Intraseasonal Variability in the Tropics

Madden and Julian, 1972



FIG. 2. Individual variance spectra for the 850- and 150-mb zonal wind component and station (sfc) pressure for the Canton Island record. The use of a logarithmic ordinate permits a constant scaling to be used for the chi-square degrees of freedom sampling analysis. This scaling $[\chi^2(0.1\%)/51]$ and the bandwidth of the analysis, $\Delta f = 0.0081$ day⁻¹, are shown by the cross. Spectral densities are normalized to unit bandwidth (m² sec⁻² day⁻¹).

Madden and Julian, 1972



FIG. 2. Variance spectrum for station pressures at Nauru I. $(0^{\circ}24'S, 161^{\circ}0'E)$. Ordinate is logarithmic and abscissa (frequency) is linear. The 40–50 day period range is indicated by the dashed vertical lines. Prior 95% confidence limits and the bandwidth of of the analysis (0.008 day⁻¹) are indicated by the cross.



Fig. 6. Zonal wind oscillation in the equatorial plane at the time when the station pressure is a maximum at Canton based on the phase angles of Table 4. The unit length represents the maximum excursion at each location. The +, -, or 0 at the tail of each wind arrow represents the sign of the instantaneous local change of the zonal wind. Arrows are plotted only at levels whose coherence squares from Table 4 are above their background coherence square, and whose spectra, as tabulated in Table 3, indicate a peak. Heavy arrows at the top and bottom represent a schematic of the upper and lower tropospheric wind disturbance that is consistent with the plotted wind arrows and that will satisfy the local changes if it propagates eastward.





Figure 2. Latitude-lag map of (a) 200 mbar winds (vectors), MSUT (contours), and OLR (shaded) and (b) 1000 mbar winds (vectors) and divergence (contours) regressed onto OLR at 0°N, 125°E. All data are windowed according to significant activity at the equator (Figure 1b) and band-pass filtered to eastward wavenumbers 1-3 with 35 to 95-day periods. The regressed fields are shown for a one standard deviation fluctuation of the reference time series. Maximum vectors are 2.7 ms^{-1} at 200 mbar and 0.7 ms^{-1} at 1000 mbar. Contour interval for MSUT is 5.0×10^{-2} K, and for 1000-mbar divergence it is $0.9 \times 10^{-7} \text{ s}^{-1}$. Shading levels for OLR (positive means wet convection) start at 0.5 K, 1.75 K, 3.0 K, and 4.25 K.



FIG. 16. A descriptive model of the kinematic, thermodynamic, and surface properties of the December to early January westerly wind burst as it passed the IFA. Day 0 is time of maximum low-level westerlies, with earlier times indicated by negative days (placed to the right so that the left portion of the diagram is to the west: see caution in text, however, about fully interpreting diagram as west-east section). Letters in figure refer to anomalies W: warm, C: cool, M: moist, and D: dry. Heavy arrows indicate strong vertical motion; light arrows weak vertical motion. Clouds are schematic, horizontal scales exaggerated. Temperatures corresponding to pressure levels are indicated on right.





FIG. 4. Coherence squared (contoured) and phase (vectors) between OLR and itself at $(0^{\circ}, 84^{\circ}E)$ for ER8 truncation. Coh² > 0.1 significant at the 99% level; see text. Zero phase indicated by upward-directed vector. Clockwise rotation implies direction of phase propagation.



Fig. 1. Composite seasonal cycle of OLR variance for eastward wavenumbers 1-3 and 35-95 day periods. Time series at each latitude is smoothed with 100 day running mean before plotting. Units are arbitrary. Stripes represent times when signal is significant at 95% level relative to a composite background spectrum based on a Chi-square test. Adapted from Salby and Hendon (1994).















Fig. 9. (a) Time-longitude section of the OLR anomalies for the MJO-filtered band for the same 6-month sample period as Fig. 8, averaged for the latitudes from 10°S to 2.5°N. The zero contour



Fig. 9 (*Continued*) (e) The n = 0 EIG wave-filtered plus the s = 0, n = 0 IG wave-filtered band. (f) The n = 1 WIG wave-filtered plus the n = 2 WIG wave-filtered bands.











DAY 0





Perturbations in surface properties and OLR/precip confined to Indian Ocean and western Pacific, but upper tropopsheric wind signals are global

European Center forecast 200 hPa velocity potential, 4 January to 13 February, 2005



CHI 200 hPa 40—DAY forecast (00z04jan2005—13feb2005) (based on EWP zonal harmonics)

Observational analyses courtesy of:

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from Wheeler and Kiladis, 1999

OLR power spectrum, 15°S-15°N, 1979–2001 (Antisymmetric)



from Wheeler and Kiladis, 1999



from Wheeler and Kiladis, 1999



from Wheeler and Kiladis, 1999













from Wheeler and Kiladis, 1999



OLR power spectrum, 1979–2001 (Symmetric)


OBSERVATIONS OF KELVIN WAVES AND THE MJO Time-longitude diagram of CLAUS Tb (5S-equator), February 1987





OBSERVATIONS OF TWO DAY (WIG), KELVIN WAVES AND THE MJO Time-longitude diagram of CLAUS Tb (2.5S-2.5N), October-December 1992



OBSERVATIONS OF TWO DAY (WIG), KELVIN WAVES AND THE MJO Time-longitude diagram of CLAUS Tb (2.5S-2.5N), October-December 1992



OBSERVATIONS OF TWO DAY (WIG), KELVIN WAVES AND THE MJO Time-longitude diagram of CLAUS Tb (2.5S-2.5N), October-December 1992



OBSERVATIONS OF KELVIN WAVES AND THE MJO Time–longitude diagram of CLAUS Tb (2.5S–7.5N), January–April 1987







Kelvin waves (15 m s⁻¹)



1998 OLR 2.5°S-2.5° N



OLR shading starts at - 10 W m⁻² at 20 W m⁻² intervals, negative only

Regression Model

Simple Linear Model:

y = ax + b

where: x= predictor (filtered OLR) y= predictand (OLR, circulation) Convective Fraction from TRMM TMI Regressed against MJO-filtered OLR at Eq, 80°E (scaled -40 W m²) for 1998-2003



OLR (contours, 10 W m⁻²) Convective Fraction (shading, ± 2 and 5%), red positive

Convective Fraction from TRMM TMI Regressed against Kelvin-filtered OLR at 7.5°N, 172.5°W (scaled -40 W m²) for 1998-2003



OLR (contours, 10 W m⁻²) Convective Fraction (shading, ± 2 and 5%), red positive

Convective Fraction from TRMM TMI Regressed against MRG-filtered OLR at 7.5°N, 172.5°W (scaled -40 W m²) for 1998-2003



OLR (contours, 10 W m⁻²) Convective Fraction (shading, ± 2 and 5%), red positive

Morphology of a Tropical Mesoscale Convective Complex in the eastern Atlantic during GATE (from Zipser et al. 1981)



System Motion Is Left to Right at 3 m s⁻¹. Arrows Show Relative Wind.



Observed Kelvin wave morphology (from Straub and Kiladis 2003)



Morphology of MJO (from Wang 2005)



Equatorial Wave Structure

Consistent with a progression of shallow to deep convection, followed by stratiform precipitation for the Kelvin, Westward Inertio-gravity (2-day) Waves, and Easterly Waves

This was also observed during COARE for the MJO (e.g. Lin and Johnson 1996; Johnson et al. 1999; Lin et al. 2004)

This evolution is similar to that occurring on the Mesoscale Convective Complex scale



ECMWF reanalysis regression

OLR (shading), 1000-hPa Z (contours), 1000-hPa winds (vectors)



Dynamical fields are symmetric with respect to the equator; convection is centered in NH along ITCZ

a) OLR



b) Temperature



c) Specific humidity



Kelvin wave in Majuro radiosonde data: 1979-1999 (Straub and Kiladis, 2003)





Zonal Wind Anomaly over the IFA

Two day waves during COARE, from Haertel and Kiladis, 2004



Temperature Anomaly over the IFA



Moisture Anomaly over the IFA



Heating Anomaly over the IFA







Day-12 Streamfunction (contours 4 X $10^5 \text{ m}^2 \text{ s}^{-1}$)

Wind (vectors, largest around 2 m s⁻¹) OLR (shading starts at +/- 6 W s⁻²), negative blue



Day-8 Streamfunction (contours 4 X 10⁵ m² s⁻¹) Wind (vectors largest around 2 m s⁻¹)

Wind (vectors, largest around 2 m s⁻¹) OLR (shading starts at +/- 6 W s⁻²), negative blue



Day-4



Day 0



Day+4



Day+8



Day+12

Zonal Wind at Honiara (10°S, 160°E) Regressed against MJO-filtered OLR (scaled -40 W m²) for 1979-1999



Zonal Wind at Seychelles (5°S, 55°E) Regressed against MJO-filtered OLR (scaled -40 W m²) for 1979-1999



Zonal Wind at Diego Garcia (7.5°S, 72°E) Regressed against MJOfiltered OLR (scaled -40 W m²) for 1979-1999



Temperature at Tarawa (Eq, 172.5°E) Regressed against MJO-filtered OLR (scaled -40 W m²) for 1979-1999


Temperature at Honiara (10°S, 160.0°E) Regressed against MJO-filtered OLR (scaled -40 W m²) for 1979-1999



Temperature at Diego Garcia (7.5°S, 72°E) Regressed against MJOfiltered OLR (scaled -40 W m²) for 1979-1999



Temperature at Seychelles (5°S, 55°E) Regressed against MJO-filtered OLR (scaled -40 W m²) for 1979-1999



Specific Humidity at Diego Garcia (7.5°S, 72°E) Regressed against MJO-filtered OLR (scaled -40 W m²) for 1979-1999



Specific Humidity at Medan (2.5°N, 97.5°E) Regressed against MJOfiltered OLR (scaled -40 W m²) for 1979-1999



Specific Humidity at Tarawa (Eq, 172.5°E) Regressed against MJOfiltered OLR (scaled -40 W m²) for 1979-1999



Specific Humidity at Truk (7.5°N, 152.5°E) Regressed against MJOfiltered OLR (scaled -40 W m²) for 1979-1999



Q1 Regressed against MJO-filtered OLR over the IFA during COARE

a) Q1 Total



Q1 Regressed against MJO-filtered OLR over the IFA during COARE

b) Q1 First Mode (49 m/s)



Q1 Regressed against MJO-filtered OLR over the IFA during COARE

c) Q1 Second Mode (23 m/s)



Dynamical Structures

All equatorial waves studied have tilted vertical structures, with:

Easterlies ahead of and westerlies following the convective region

Warm lower tropospheric temperatures ahead of the wave, with cooling behind. The mid-troposphere is warm within the convective region, indicating that latent heating more than compensates for vertical motion.

Waves are moist ahead (high CAPE) and dry following the deep convection

AVHRR OLR











ECHAM4_OPYC3 OLR



1.2

1.4

1.6

1.8



















Conclusions

Although the MJO is comprised of a variety of higher frequency, smaller scale disturbances, the dynamical structures of all of these waves resemble each other in many important aspects, all consistent with shallow cumulus leading to deep convection followed by stratiform precipitation

There is a high degree of self-similar behavior seen in equatorial waves across a wide variety of scales

General Circulation Models do not represent such scale interactions, and most do not adequately represent the MJO or other equatorial modes sufficiently well.



Figure 3. As for Figure 1, but for u at 940 mb, wavenumber 1. The vertical scale is one quarter that of Figure 1.

Neelin et al., JAS, 1987



Composite Evolution of 200-hPa Velocity Potential Anomalies (10⁶m²s⁻¹) and points of origin of tropical systems that developed into hurricanes / typhoons



Tropical cyclone activity and the MJO