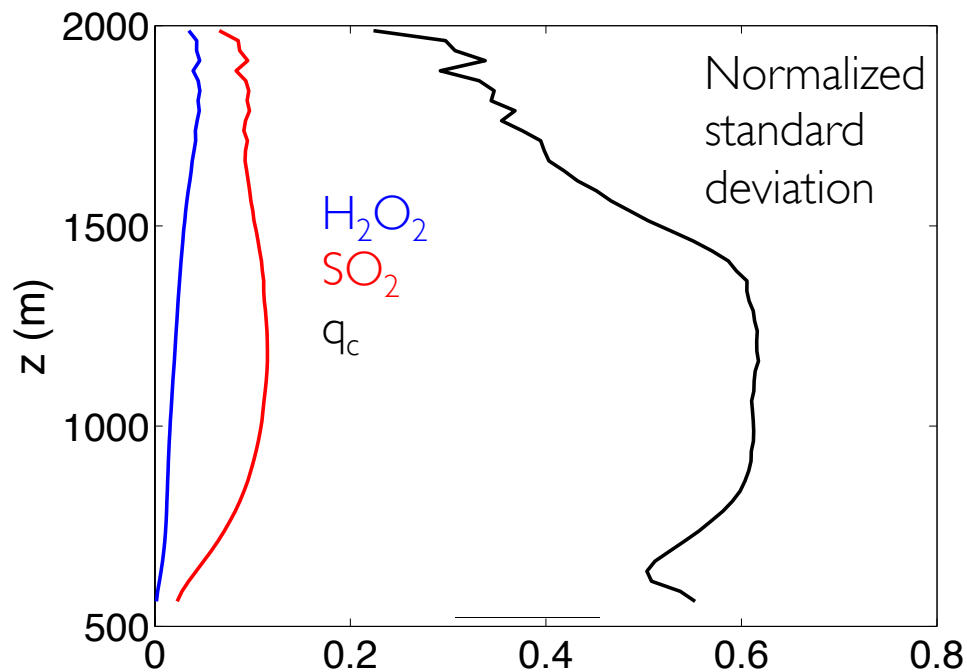


Segregation error

$$\frac{R_{LES}}{R_{plume}} - 1 \approx \sigma_{SO_2} \sigma_{q_c} - \sigma_{H_2O_2} \sigma_{q_c} - \sigma_{H_2O_2} \sigma_{SO_2}$$

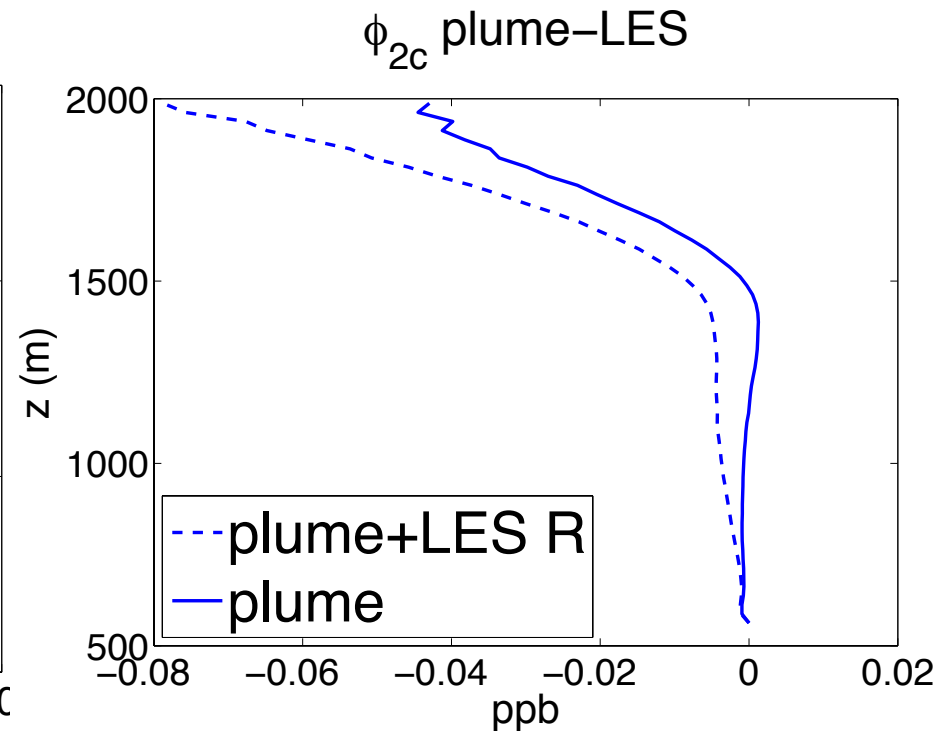
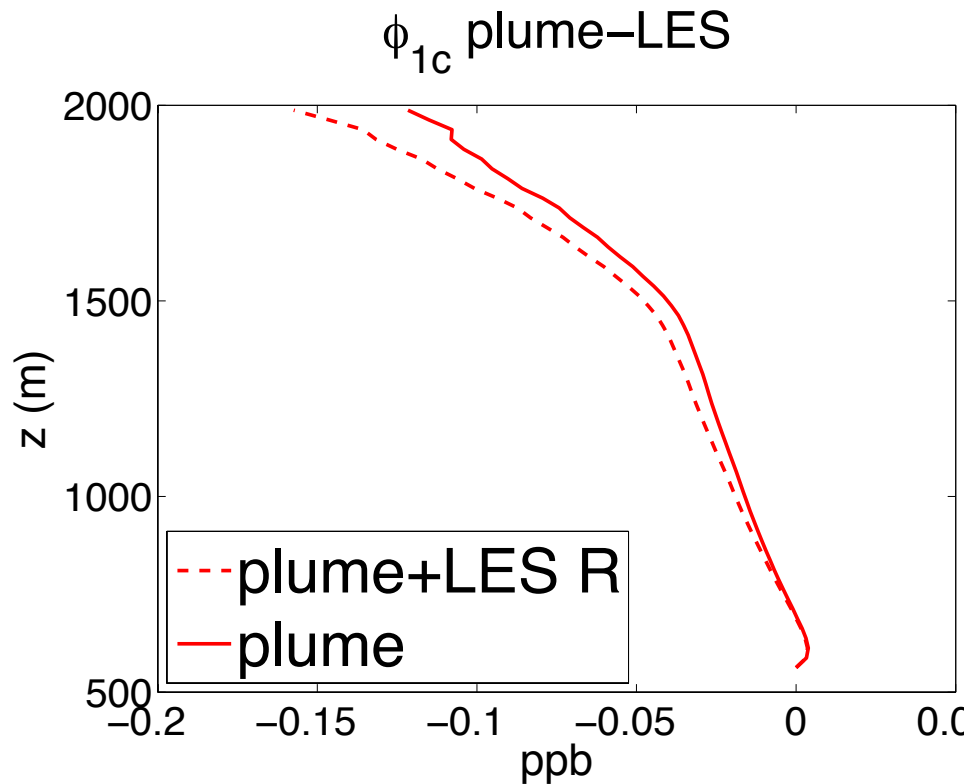
$$k = 0.001 s^{-1} ppb^{-1} (g / kg)^{-1}$$

$$k = 0.1 s^{-1} ppb^{-1} (g / kg)^{-1}$$



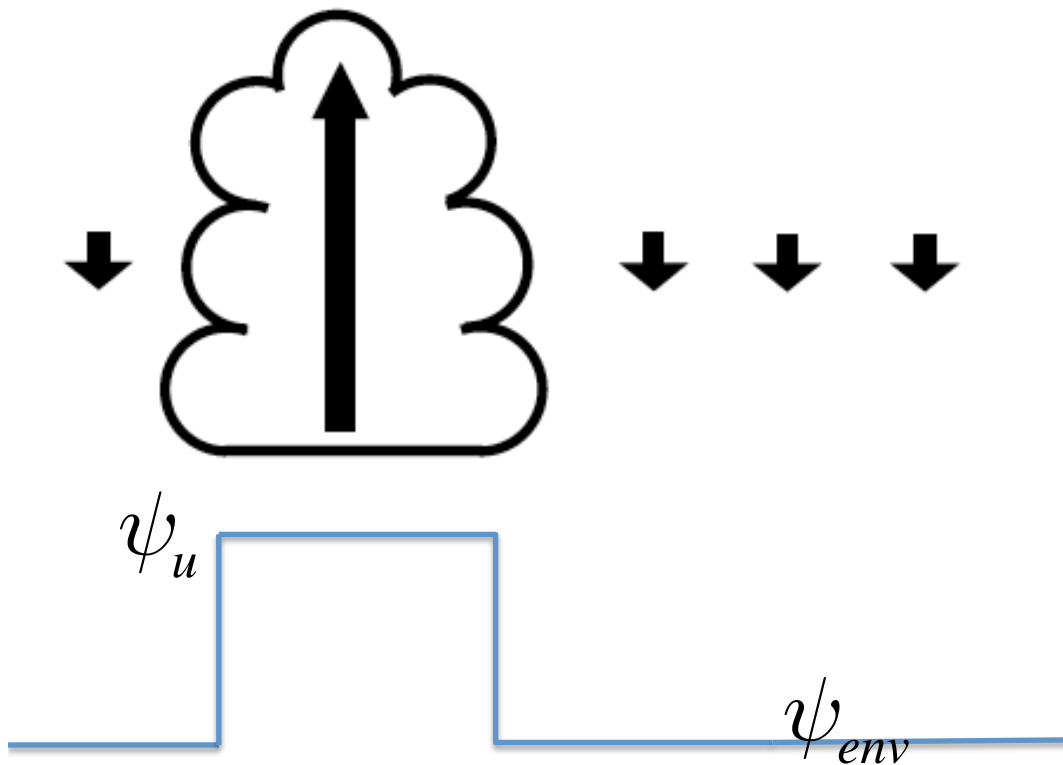
Why are in-cloud tracer values well-modeled with the bulk plume?

3. Error in reaction rates partly cancels error in entrainment



Bulk plume representation for mass flux

Assume clouds and environment at a given height have uniform properties within each category



$$\frac{\partial M_u}{\partial z} = M_u(\varepsilon - \delta),$$

$$\frac{\partial \psi_u}{\partial z} = \varepsilon(\bar{\psi} - \psi_u) + S_\psi.$$