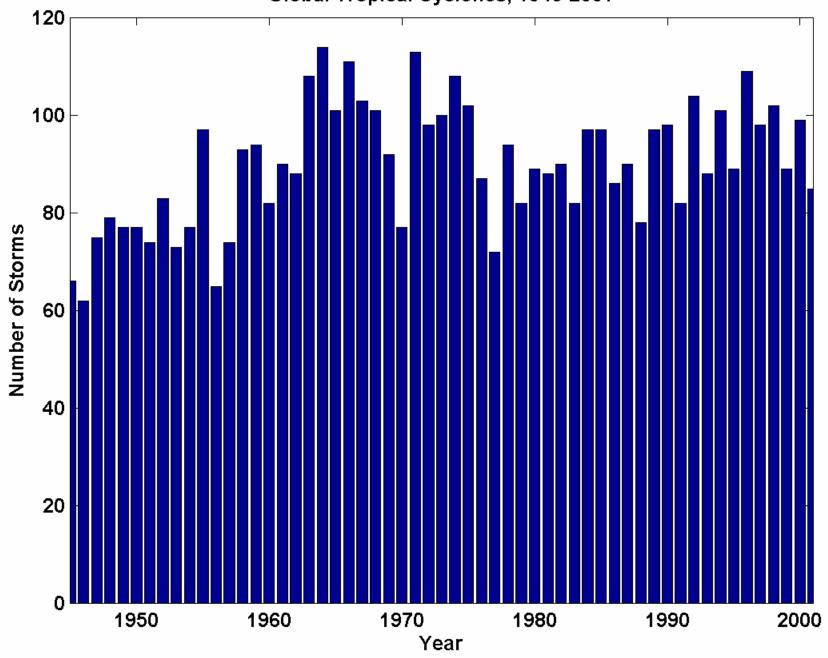
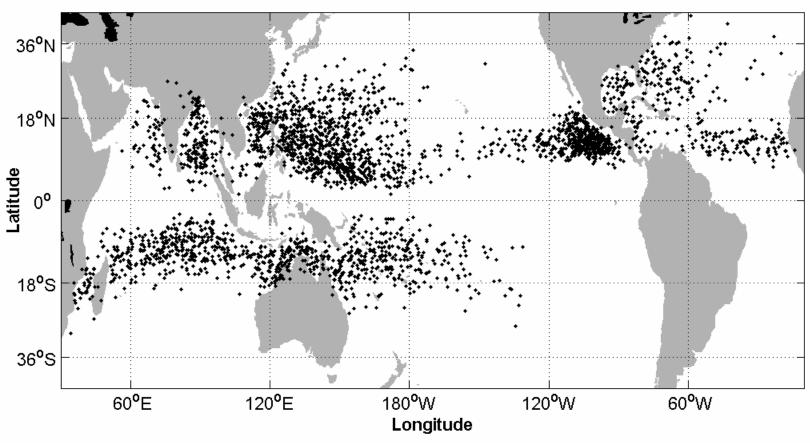
Genesis

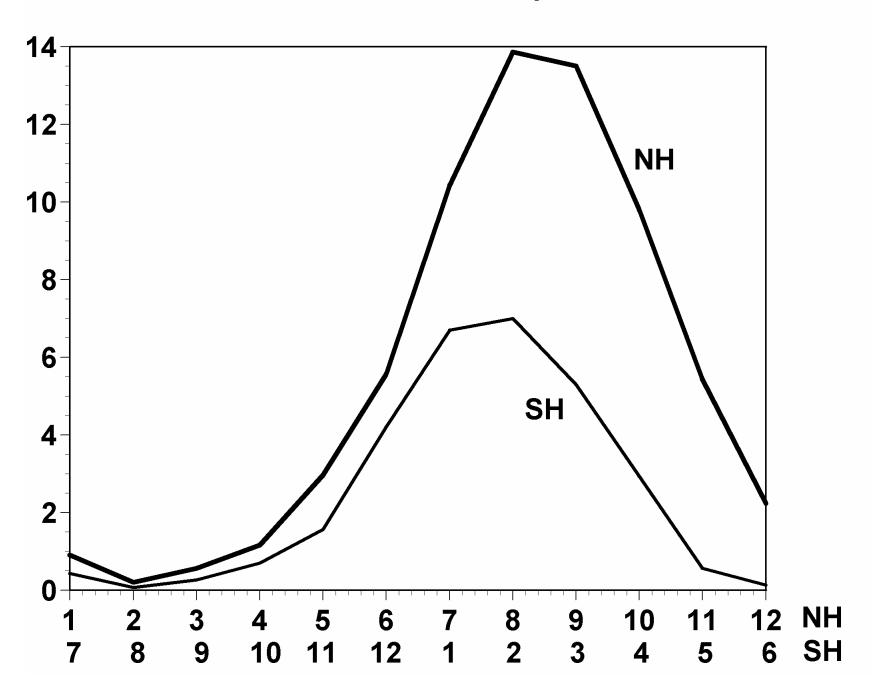
Global Tropical Cyclones, 1945-2001



Global Genesis Events 1971-2001



Number of Genesis Events per Month



Key Aspects of Genesis:

- Tropical cyclones results only from finiteamplitude perturbations of the normal state of the tropical atmosphere
- Empirically favors regions of large potential intensity, large low-level vorticity, and humid middle tropopshere with little vertical wind shear
- Evidence from field experiments suggest that TCs often originate in cold-core mesosystems

Empirical Approach to Environmental Control of TC Frequency:

Develop an empirical index based on monthly re-analysis data

Test index against geographic, seasonal and interannual variability

Empirical Index:

$$I = \left| 10^{5} \eta \right|^{3/2} \left(\frac{\mathcal{H}}{50} \right)^{3} \left(\frac{V_{pot}}{70} \right)^{3} \left(1 + 0.1 V_{shear} \right)^{-2},$$

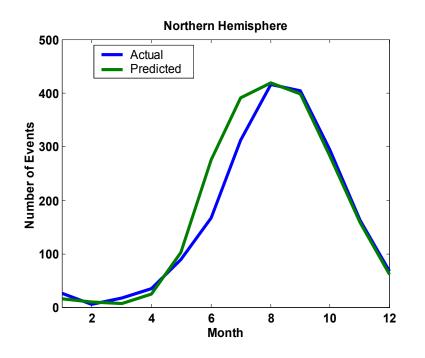
$$\eta \equiv absolute vorticity (s^{-1}),$$

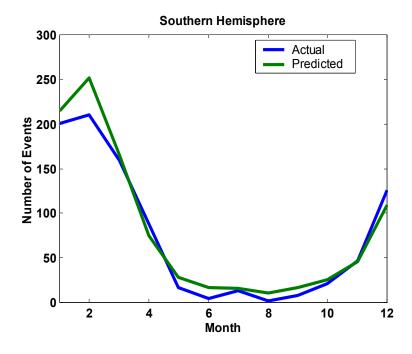
$$V_{pot} = Potential \ wind \ speed \ (ms^{-1}),$$

$$\mathcal{H} \equiv 600 \, mb \, relative \, humidity \, (\%),$$

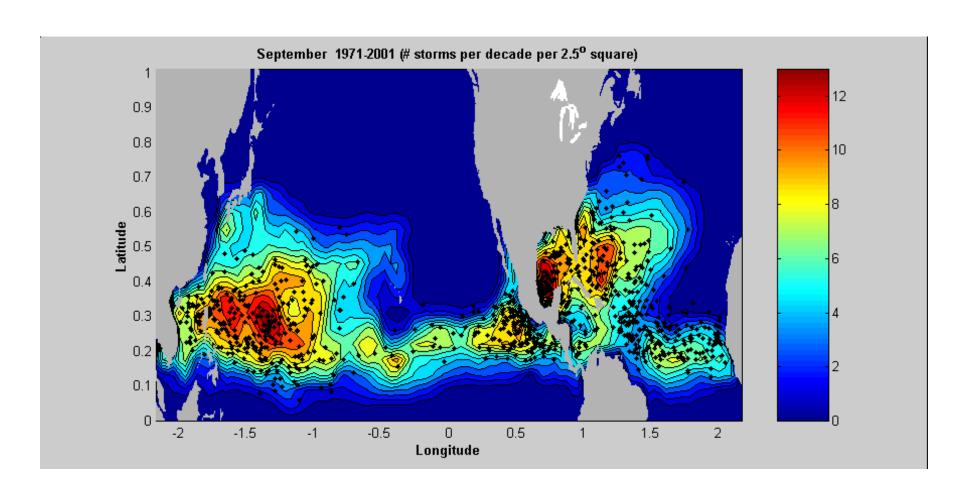
$$V_{shear} \equiv \begin{vmatrix} \mathbf{V}_{850} - \mathbf{V}_{250} \end{vmatrix}$$
 $(ms^{-1}).$

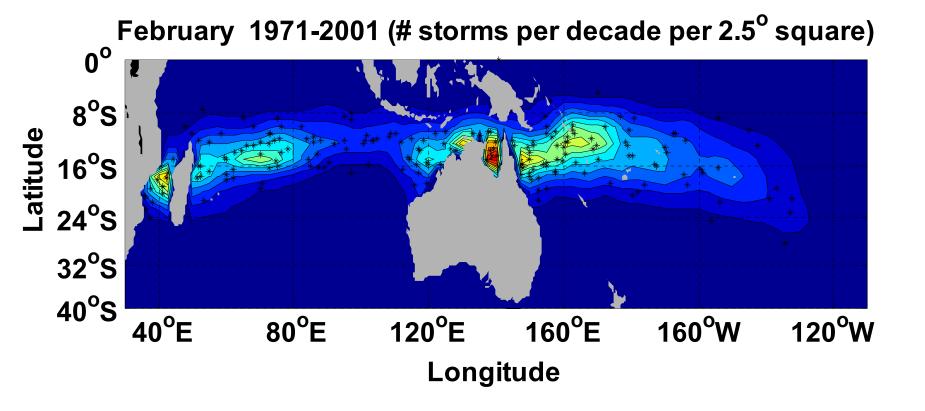
Seasonal Variability:





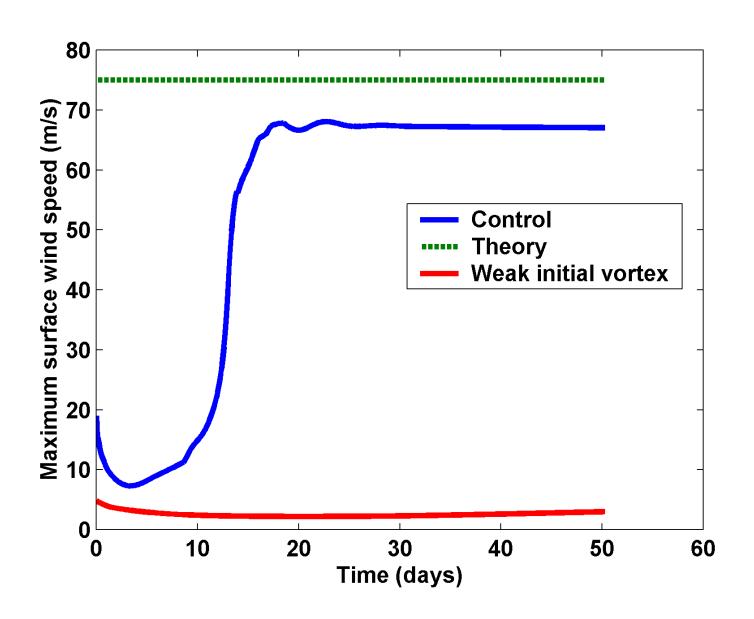
Spatial Variability:

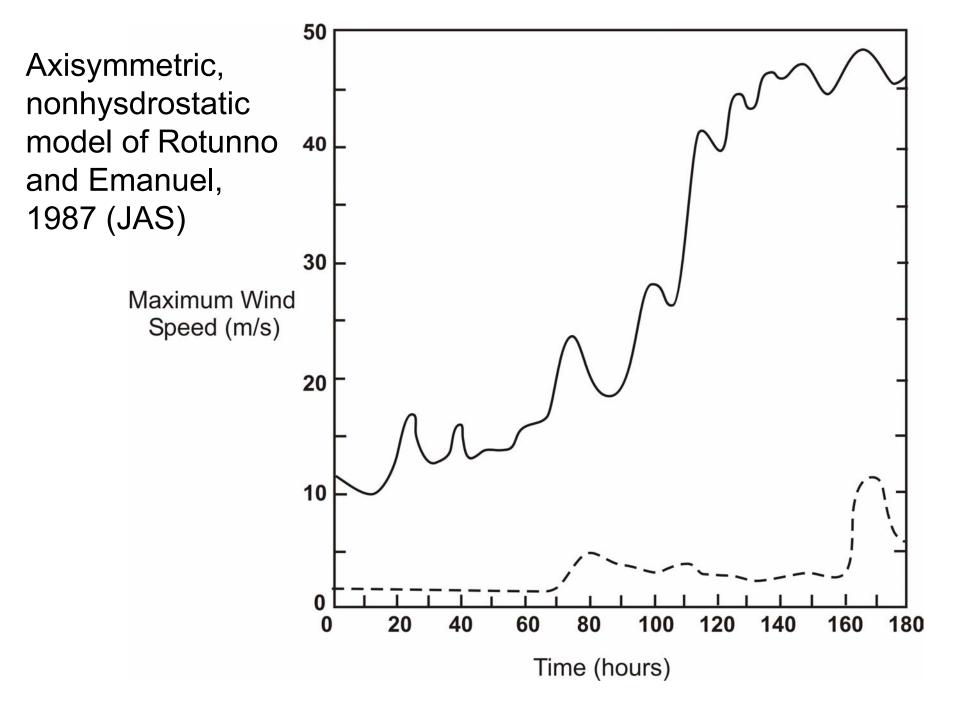




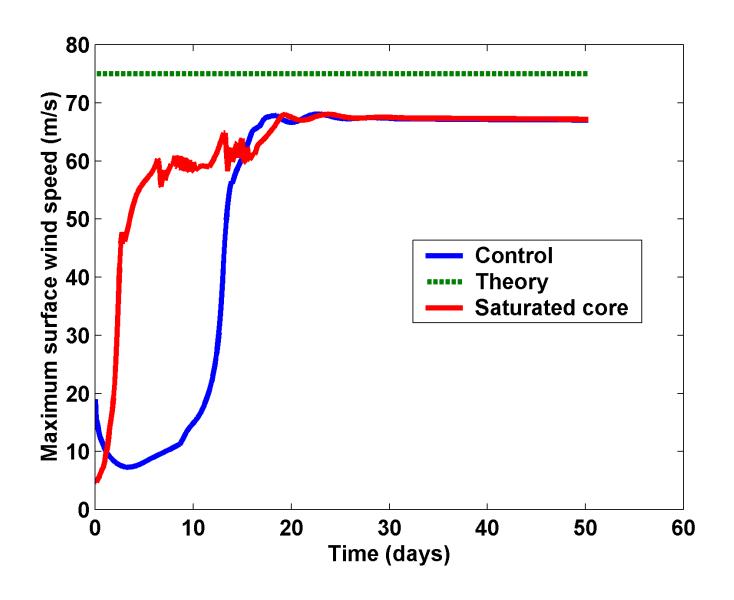
Model Behavior

Model behavior

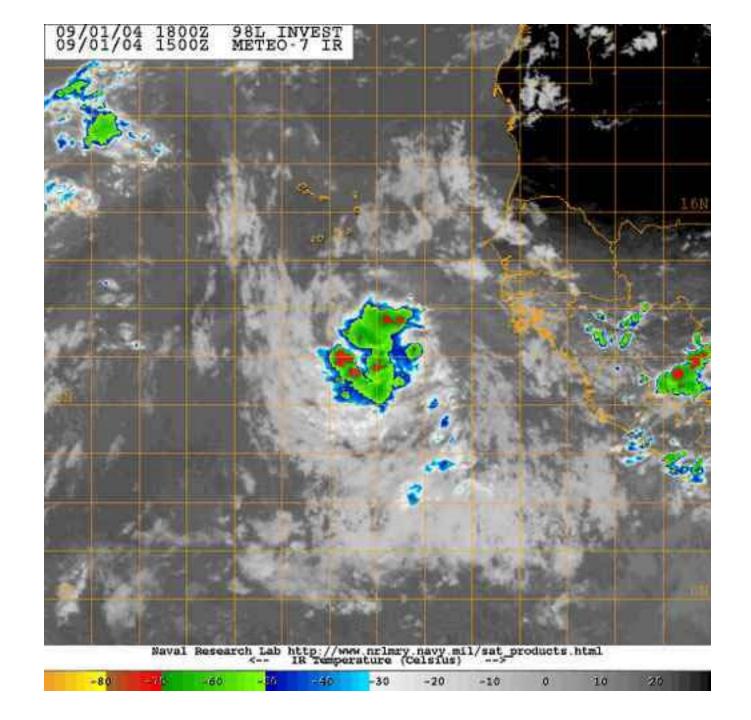


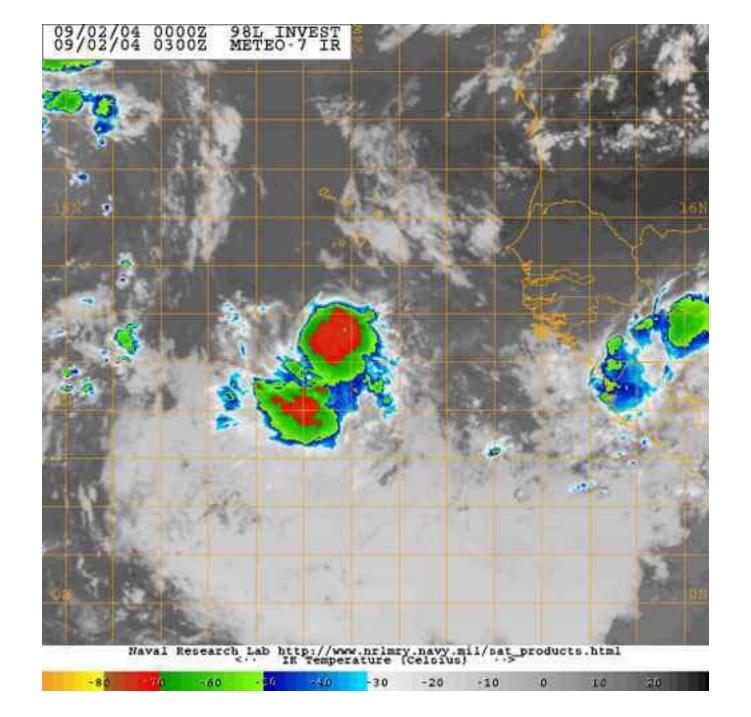


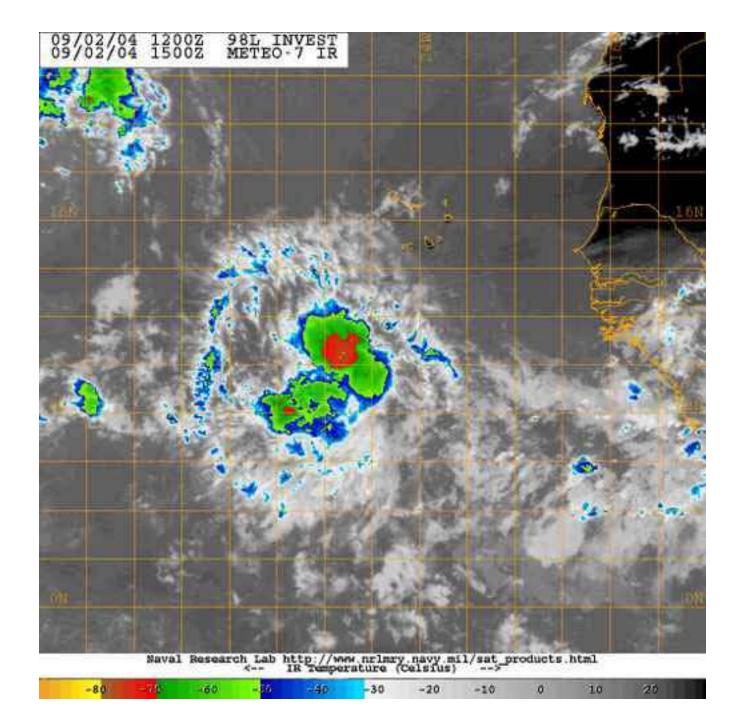
Saturate troposphere inside 100 km in initial state:

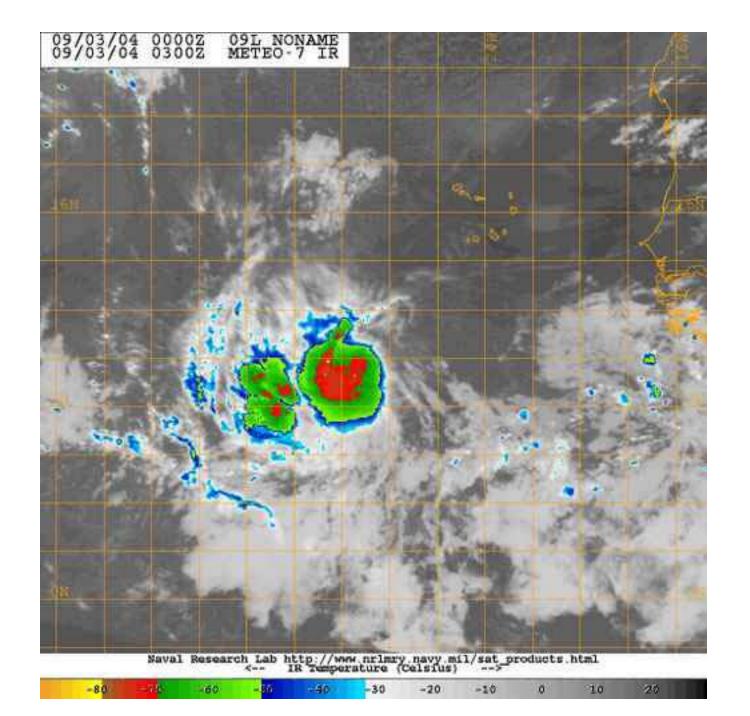


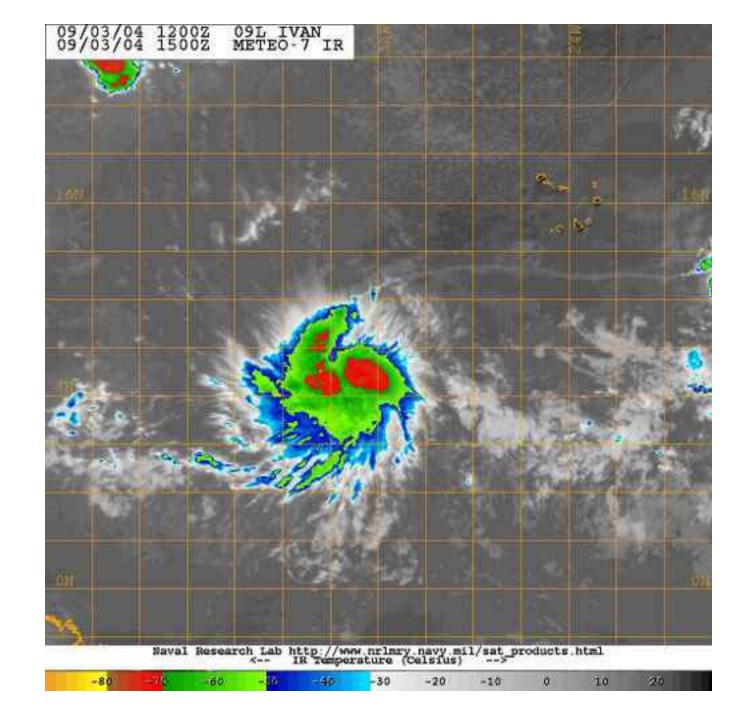
The View from Space: Ivan, 2004

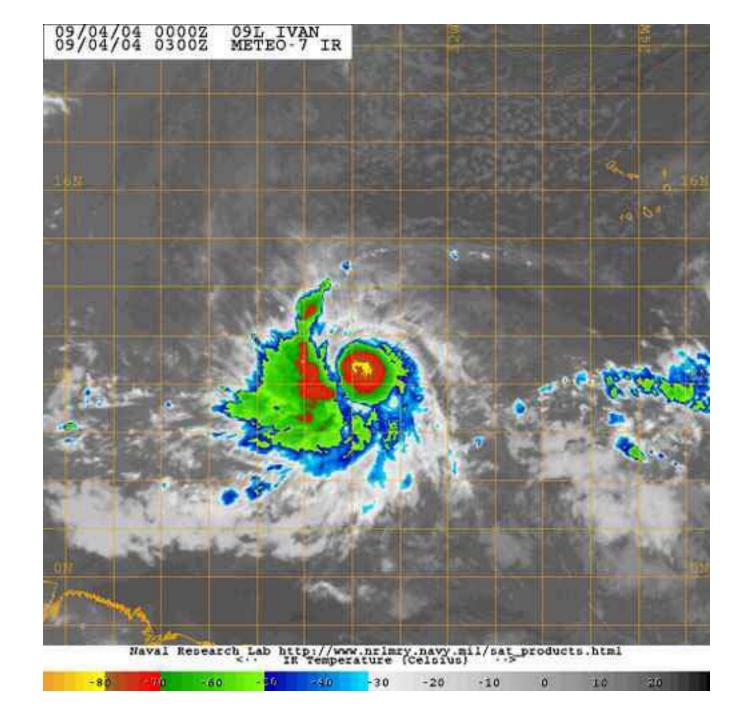


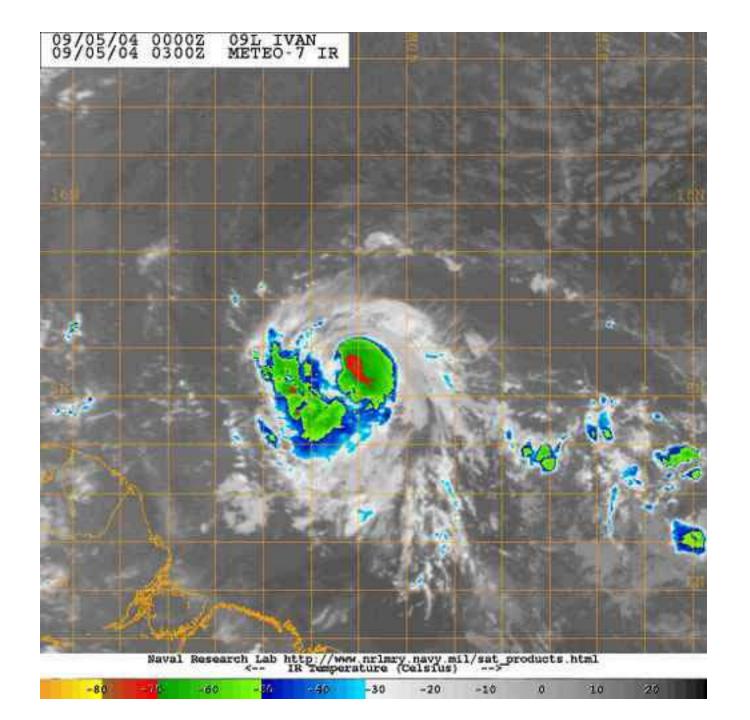


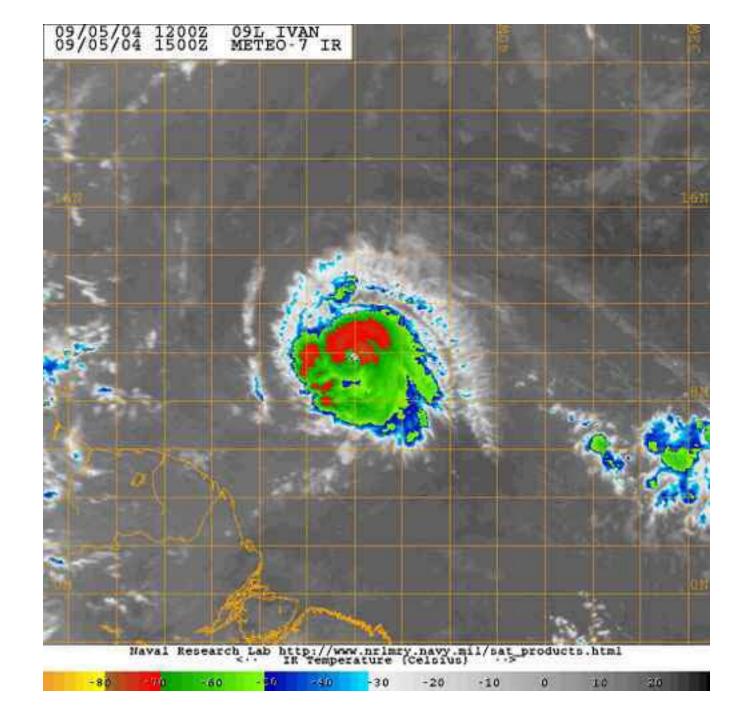


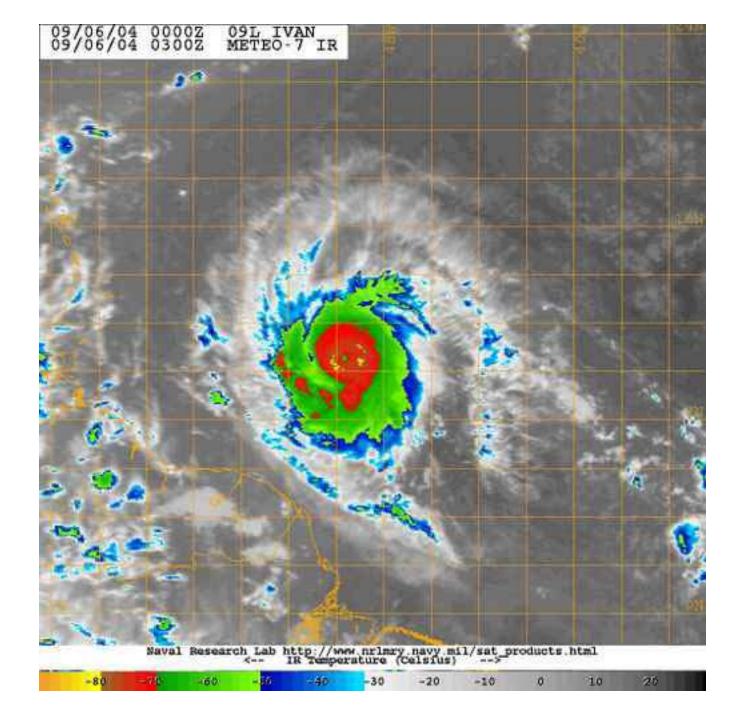


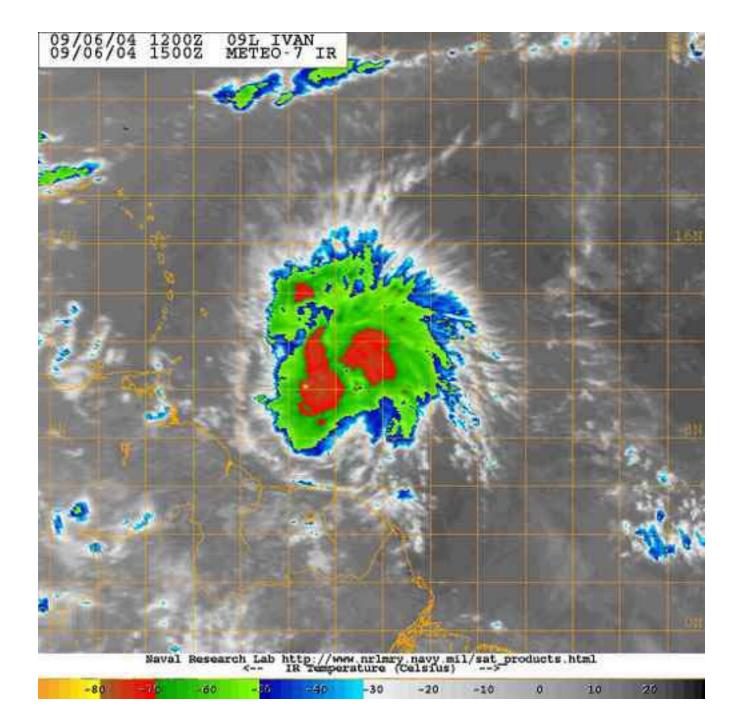


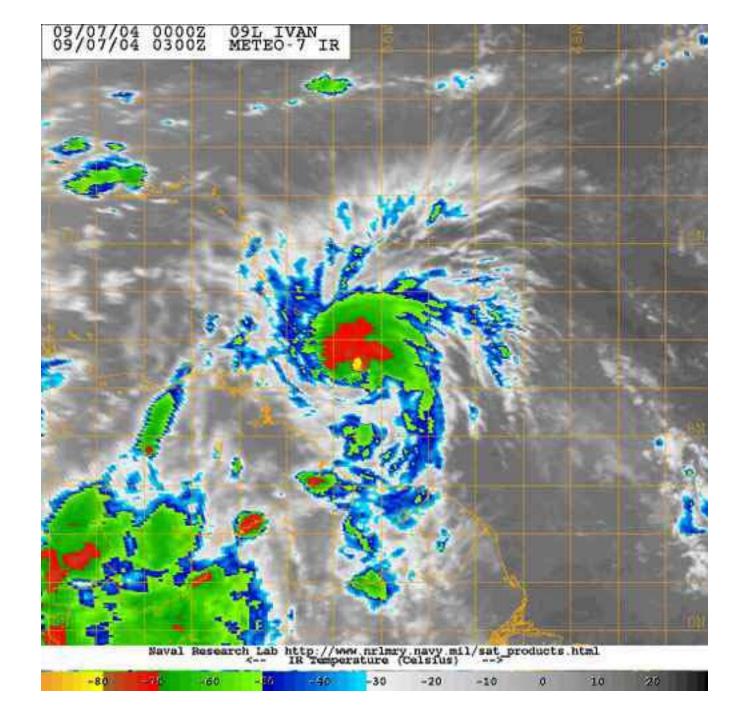


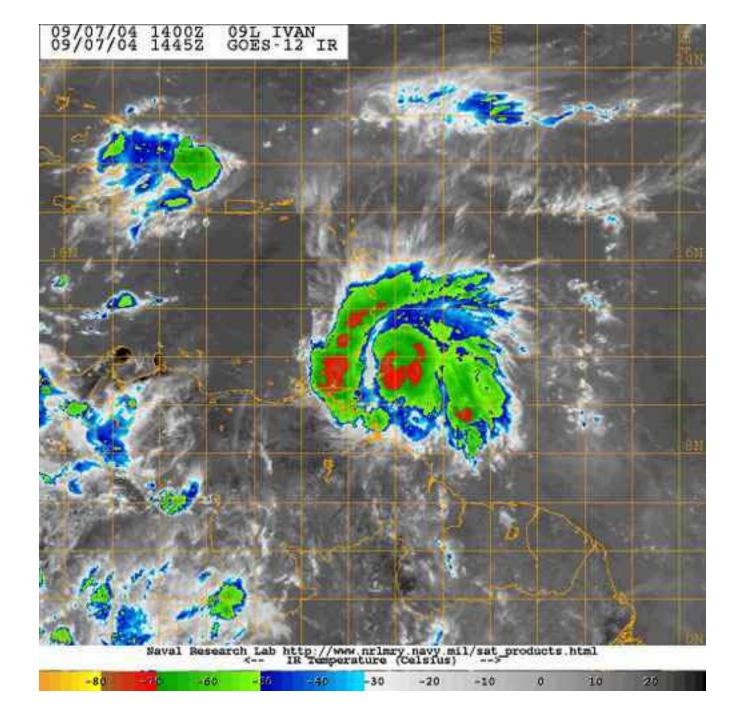


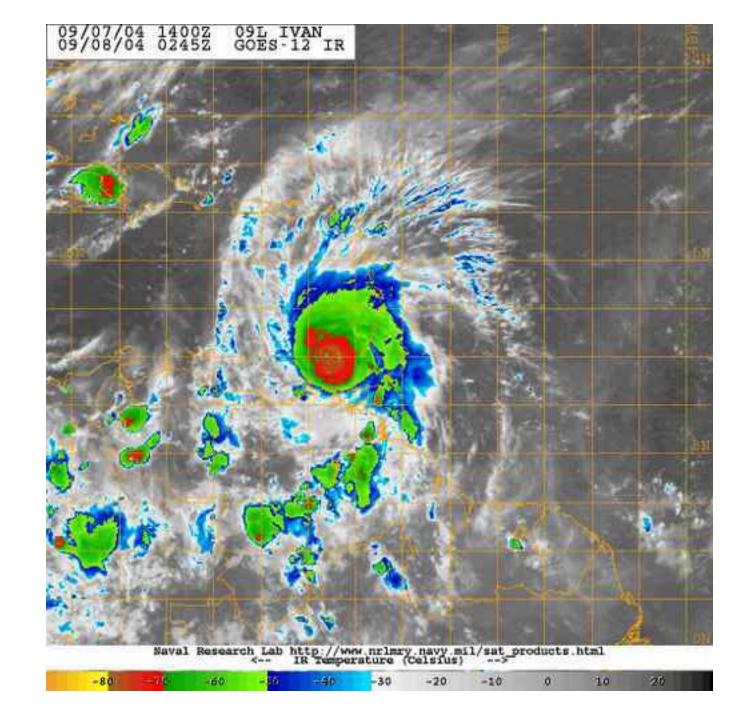


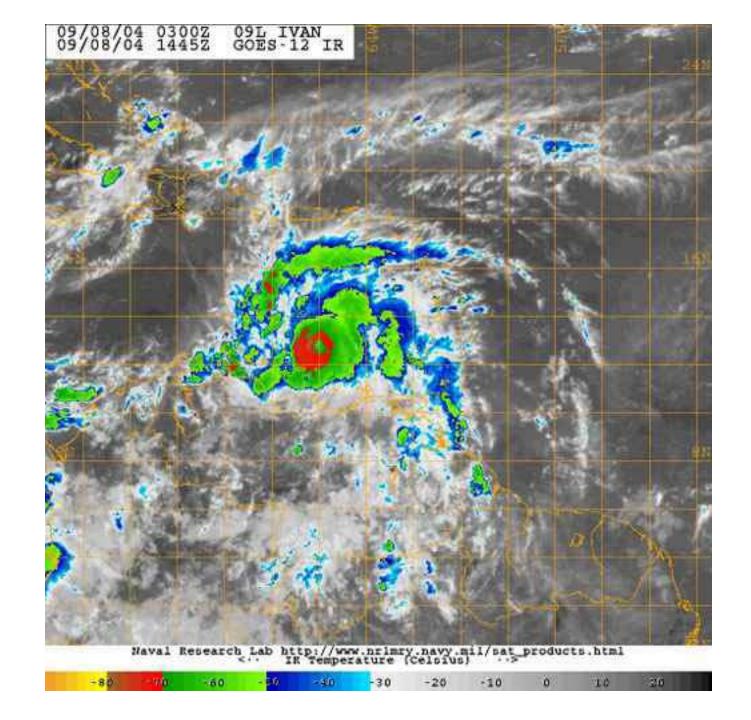


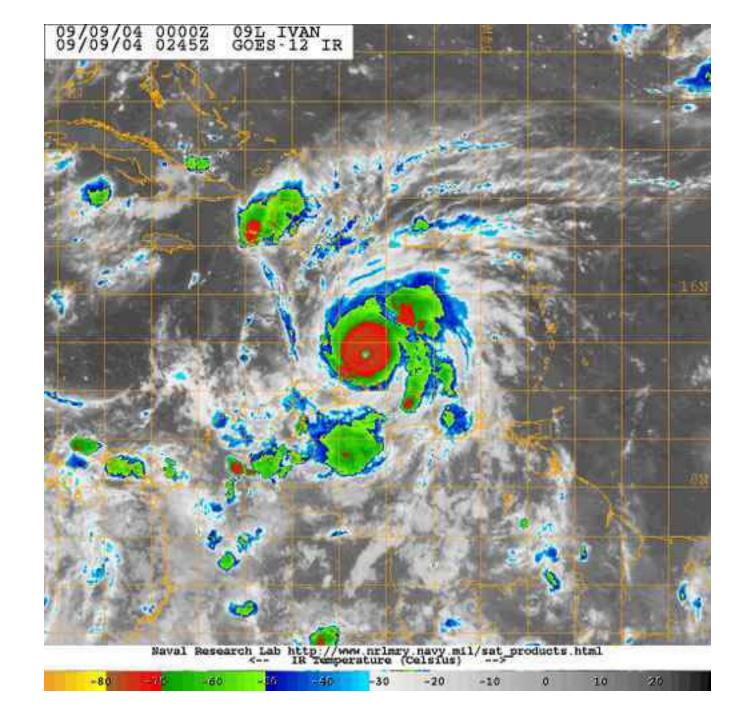


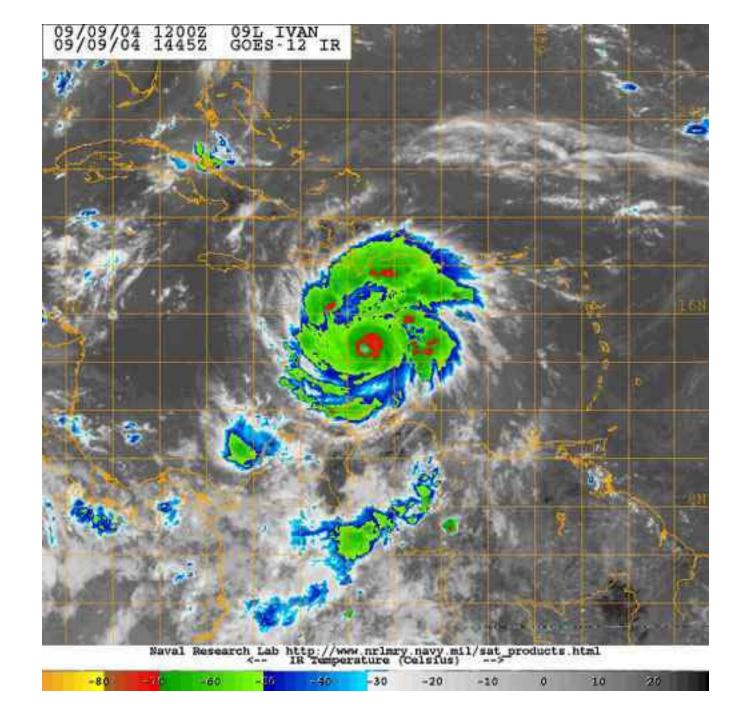


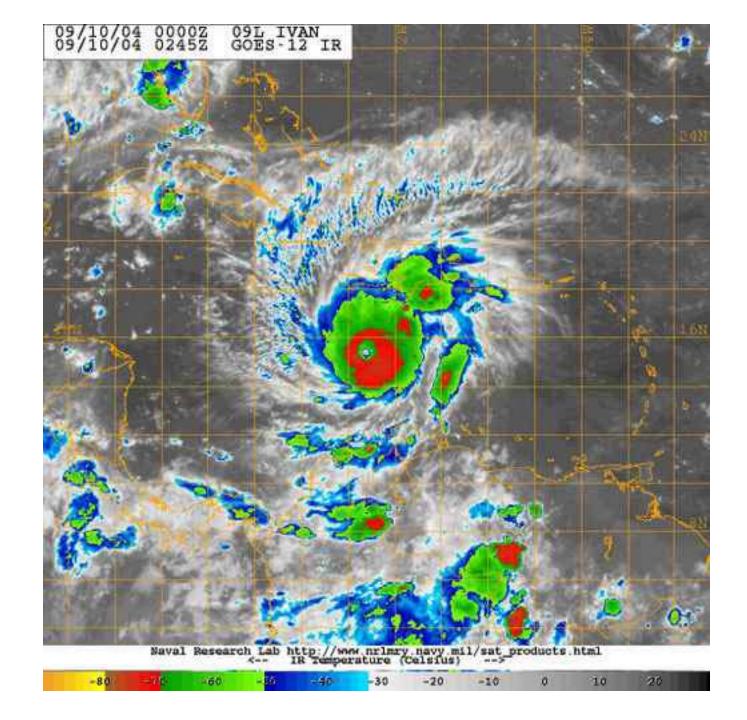


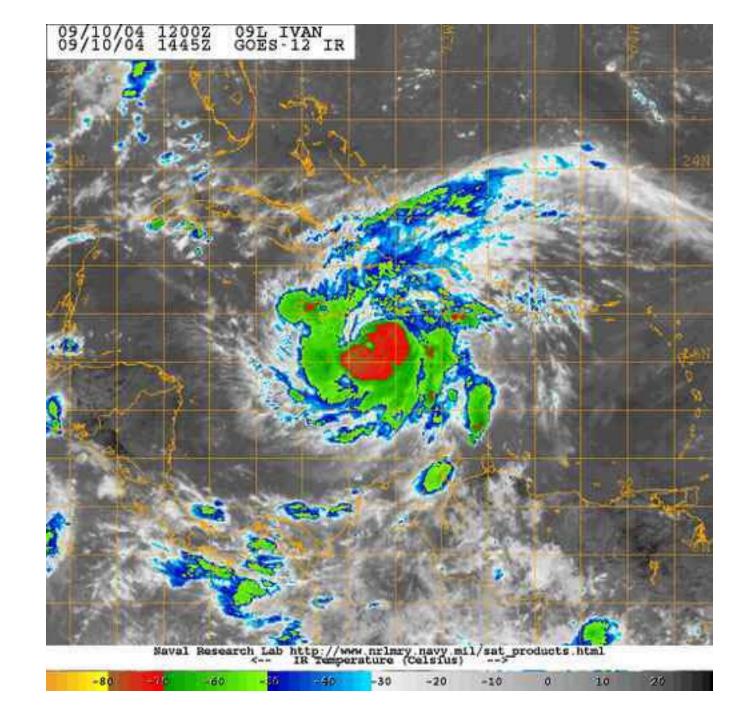


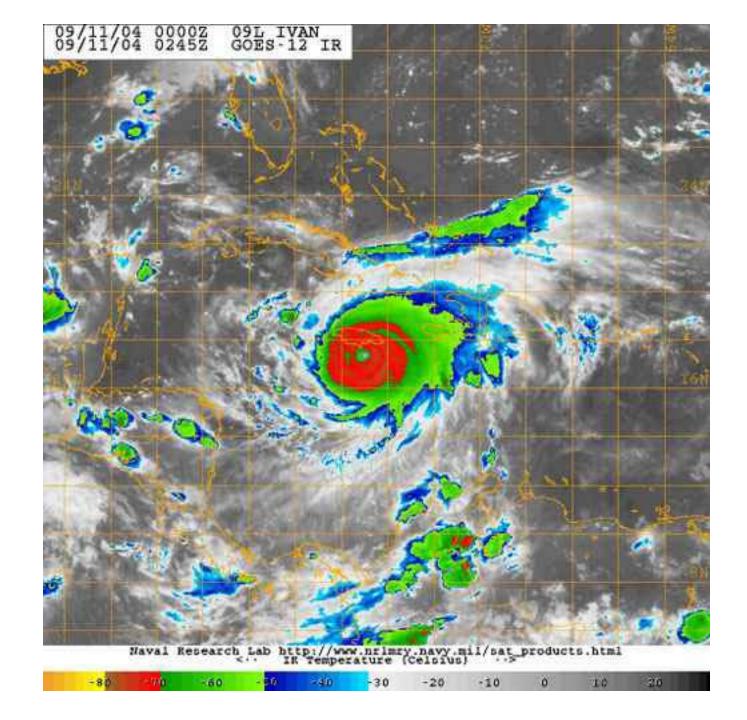


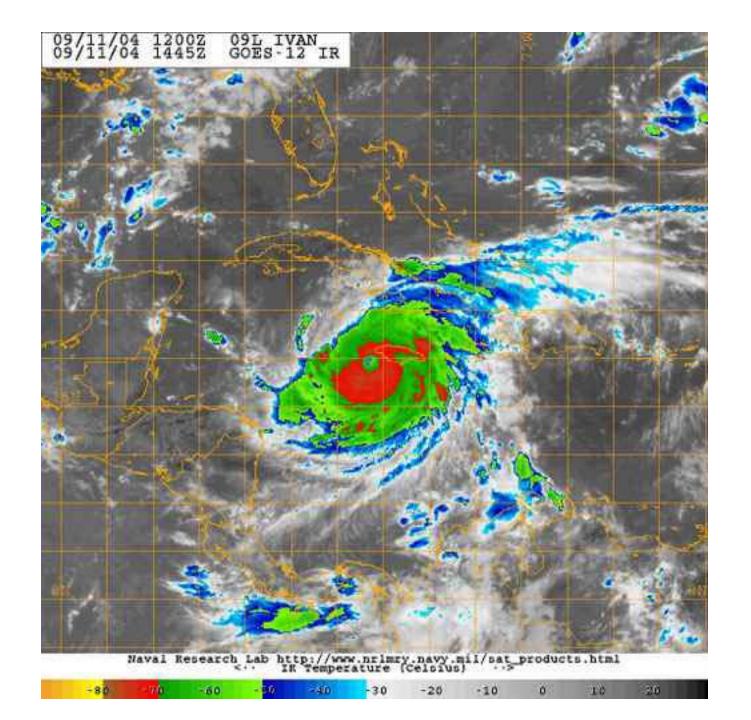


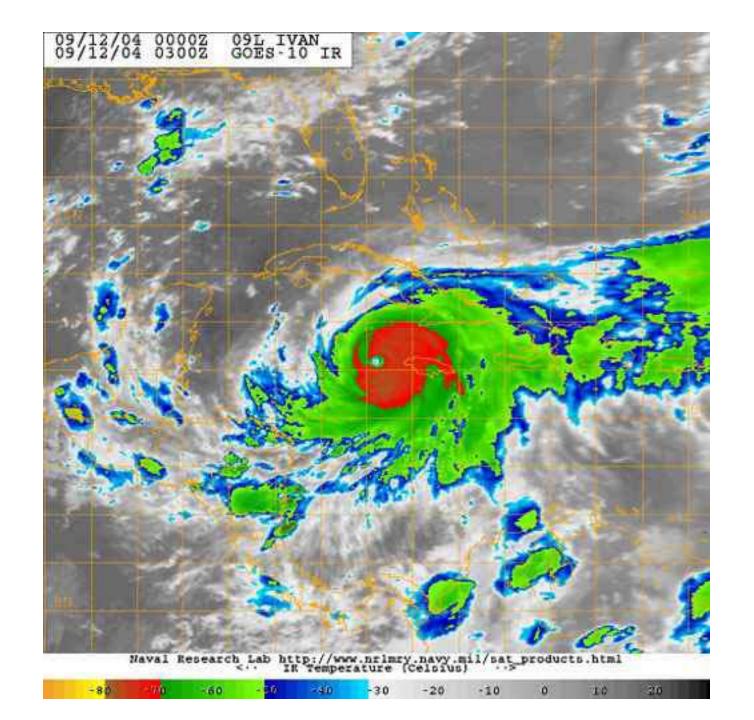




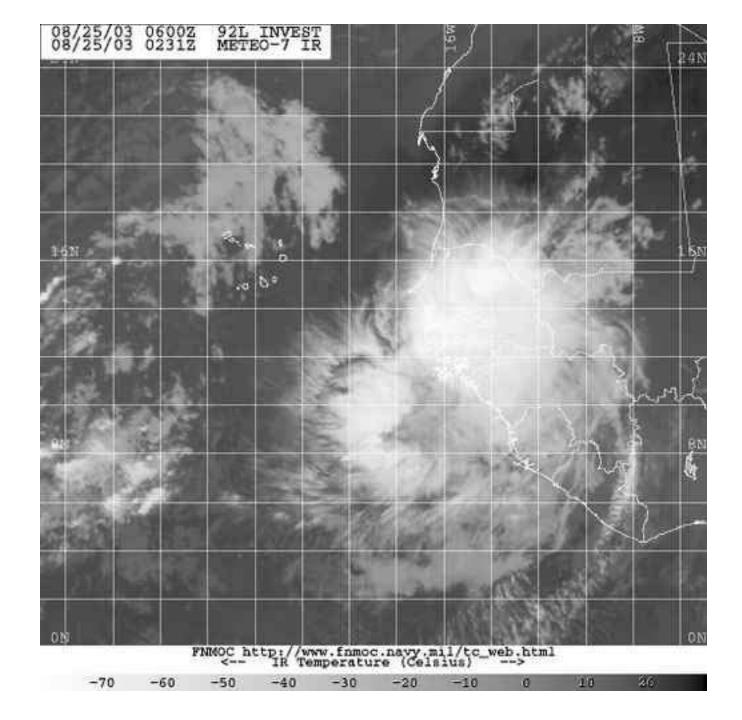


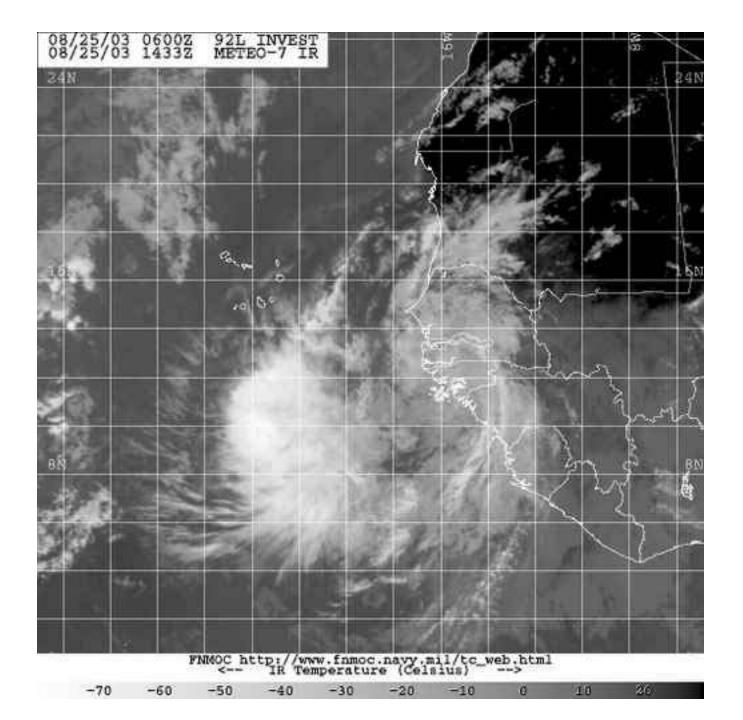


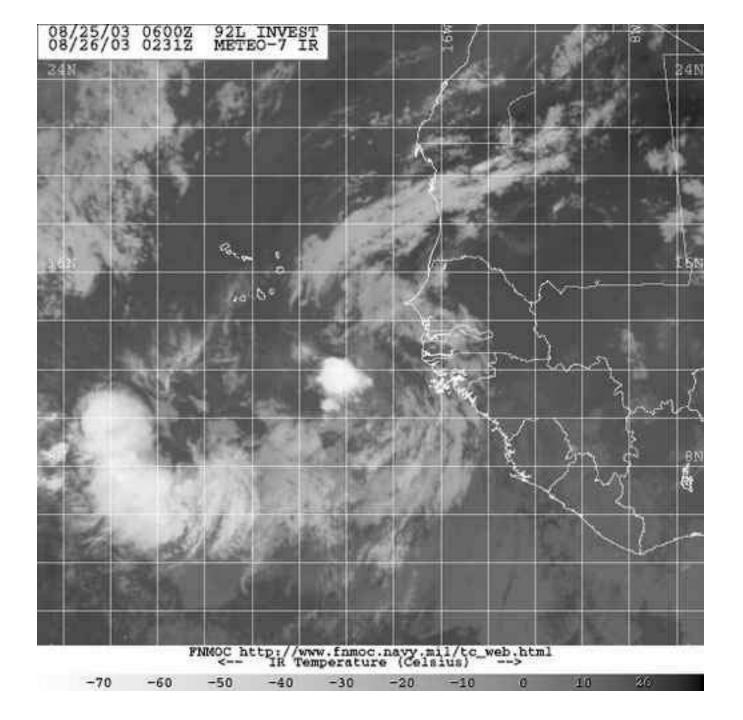


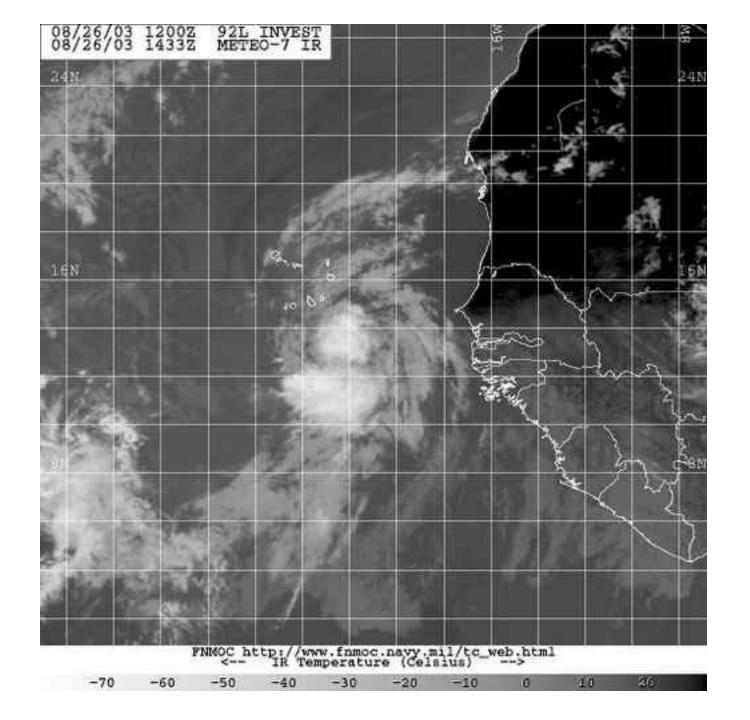


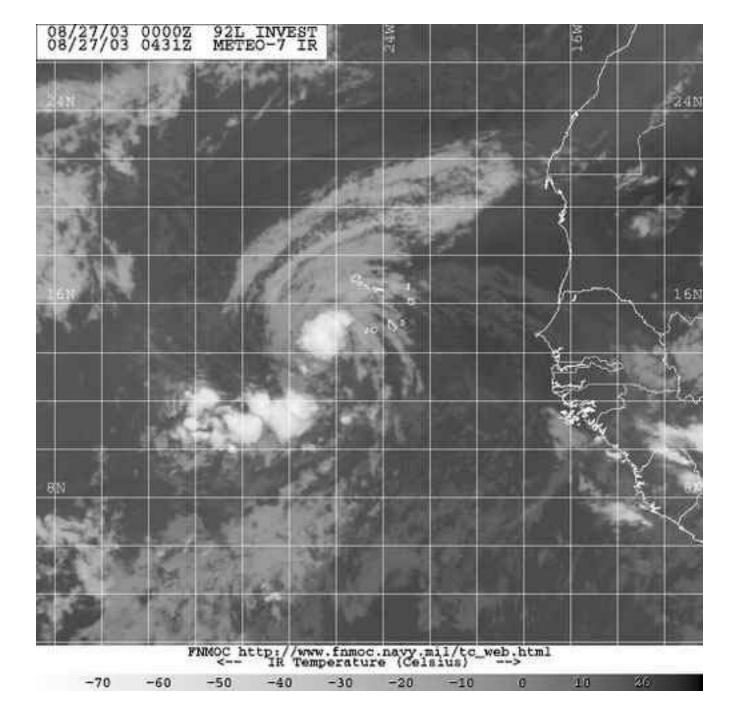
Genesis as seen in satellite and 850 hPa NCEP re-analysis data: Fabian, 2003

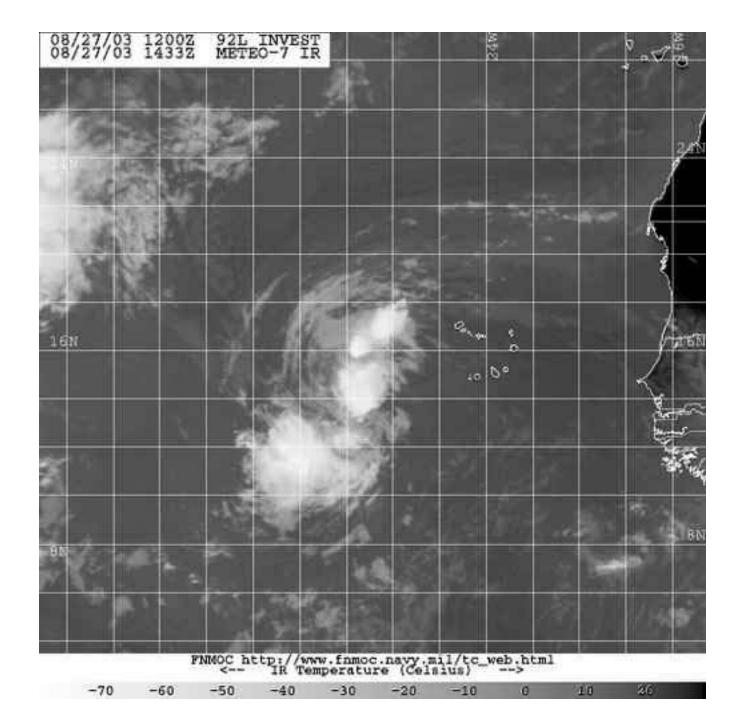


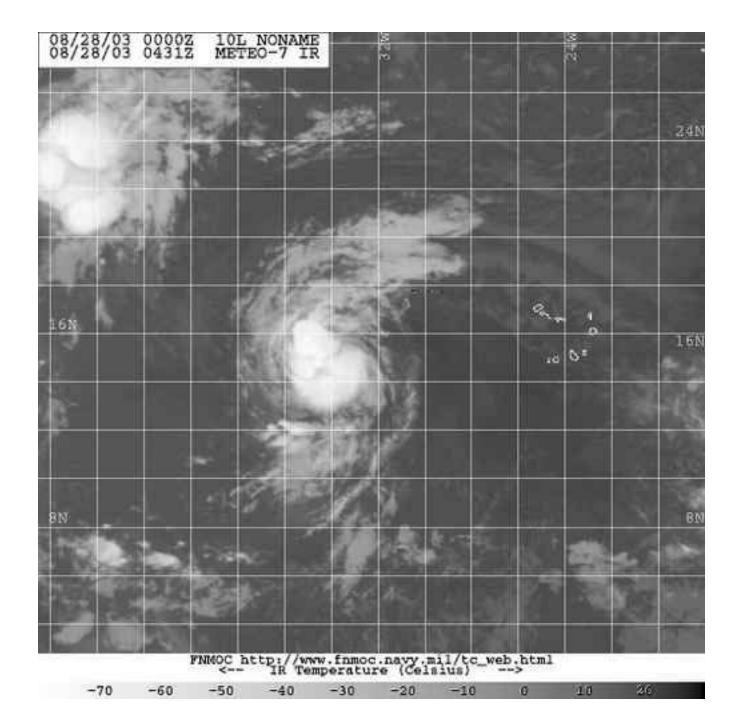


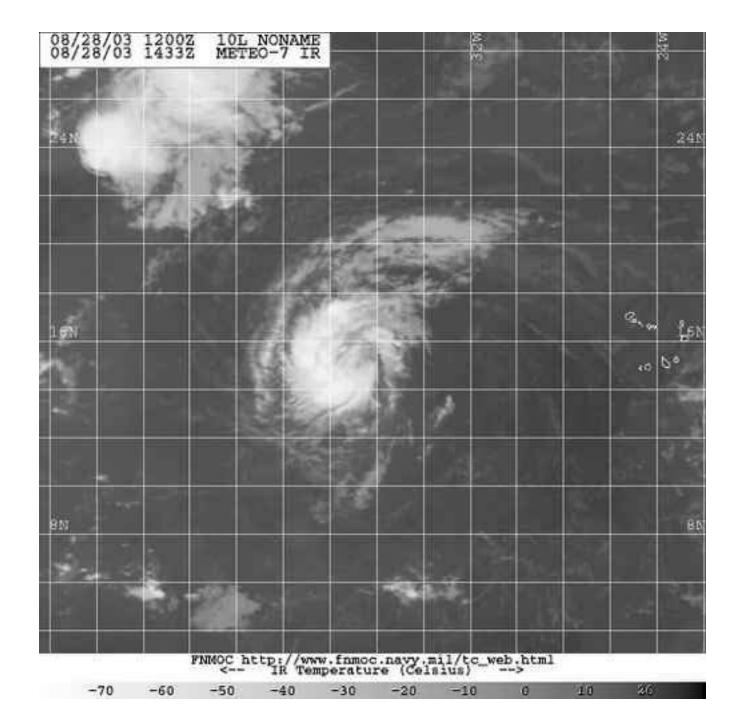


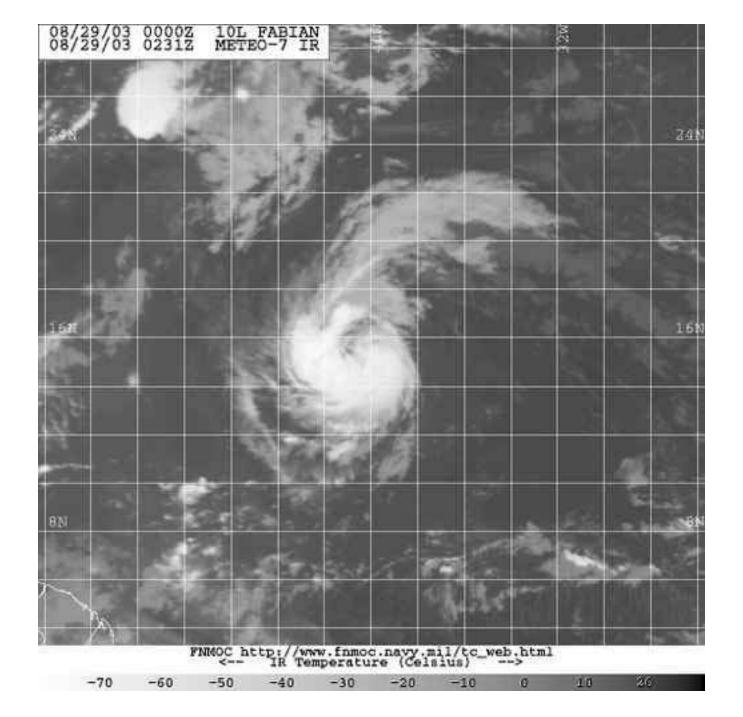


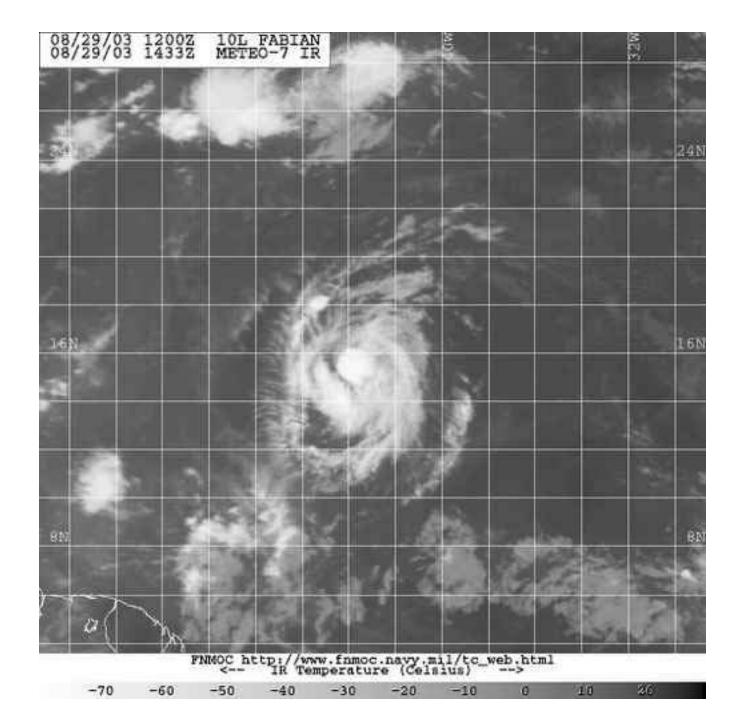


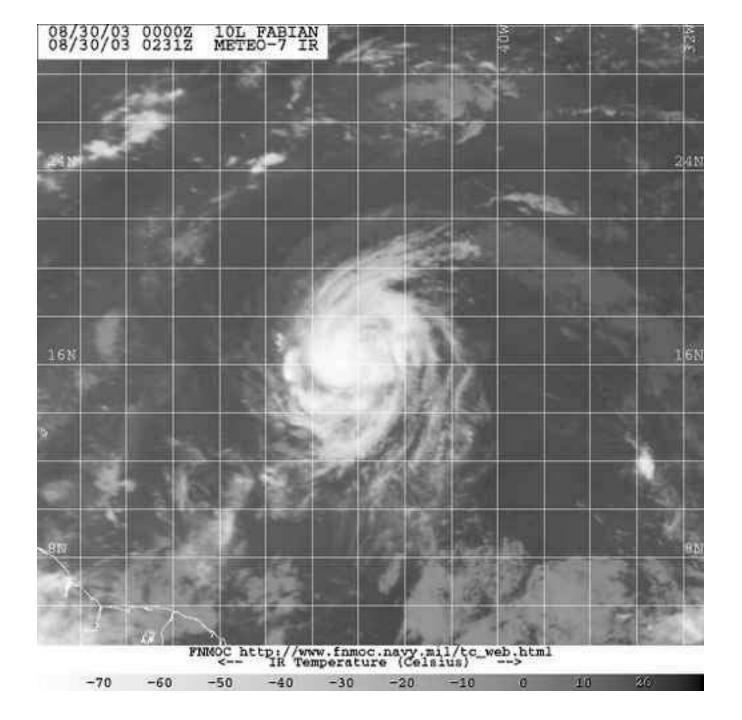


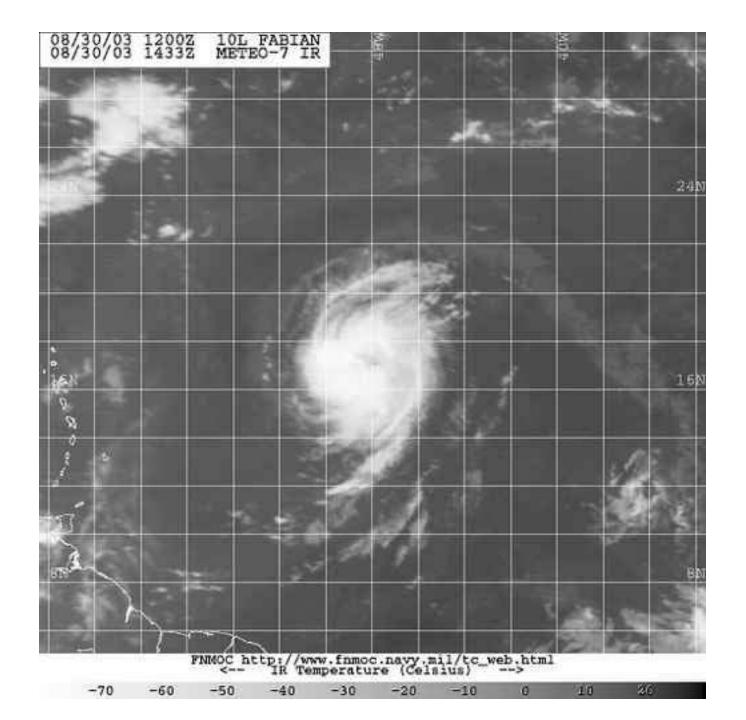


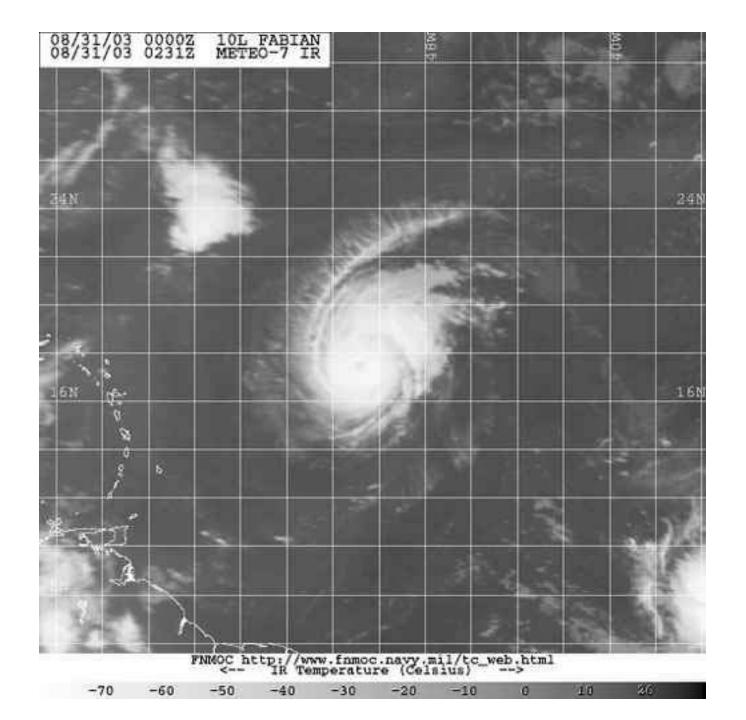






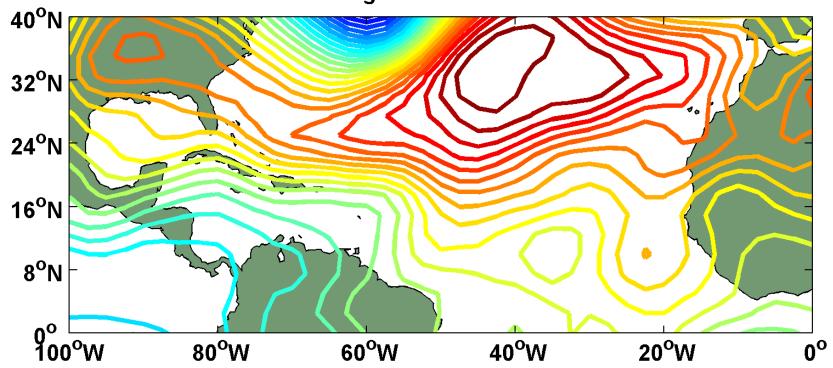


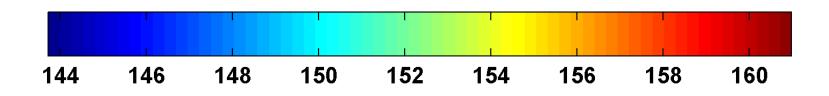




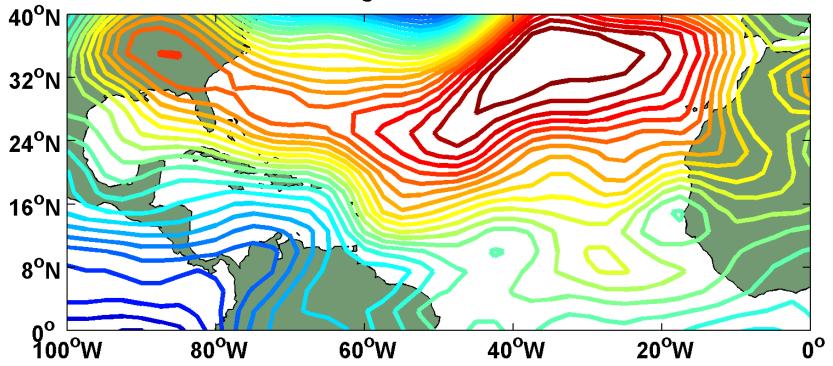


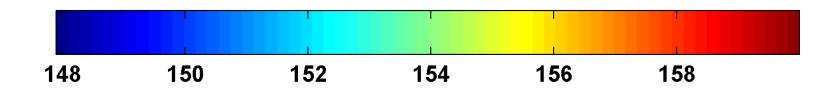
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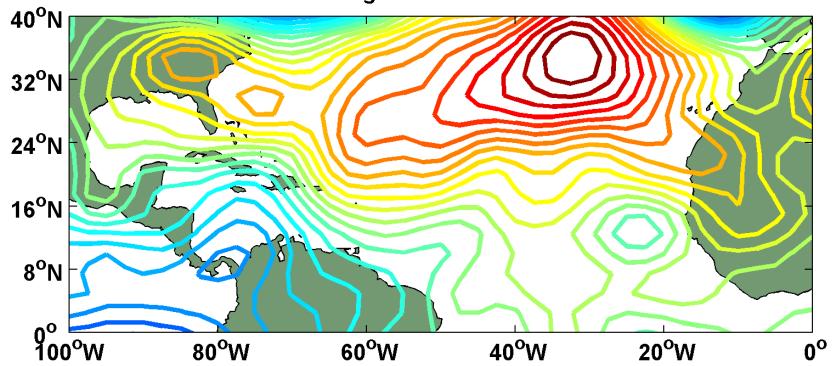


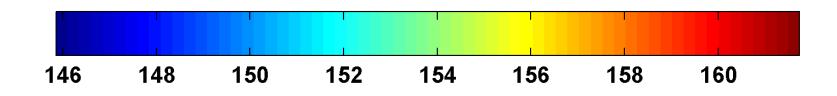
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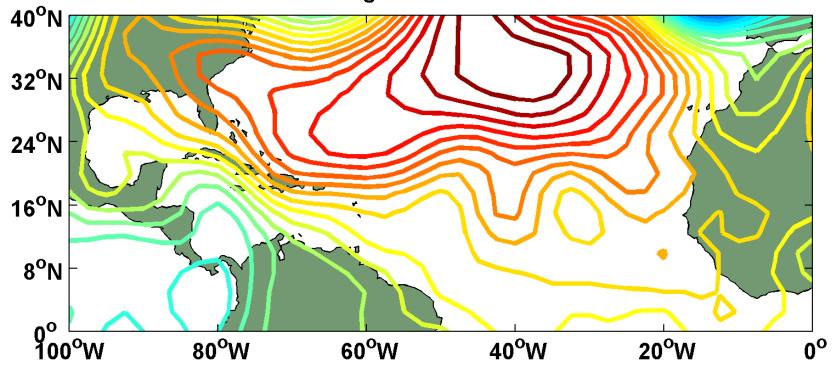


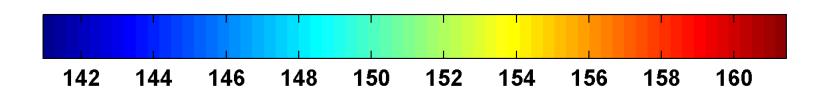
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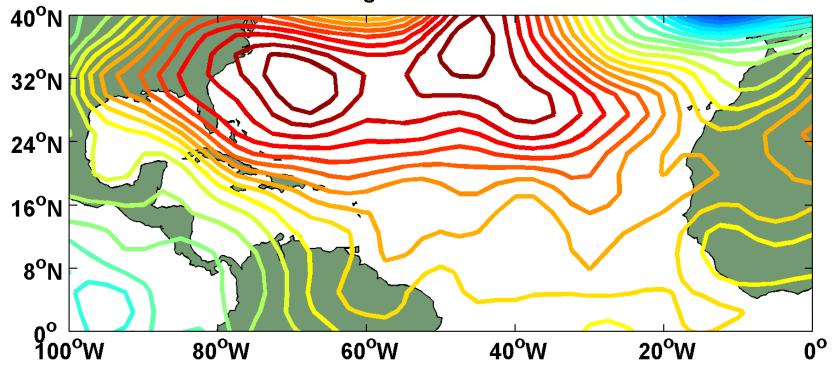


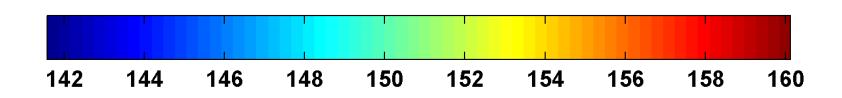




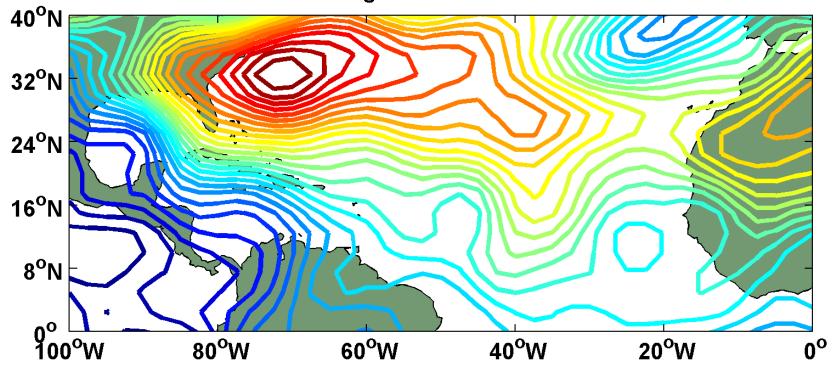


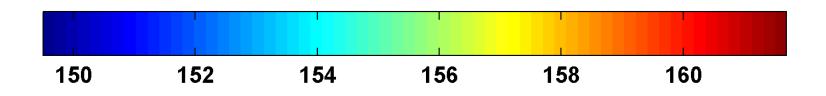
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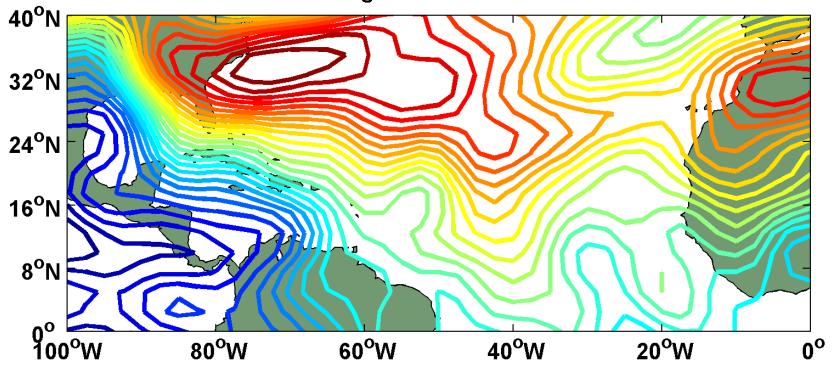


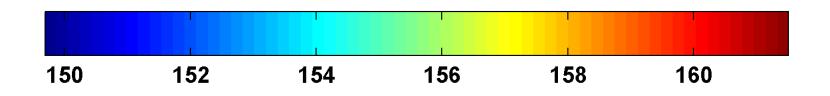
850 hgt 00 UTC0830 2003



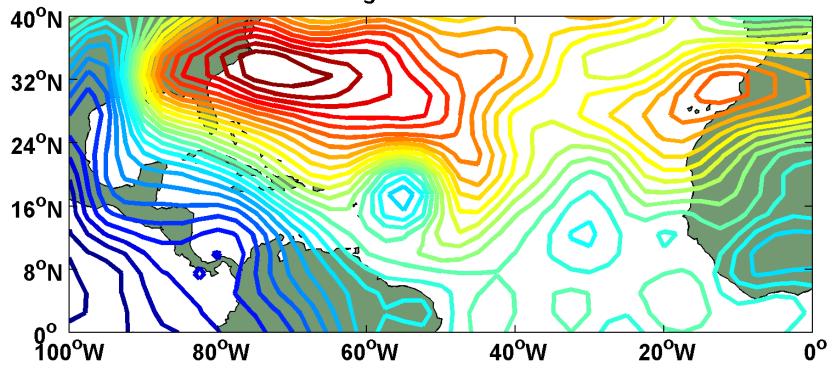


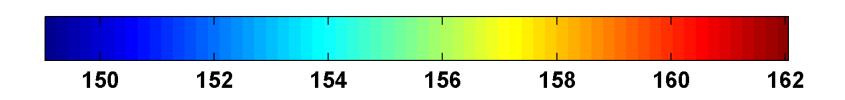






850 hgt 00 UTC0901 2003





A case study using rawinsonde and surface observations (From Hawkins and Rubsam, *Mon. Wea. Rev.*, 1968)

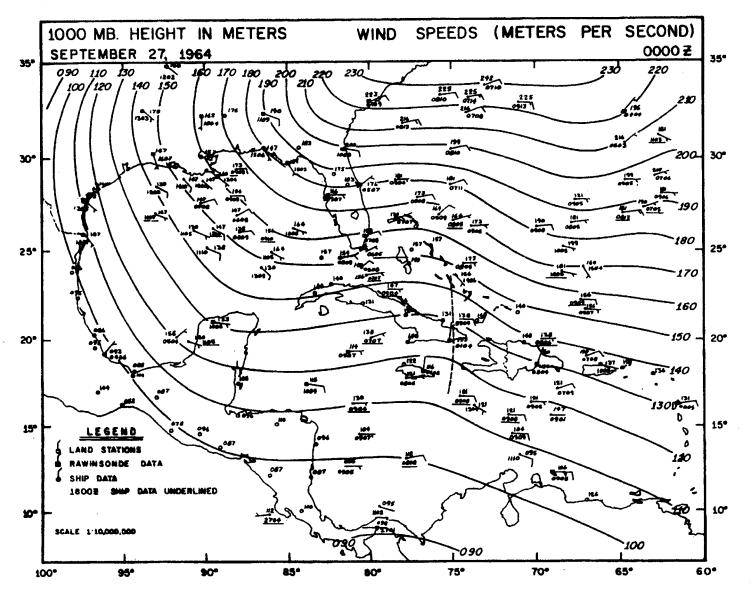


FIGURE 5.—1000-mb. contours for 00 gmr, September 27 showing a barely perceptible easterly wave over eastern Cuba, the first low level evidence of the disturbance that became hurricane Hilda.

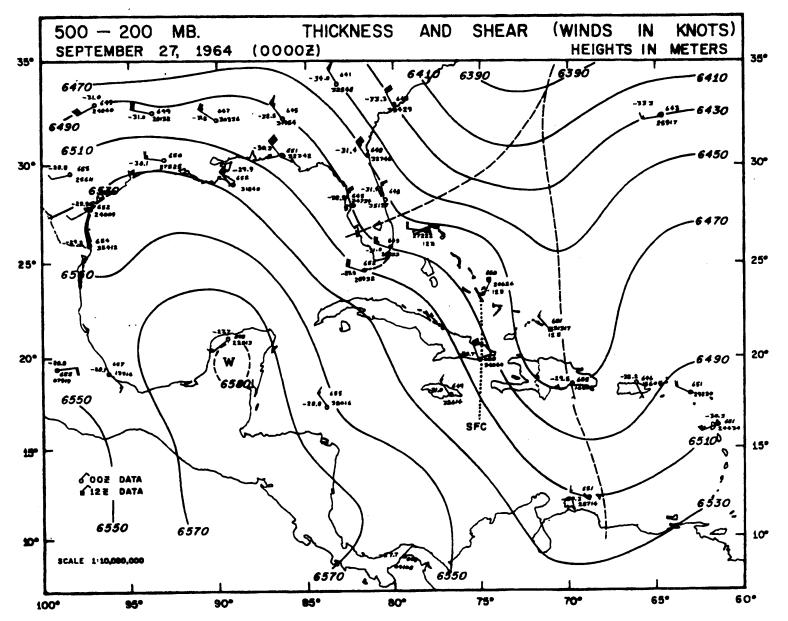


FIGURE 6.—The surface trough (dotted) almost midway between the trough-ridge system in the 500- to 200-mb. thickness.

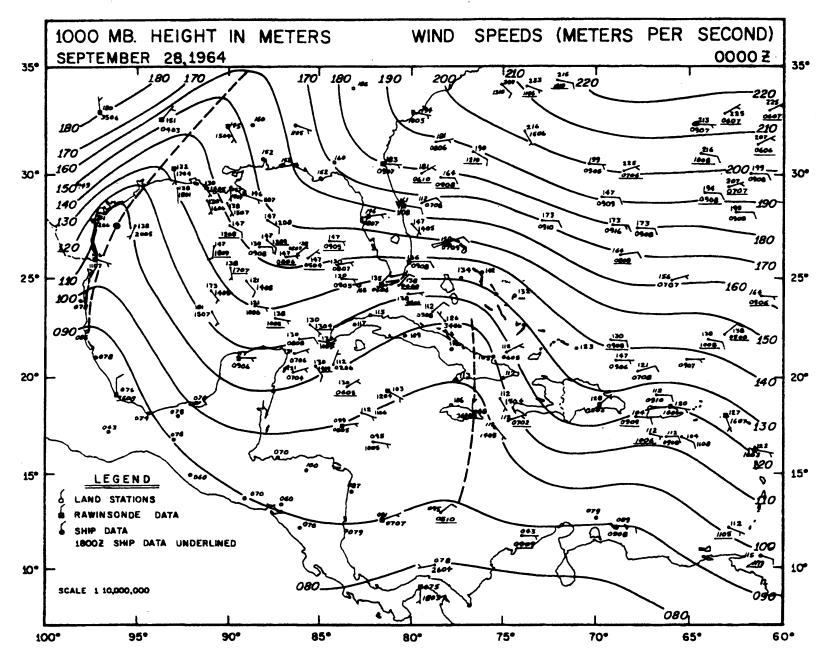


FIGURE 8.—Developing easterly wave over eastern Cuba and Jamaica.

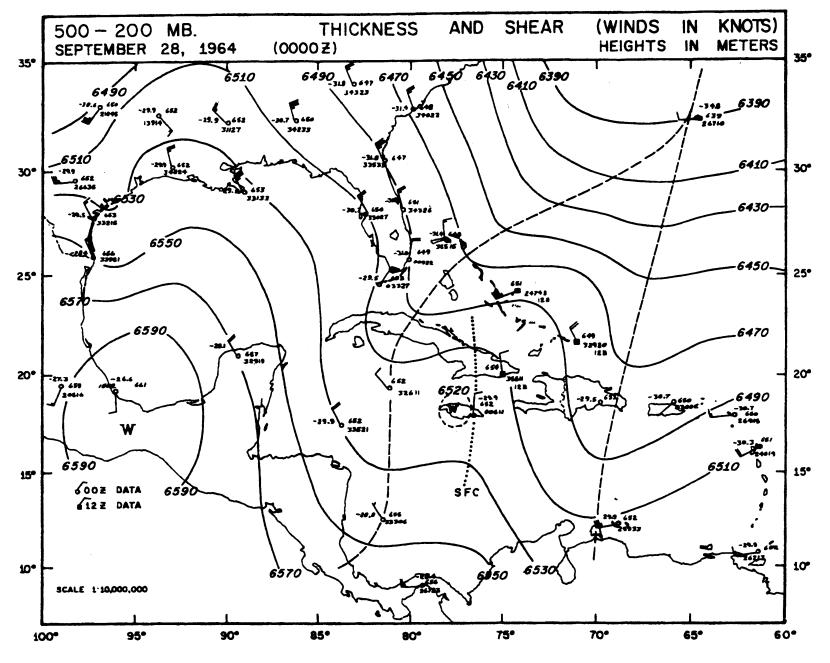


FIGURE 9.—Surface trough (dotted) now located in the 500- to 200-mb. thickness ridge.

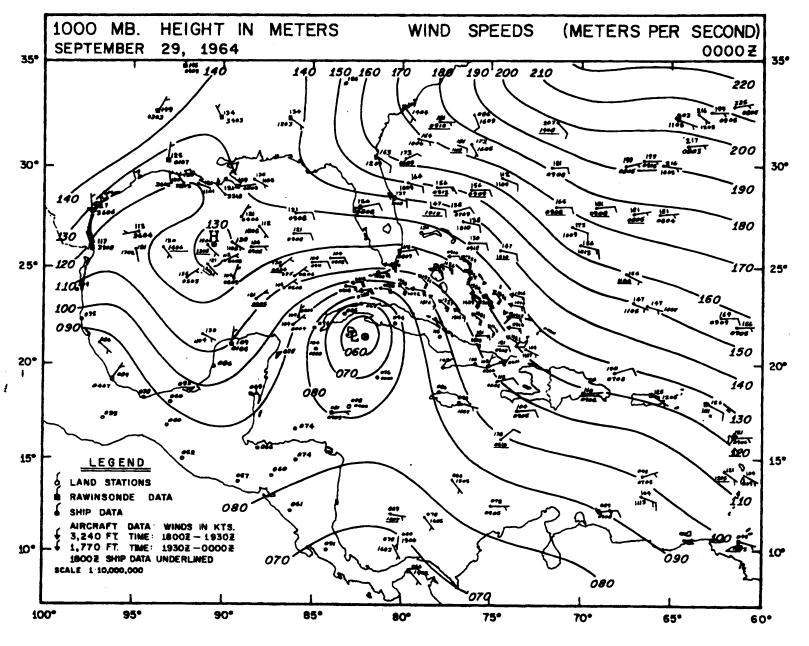


FIGURE 12.—Tropical depression (Hilds) at 00 GMT, September 29.

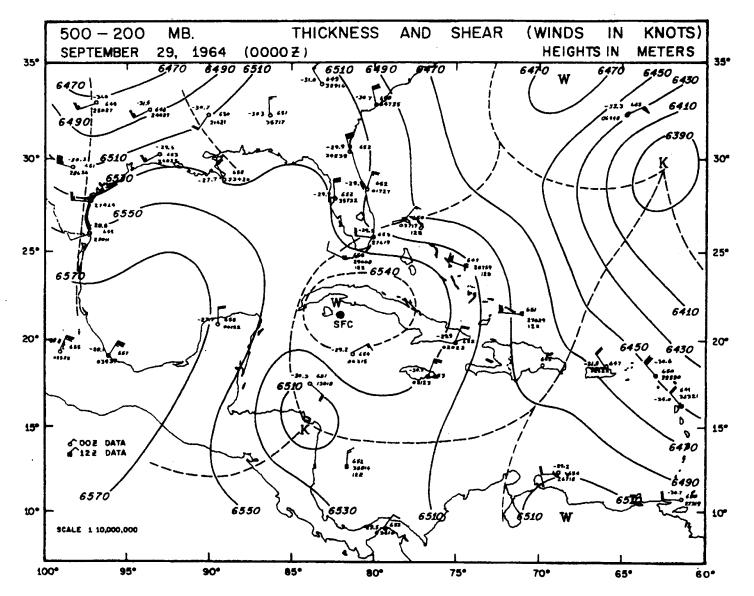


FIGURE 14.—Further marked warming in the 500- to 200-mb. layer. Shear winds clearly define the presence of anomalous warmth.

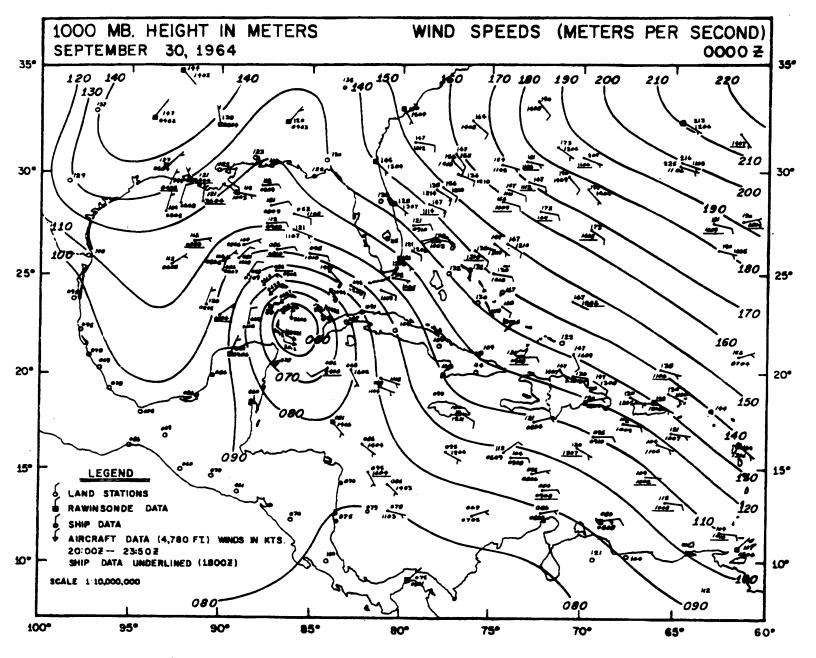


FIGURE 17.—Tropical storm Hilds at 00 gmr, September 30.

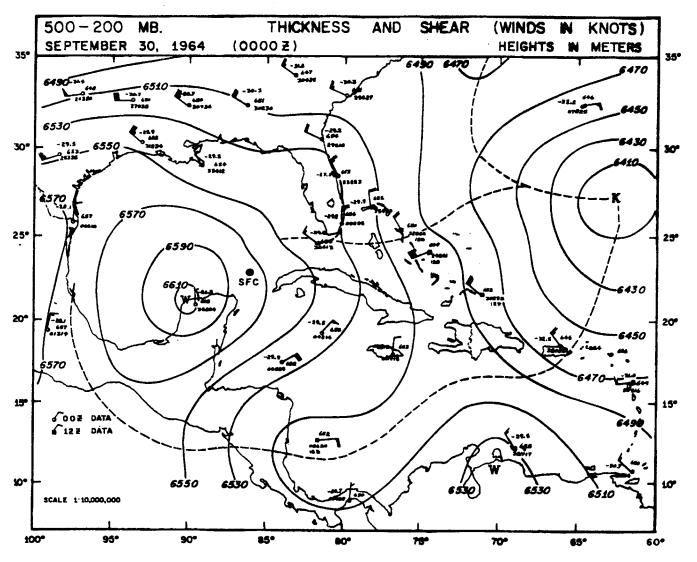


Figure 21.—Lack of data prohibited more detailed analysis of thickness over the storm area, but there is little doubt of the larger scale pattern.

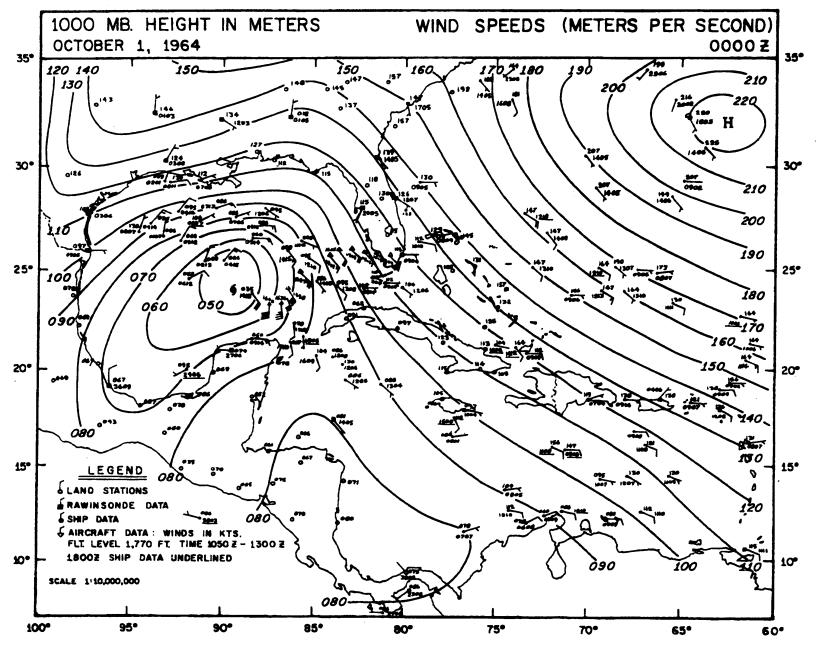


FIGURE 23.—Hurricane Hilds deepening rapidly near the center of the Gulf of Mexico.

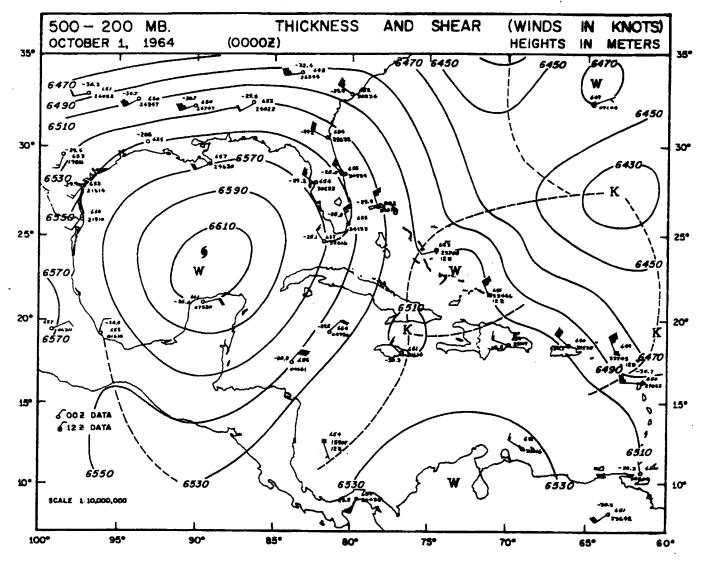


FIGURE 25.—The only shear winds that do not reflect the prevailing upper level warm pool accompanying Hilda over the Gulf are the weak winds at Vera Crus, Mexico.

Tropical Experiment in Mexico (TEXMEX) July 1 – August 10 1991

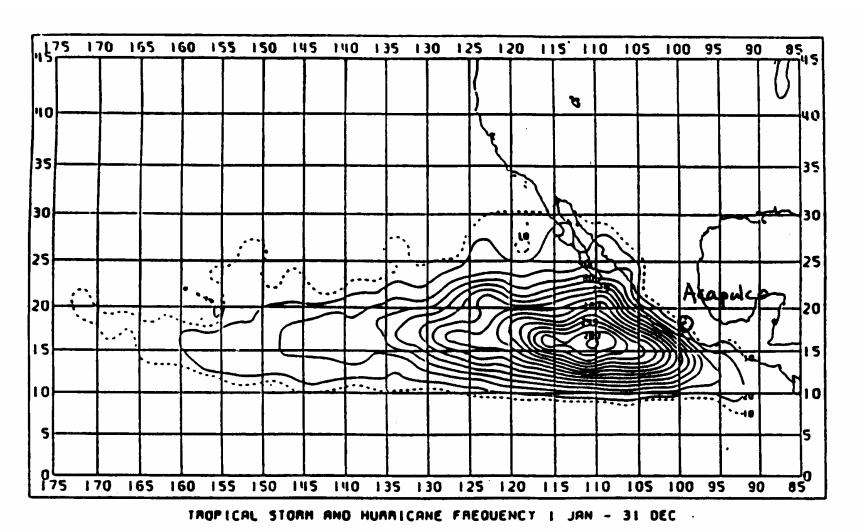


Fig. 3 Number of storms passing within 75 n. mi. of points in the eastern North Pacific region. The numbers of storms passing through a grid of overlapping equal-area circles was counted for the years 1966-1984 and normalized to give the frequency for a 100-year period (provided by C. Neumann).

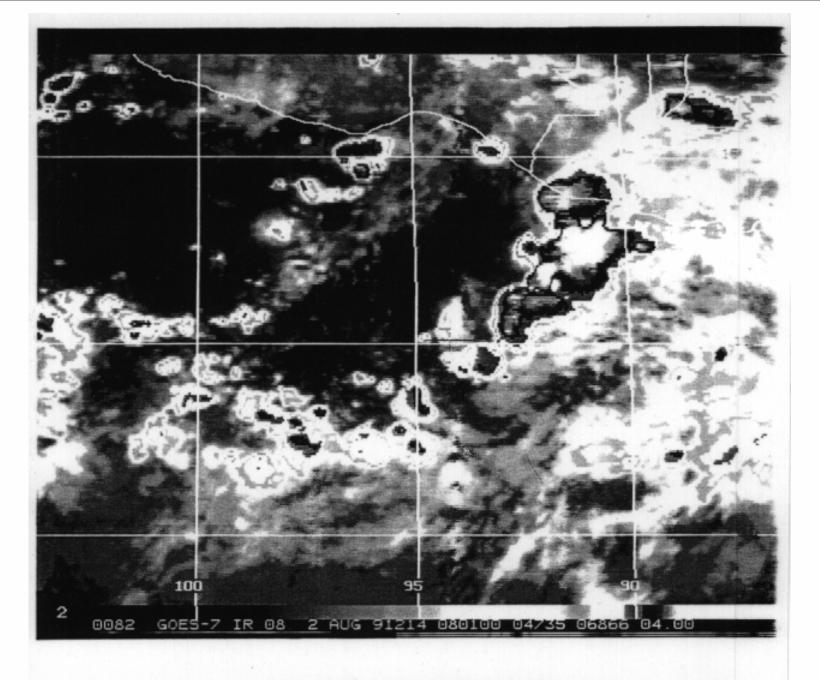


Figure 50: IR Satellite picture for 08:01 UTC 2 August.

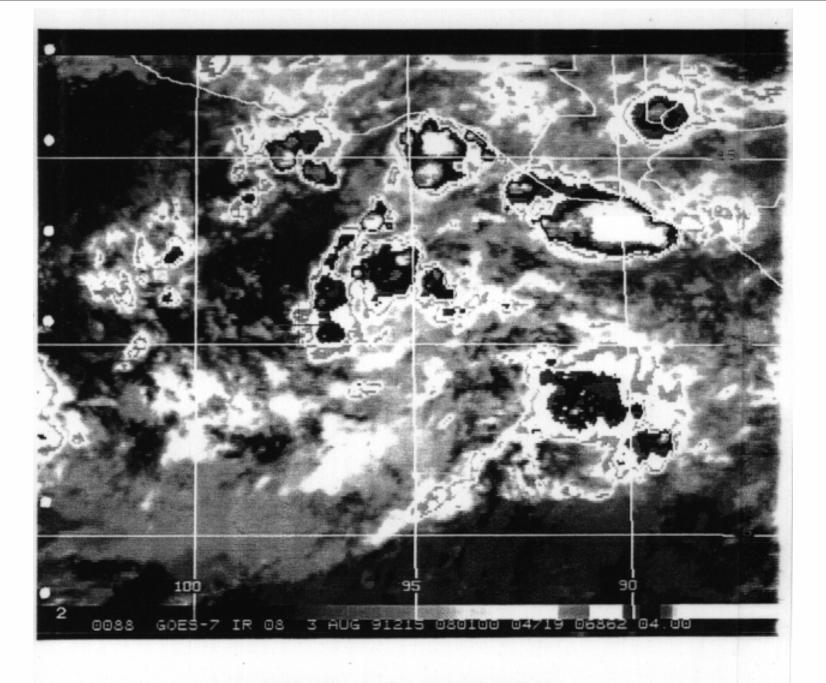


Figure 51: IR Satellite picture for 08:01 UTC 3 August.

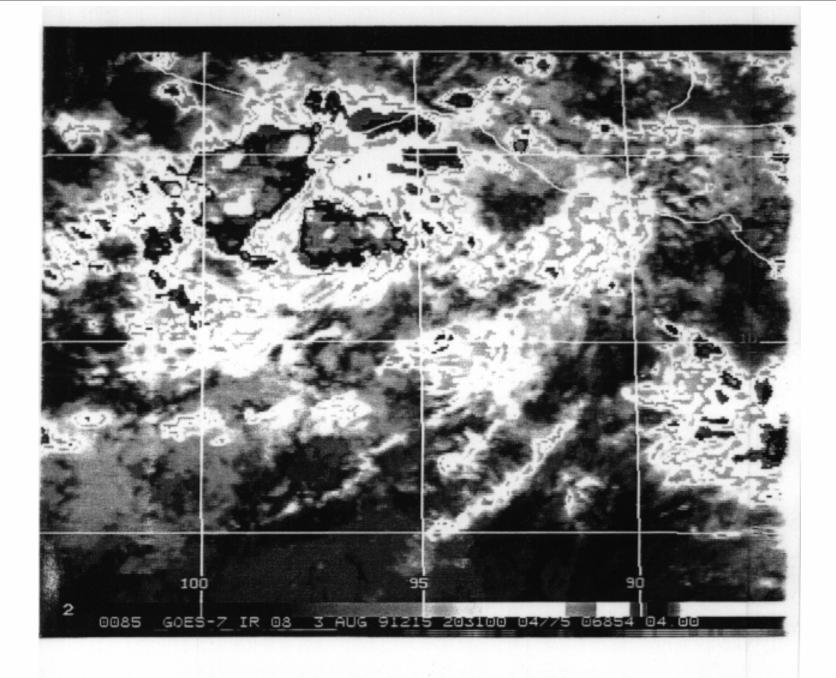


Figure 52: IR Satellite picture for 20:31 UTC 3 August.

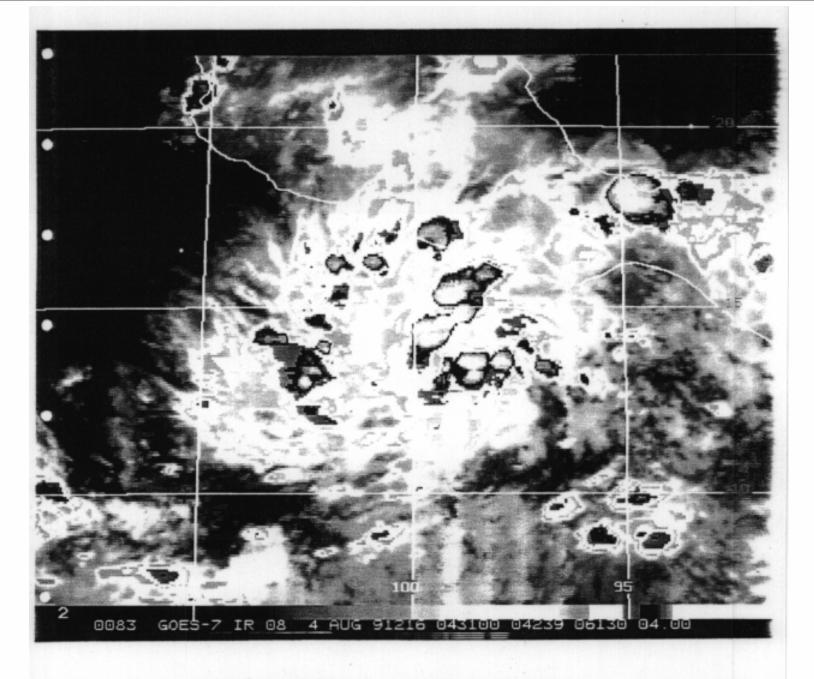


Figure 53: IR Satellite picture for 04:31 UTC 4 August.

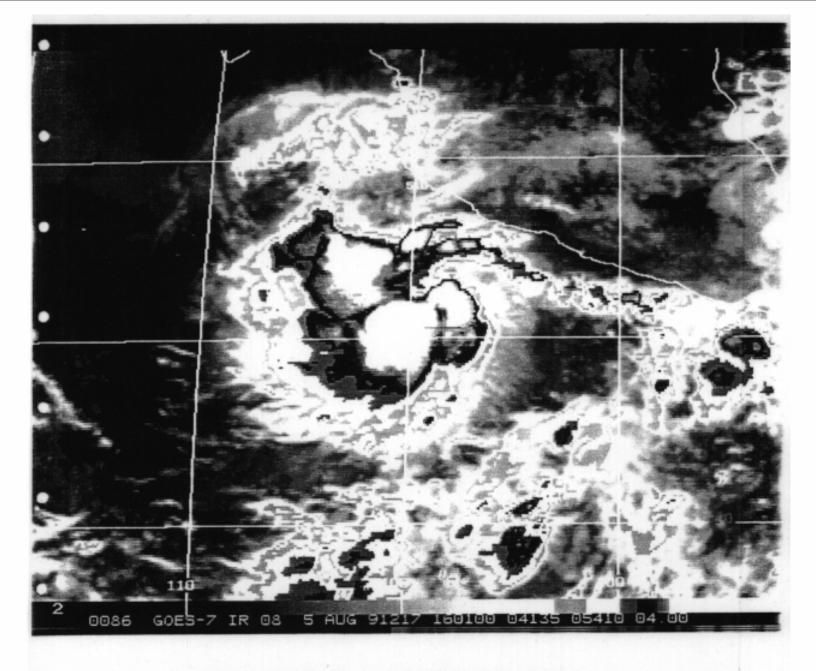
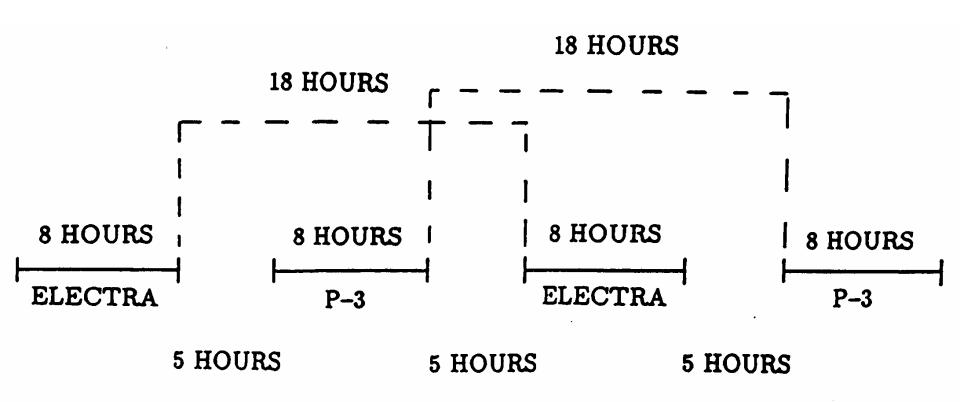
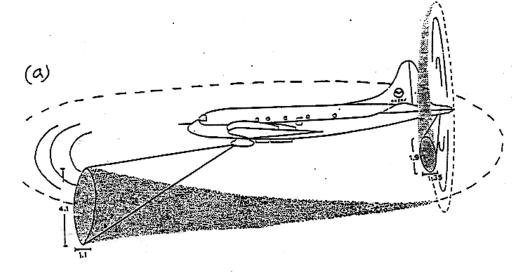


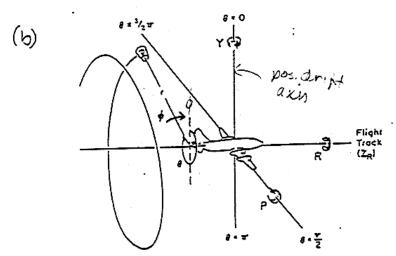
Figure 55: IR Satellite picture for 16:01 UTC 5 August.

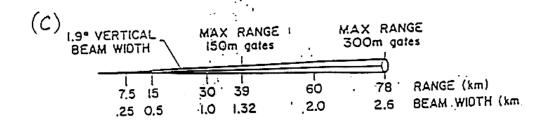
Flight Scheduling





WP-3D radars





Tail Doppler Scanning Strategy

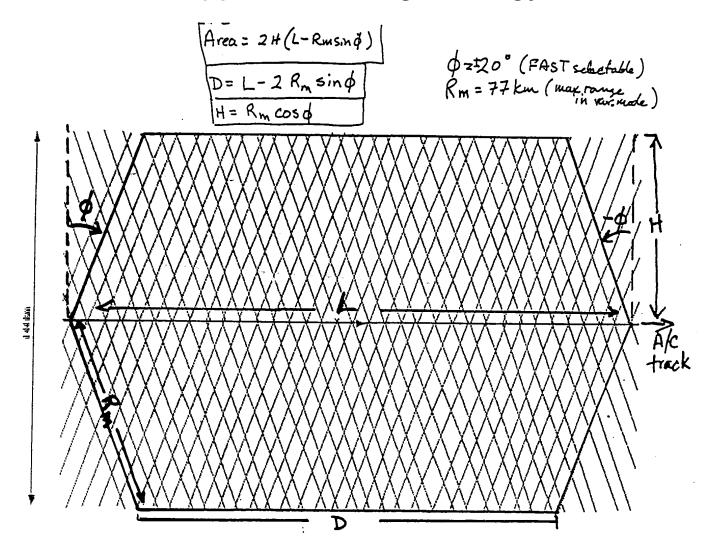
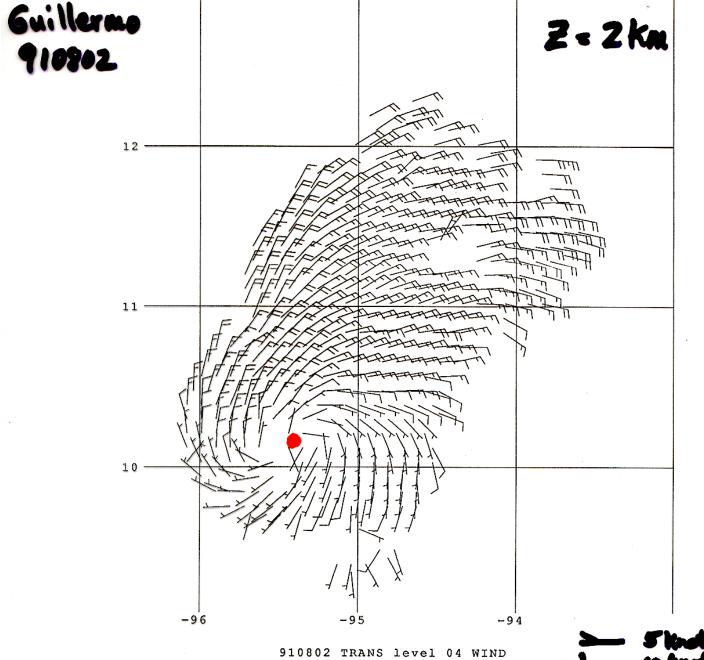


Fig. IV-2. Horizontal psuedo-dual Doppler coverage for tail Doppler radar for FAST scanning.



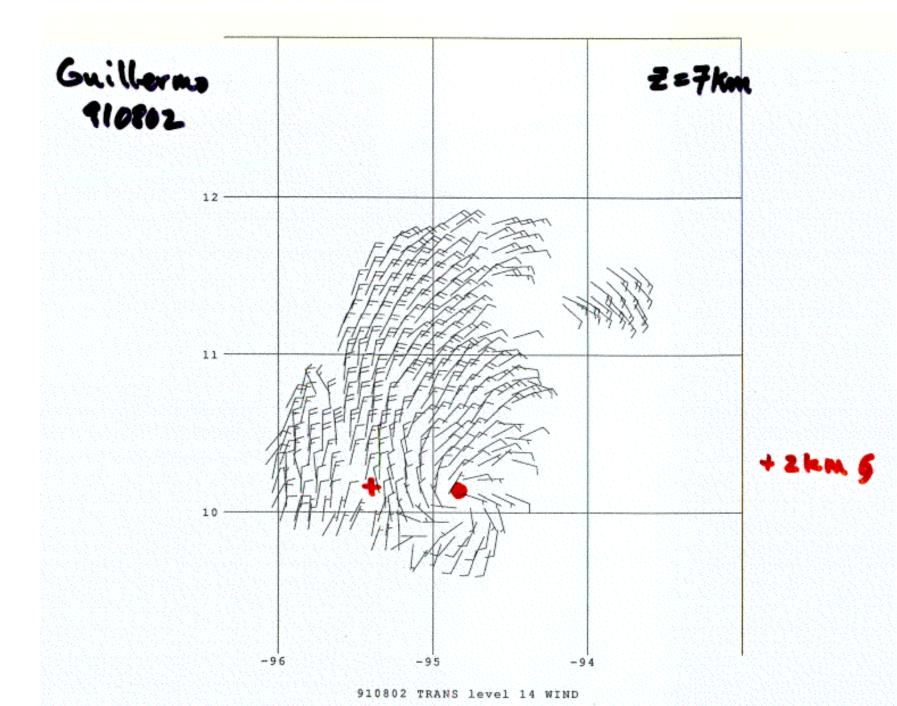


Figure 66: The 700 hPa pattern for flight 910802H.

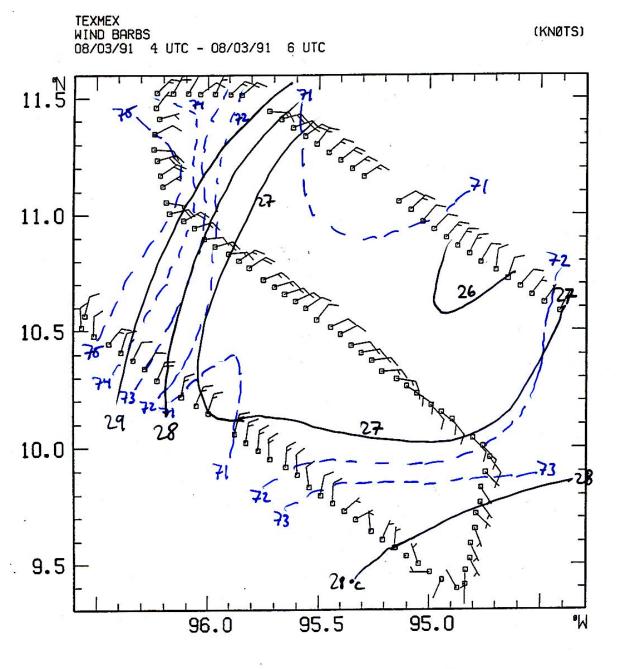


Figure 67: The 975 hPa pattern for flight 910802H.

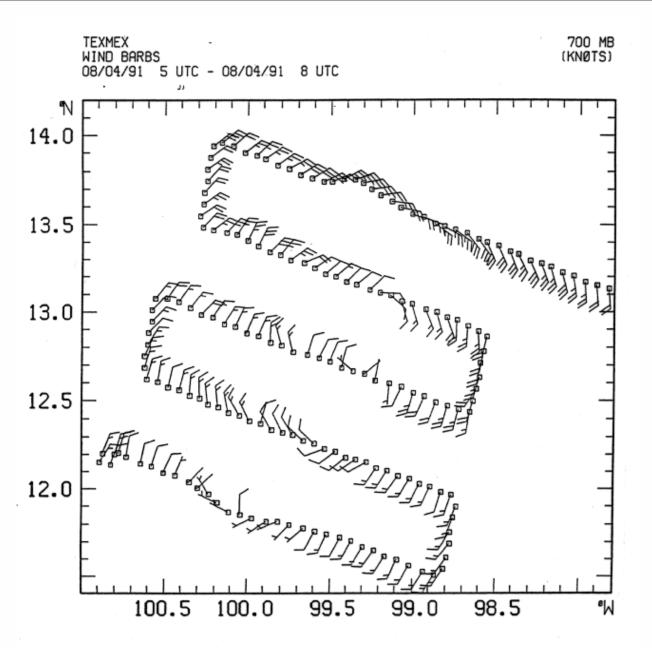


Figure 71: The 975-hPa pattern for flight 910804H.

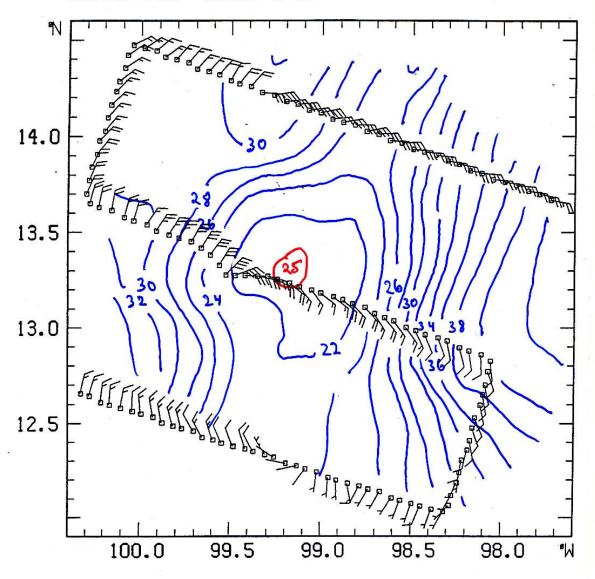


Figure 70: The 700 hPa pattern for flight 910804H.

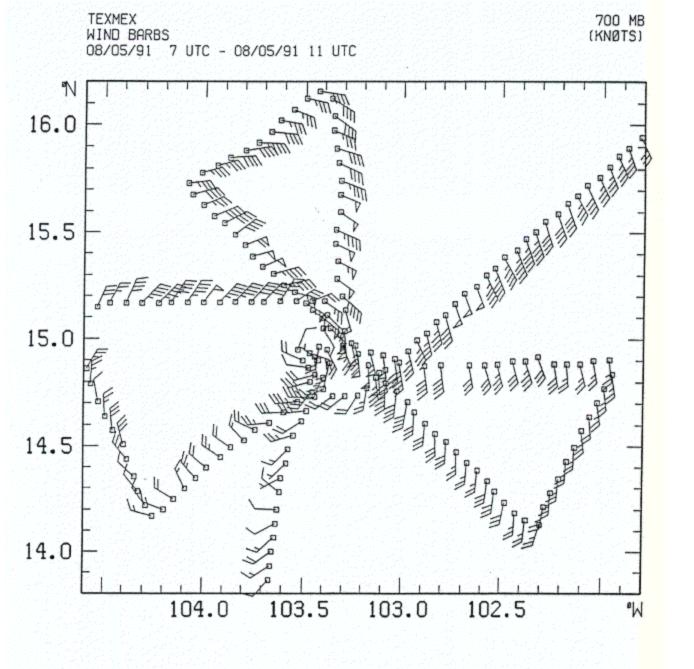


Figure 74: The 700 hPa pattern for flight 910805H.

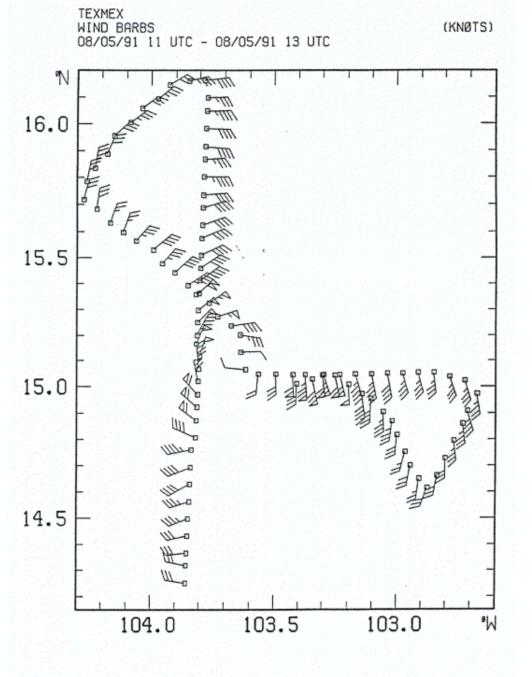
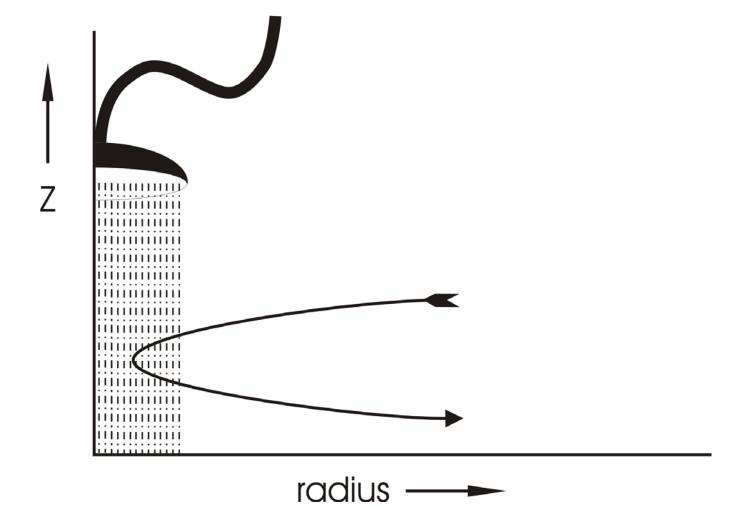
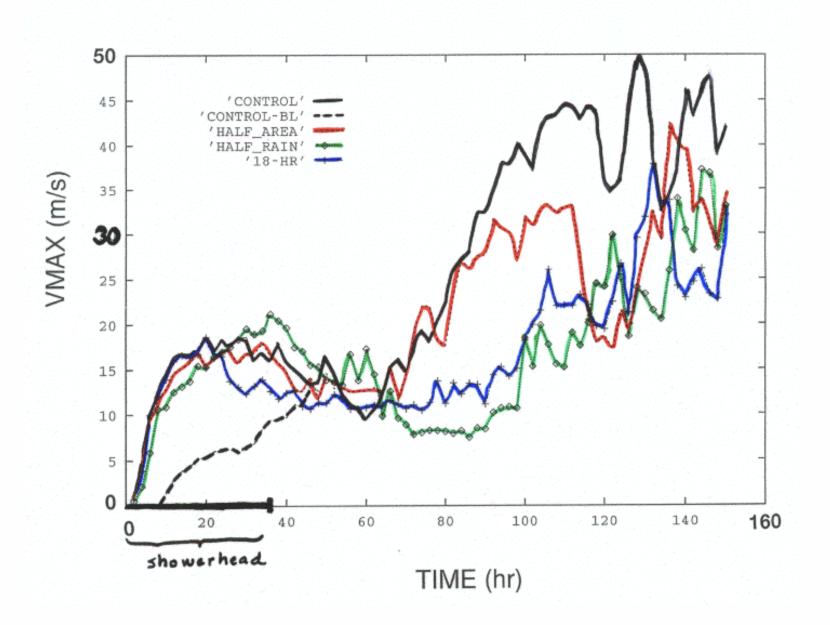


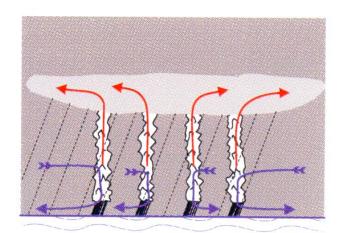
Figure 75: The 975 hPa pattern for flight 910805H.



Maximum tangential wind velocity

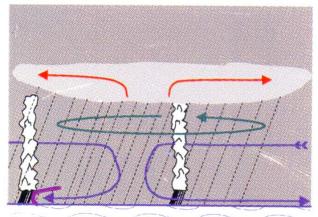


1. TRIGGERING



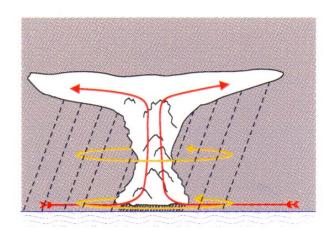
- Formation of long-lived mesoscale strattform anvil
- Appears to require large-scale ascent in the upper troposphere
- Reduction of subcloud layer entropy by downdrafts

2. Gestation



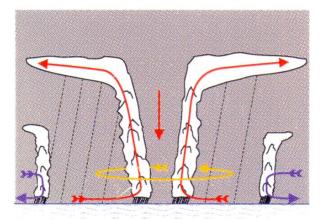
- Light to moderate strattform rain;
 little deep convection, except at periphery
- Formation of middle tropospheric mesoscale cyclone cold core in the lower troposphere
- · High relative humidity develops in core
- Subcloud layer entropy recovers

3. Ignition



- New episode of convection that is free of downdraft forms near core
- Strong surface Inflow, strong surface heat fluxes
- Carnot engine switched on

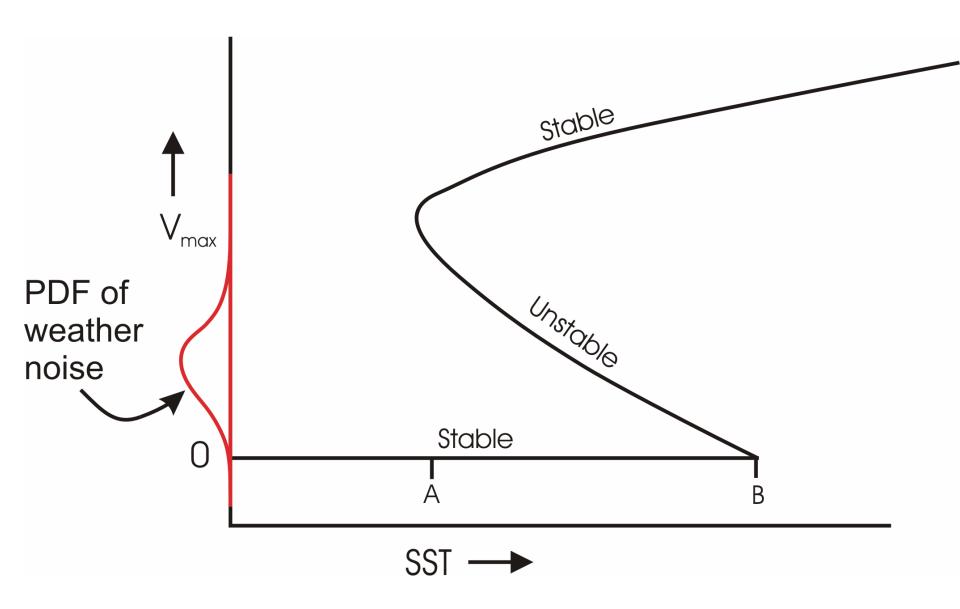
4. Intensification



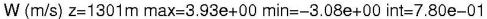
- 5. Maturity
- 6. Dissipation

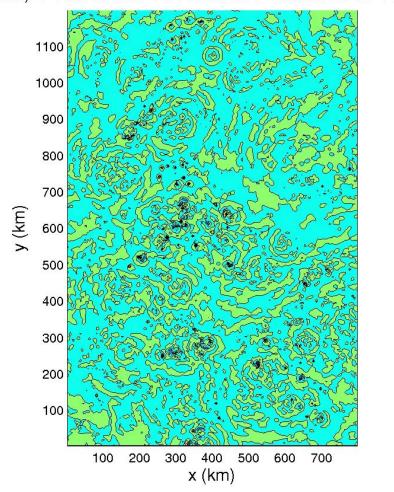
Hypothesis

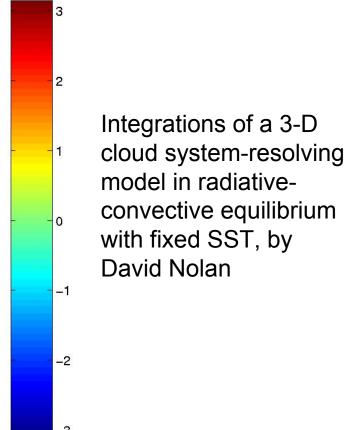
 Tropical cyclones result from a subcritical bifurcation of the normal state of the tropical atmosphere



$SST = 30^{\circ}C$

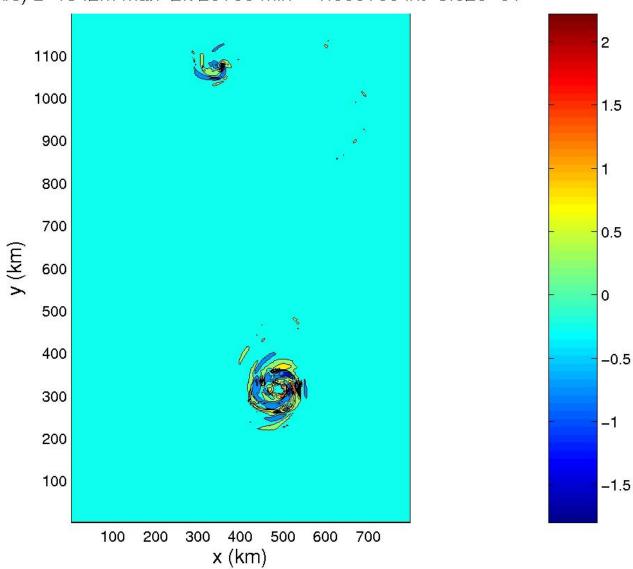






$SST = 35^{\circ}C$

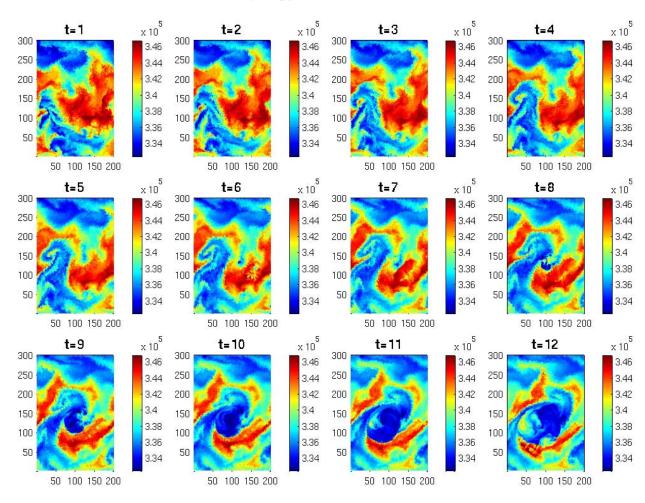
W (m/s) z=1342m max=2.72e+00 min=-1.80e+00 int=5.02e-01



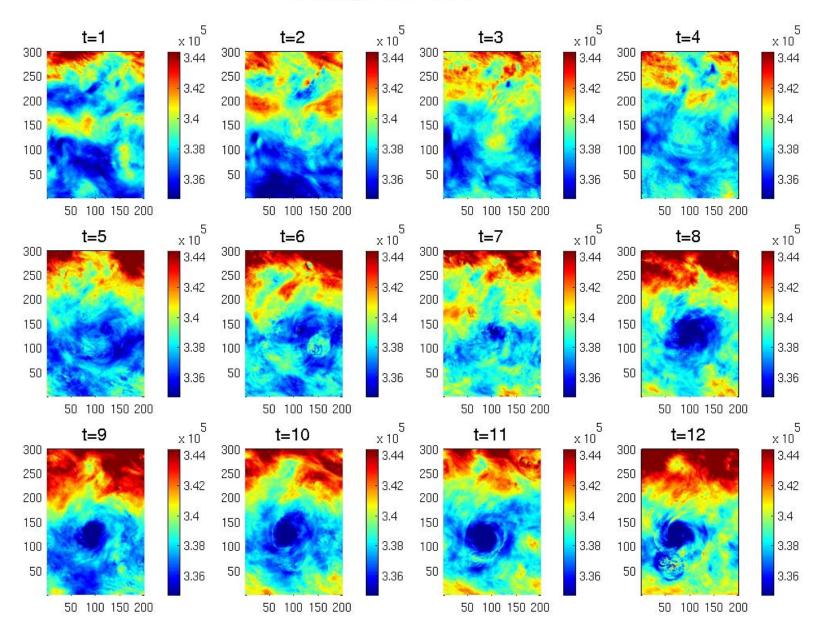
Initial development begins as mesoscale downdraft:

h(J/kg) at z=24.6

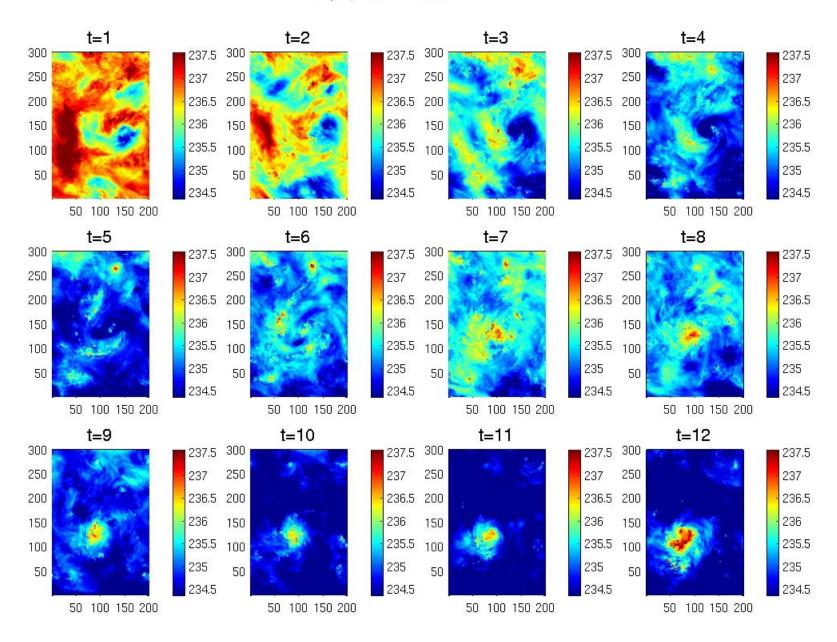
Moist static energy at 25 m



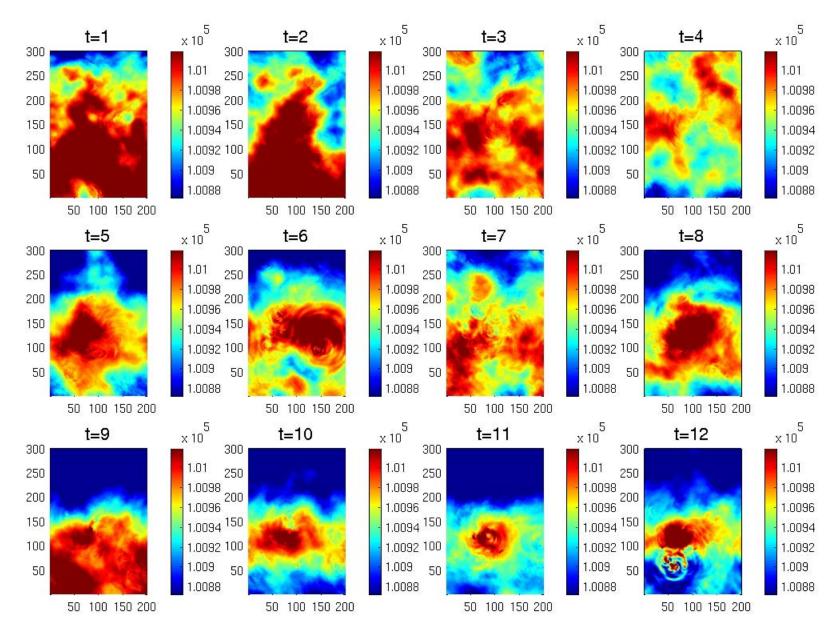
h*(J/kg) at z=3810



T(K) at z=9675

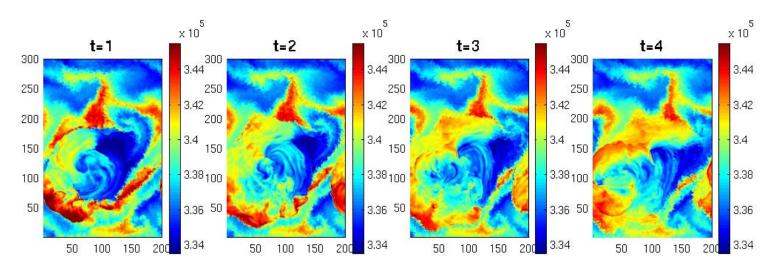


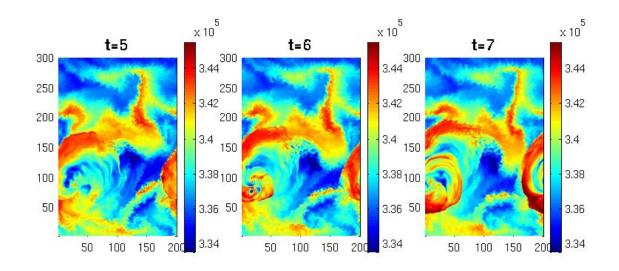
P(Pa) at z=24.6



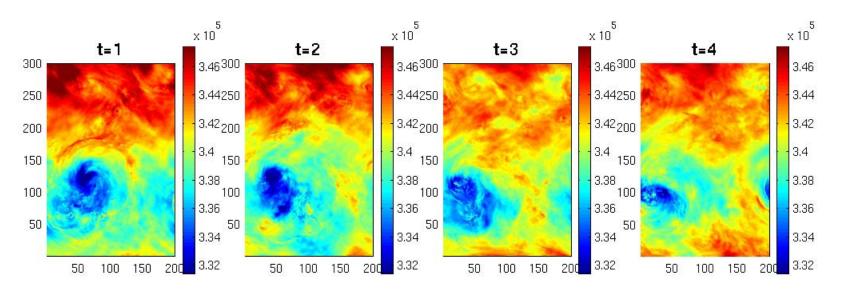
Intermediate times

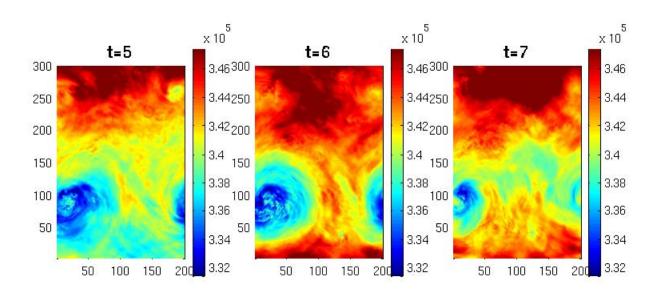
h(J/kg) at z=24.6m



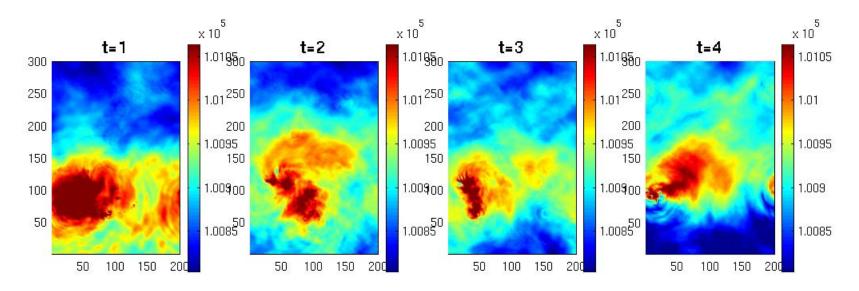


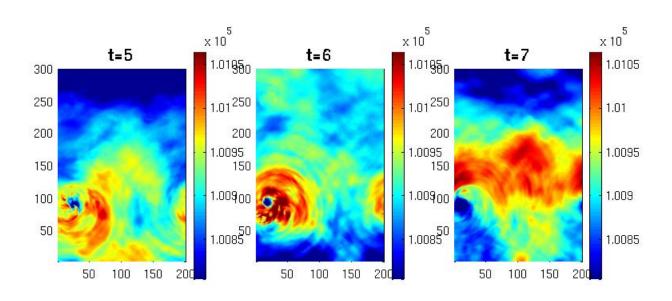
$h^*(J/kg)$ at z=3810m





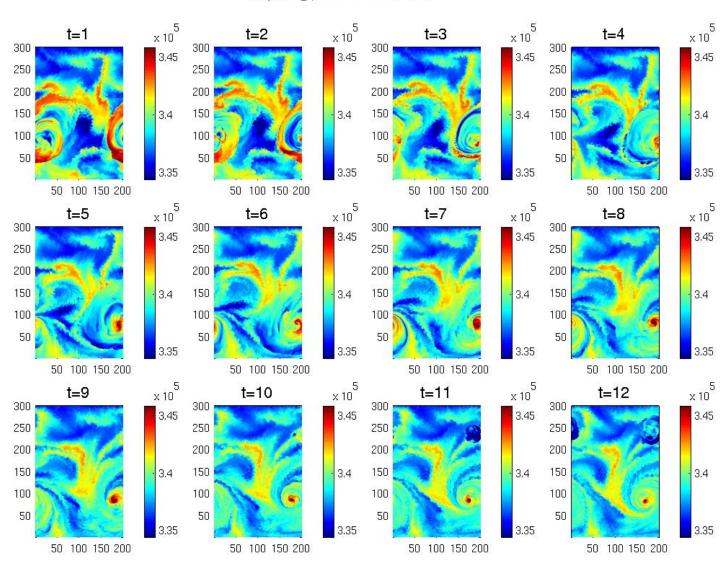
P(Pa) at z=24.6m



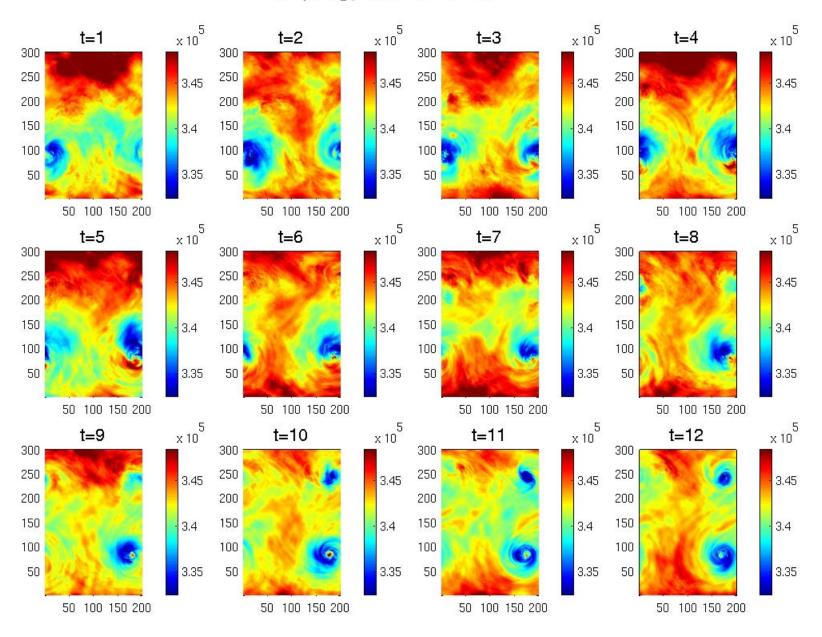


Later times:

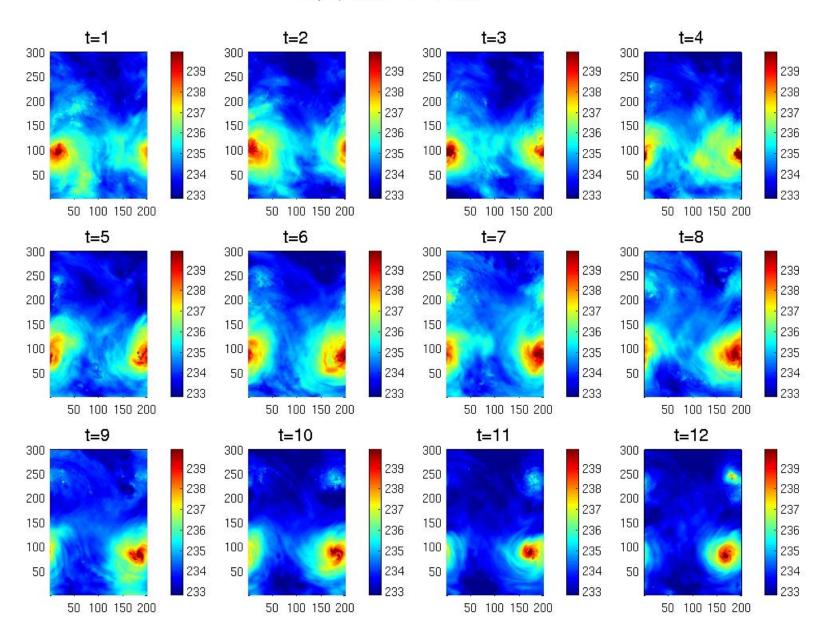
h(J/kg) at z=24.6m



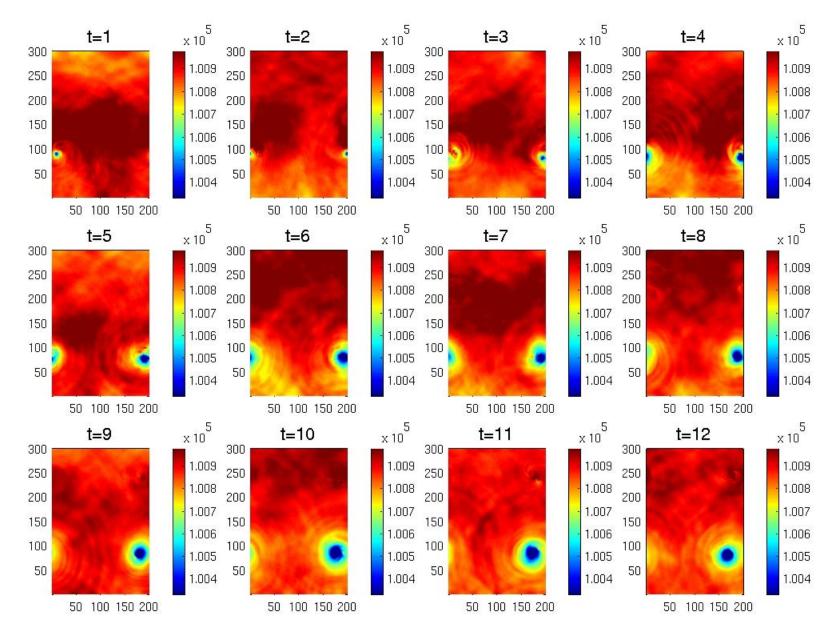
h*(J/kg) at z=3810m



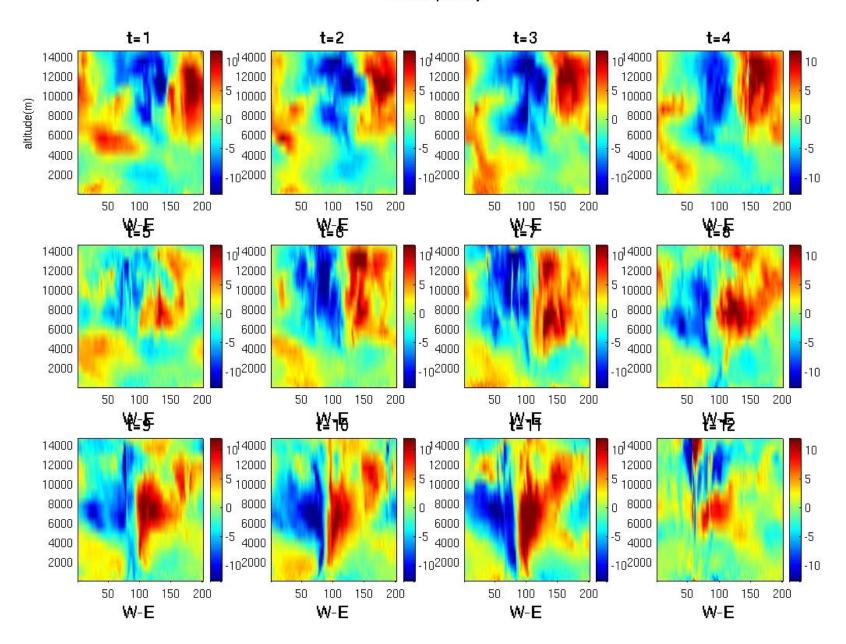
T(K) at z=9675m



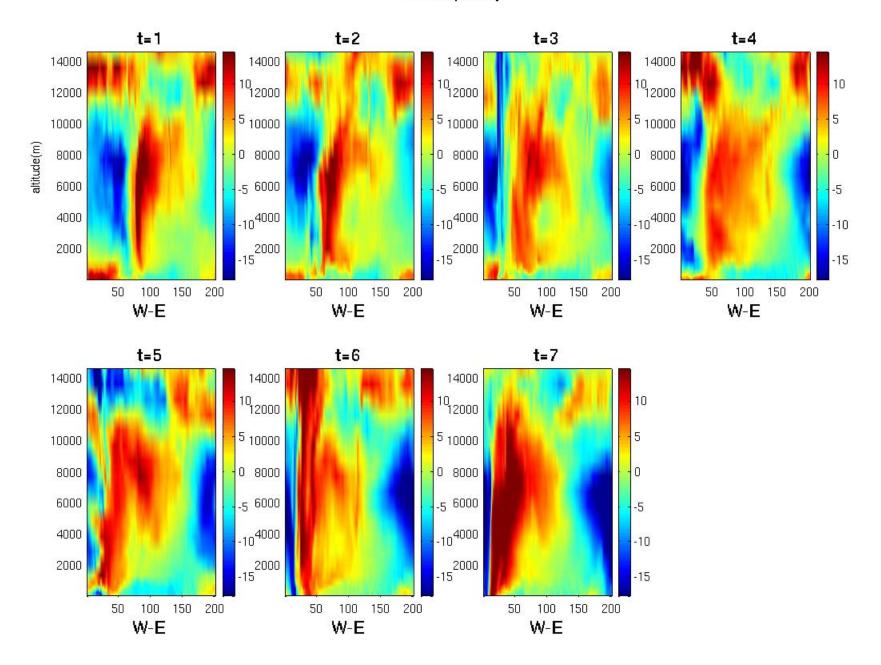
P(Pa) at z=24.6m



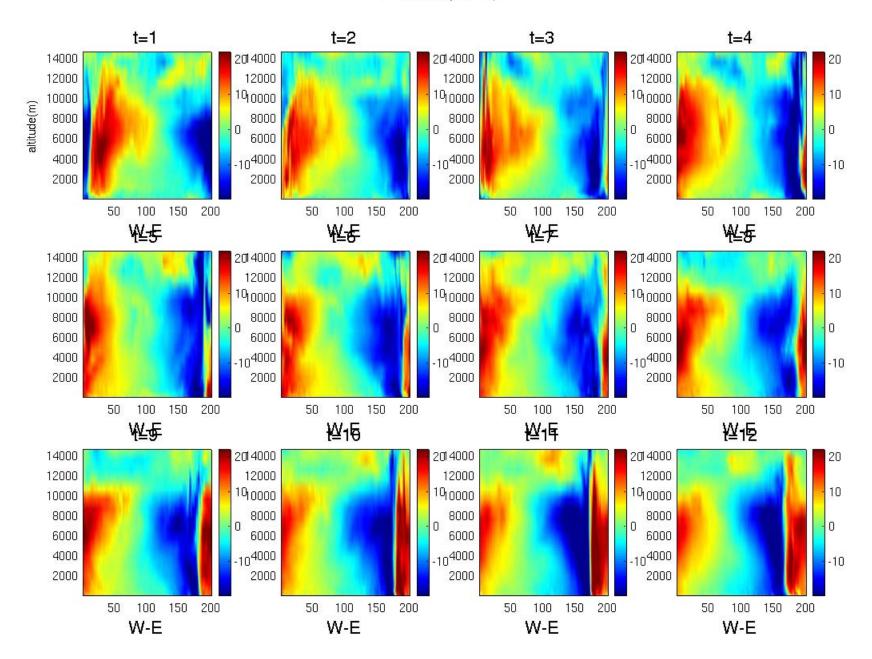
Evolution of v in E-W cross-sections v-wind(m/s)



v-wind(m/s)



v-wind(m/s)



Evolution of RH in E-W cross-sections q/q*

